



Temporal and spatial variations of the atmospheric CO₂ concentration in China

Dongqi Zhang,¹ Jie Tang,¹ Guangyu Shi,² Takakiyo Nakazawa,³ Shuji Aoki,³ Satoshi Sugawara,⁴ Min Wen,¹ Shinji Morimoto,⁵ Prabir K. Patra,⁶ Tadahiro Hayasaka,⁷ and Tazu Saeki⁷

Received 31 October 2007; accepted 6 December 2007; published 1 February 2008.

[1] In order to understand the variations of atmospheric greenhouse gases over the Chinese mainland, an air sampling network was established in March 2003 that included one GAW global station, three GAW regional stations and three cooperative stations. Flask air samples were taken every week at the stations for 3 years. The secular increase and the seasonal variation in the CO₂ concentration are clearly observable at all the stations, reflecting global as well as the regional terrestrial biospheric and human activities, as well as the local meteorological conditions. The average CO₂ concentration depends on the station, with the lowest value observed at Mt. Waliguan and relatively high values observed at Fukang and Lin-an stations, located in the westernmost and eastern parts of China, respectively. The observed CO₂ concentration variations are also discussed within the context of forward transport model simulations using CO₂ fluxes derived by time-dependent inversion. **Citation:** Zhang, D., et al. (2008), Temporal and spatial variations of the atmospheric CO₂ concentration in China, *Geophys. Res. Lett.*, 35, L03801, doi:10.1029/2007GL032531.

1. Introduction

[2] Precise measurements of the atmospheric CO₂ concentration over a wide geographical area are indispensable for quantifying its sources and sinks on the earth's surface [Tans et al., 1989, 1990; Fan et al., 1998; Bousquet et al., 2000; Patra et al., 2005a, 2005b]. The atmospheric CO₂ concentration has been monitored systematically at many places around the world since the 1950's [Keeling et al., 1976a, 1976b; Conway et al., 1994]. However, the current global CO₂ network is heavily biased toward ocean area. For a better understanding of the global carbon cycle, it is necessary to increase the number of CO₂ monitoring sites in the interior of continents. Indeed, additional inland stations

have recently been established in Europe (AEROCARB, <http://www.ipsl.jussieu.fr/~mrsce/index.html>) and North America (The North American Carbon Program Plan, <http://www.nacarbon.org/nacp/index.html>). In China, flask air sampling was first started by the Chinese Academy of Meteorological Sciences (CAMS) at Mt. Waliguan (one of the 24 World Meteorological Organization/Global Atmosphere Watch (WMO/GAW) global baseline stations) in May 1990, in cooperation with the National Oceanic and Atmospheric Administration/Climate Monitoring and Diagnostics Laboratory (NOAA/CMDL), and has continued to the present [Tang et al., 1995; Zhou et al., 2003]. To reveal temporal and spatial variations of atmospheric CO₂ in the Chinese mainland in more detail, the CAMS, the Institute of Atmospheric Physics, Chinese Academy of Sciences and the Center for Atmospheric and Oceanic Studies, Tohoku University, Japan established a network of 7 monitoring sites that includes the Mt. Waliguan station, 3 GAW regional stations and 3 cooperative stations, and started CO₂ flask sampling in March 2003.

[3] In this paper, we describe our CO₂ concentration measurements in China and discuss the results obtained during the first 3 years of the program, especially in terms of seasonal cycle and annual average value of the CO₂ concentration. The observed CO₂ variations are compared with those simulated using an atmospheric general circulation model and 22-region time-dependent inversion CO₂ fluxes.

2. Experimental Procedures

[4] Our sampling network consists of 7 sites, as shown in Figure 1. Mt. Waliguan station (36.3°N, 100.9°E, 3810 m) is situated on the northeastern edge of the Tibetan Plateau. The surrounding area is arid/semi-arid grassland and desert steppe, and there are no industrial or pollution source areas in a radius of 50 km. Therefore, the Waliguan station is a suitable site for monitoring the background CO₂ concentration in the Eurasian Continent. Three stations in Longfengshan, Shangdianzi and Lin-an are approved as WMO/GAW regional stations. The Longfengshan station (127.6°E, 44.7°N, 310 m) is located in the northeastern part of China, with temperate forests and grasslands dominating its surrounding area. The Shangdianzi station (117.1°E, 40.4°N, 293.9 m) is 150 km northeast of Beijing, and orchard and farmland surround the station. The Lin-an station (119.1°E, 30.3°N, 138 m) is situated near the Yangtze River delta, with the surrounding area characterized as subtropical forests and farmland. Two of the 3 cooperative stations are the meteorological observatory at the summit of

¹Key Laboratory for Atmospheric Chemistry, Centre for Atmosphere Watch and Services, Chinese Academy of Meteorological Sciences, Beijing, China.

²Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing, China.

³Center for Atmospheric and Oceanic Studies, Tohoku University, Sendai, Japan.

⁴Institute of Earth Science, Miyagi University of Education, Sendai, Japan.

⁵National Institute of Polar Research, Tokyo, Japan.

⁶Frontier Research Center for Global Change, Yokohama, Japan.

⁷Research Institute for Humanity and Nature, Kyoto, Japan.

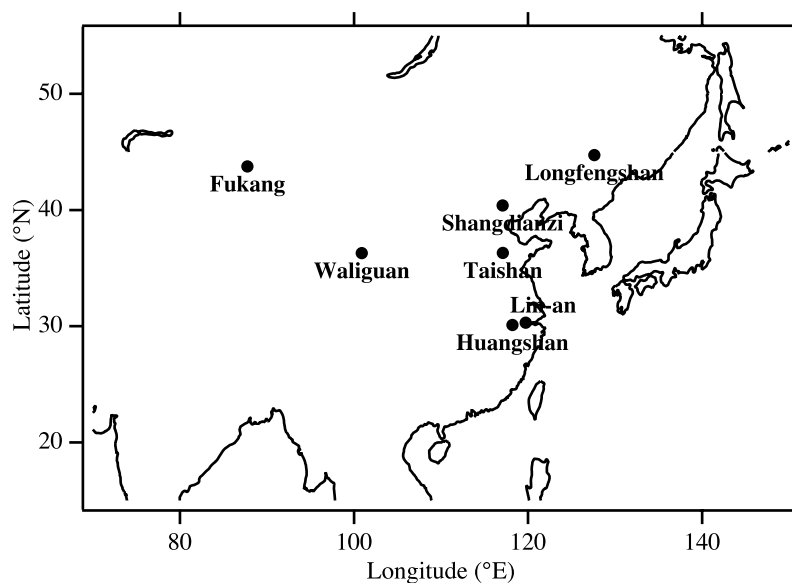


Figure 1. Map showing 7 sampling sites in China.

Mt. Taishan (117.1°E, 36.3°N, 1534 m) and Mt. Huangshan (118.2°E, 30.1°N, 1841 m). The remaining station is the observatory of the Chinese Academy of Sciences in Fukang (87.9°E, 44.3°N, 630 m) in the westernmost part of China, and the typical land surface around the station is oasis and desert.

[5] To collect air samples once per week at the stations, we have used two sets of instruments, one is 2.5 L Pyrex glass flasks and a sampling unit [Thoning *et al.*, 1995] from the NOAA/Earth System Research Laboratory (ESRL) and the other is 1.0 L stainless steel flasks and a sampling unit from Tohoku University. Both sampling units are equipped with a diaphragm pump to pressurize ambient air into the flask. The NOAA instrument was used at Waliguan, Longfengshan and Lin-an, with sample inlet located 5 m above the ground; each flask was pressurized to about 1.3×10^3 hPa. The Tohoku University instrument was used at Shangdianzi, Taishan, Huangshan and Fukang, and air samples collected at 10 m above the ground were pressurized into flasks at about 2×10^3 hPa. The glass and stainless steel flasks were rinsed by flowing ambient atmosphere through them at a rate of $5\text{--}9 \text{ L min}^{-1}$ for about 15 min and at a rate of 10 L min^{-1} for about 3 min, respectively, to ensure a complete air exchange in the flasks. The collection of air samples was usually made early in the afternoon to obtain fully mixed air. The air sampling at each surface site was conducted as much as possible when winds were from areas not influenced by local human infrastructures and heavy vegetation.

[6] The collected air samples from the stations were sent to the Key Laboratory for Atmospheric Chemistry, CAMS in Beijing, and their CO₂ concentrations were determined with a precision of 0.1 ppm against our working standard gases using a non-dispersive infra-red gas analyzer (LI-COR 7000) within 1 month of sampling. To minimize the influence of water vapor on the measured CO₂ concentration value, each air sample was dried by passing it through a water vapor trap cooled to -50°C before introducing it into the analyzer. The working standard gases were CO₂-air mixtures stored in 30 L aluminum high-pressure

cylinders. Their CO₂ concentrations were calibrated against the WMO-95 scale using 5 secondary standard gases with CO₂ concentrations of 348.09, 366.60, 379.09, 400.89 and 419.98 ppm which were purchased from the NOAA/ESRL in 2004. Each air sample was analyzed twice and their average was regarded as the sample CO₂ concentration.

[7] To extract the seasonal cycle, average value and increase rate of the CO₂ concentration, a digital filtering technique including Fourier harmonics, Reinsch-type cubic spline and Butterworth filter was applied to the CO₂ time series. Details of the technique have been described by Nakazawa *et al.* [1997]. In this study, the average seasonal CO₂ cycle was approximated by fundamental annual cycle and its first harmonics, and the cut-off period of the 26th-order Butterworth filter was chosen to be 24 months to determine the secular CO₂ trend. After the best-fit curve to the observed CO₂ values was obtained, the data lying more than 3 standard deviations from the curve were rejected from the dataset as outliers. The digital filtering was then conducted on the remaining data. These steps were repeated once or twice until no further outliers were found. The procedure rejected 2 to 8 data points as outliers for each site, corresponding to 1 to 6 % of the total data. The standard deviation of selected data from the best-fit curve (1σ) for each site is shown in Table 1.

3. Results and Discussion

[8] The CO₂ concentrations observed at the 7 sites are shown in Figure 2. It is obvious from Figure 2 that the CO₂ concentration increases with time at all sites, accompanying by clear seasonal cycle. The average CO₂ increase rates derived from the best-fit curves for the respective sites range between 3.6 and 1.7 ppm yr⁻¹, as summarized in Table 1. The different CO₂ increase rates observed in this study can be attributed to different influences of regional CO₂ sources and sinks, air transport and local meteorology at individual stations. In this regard, it is worthy of note that the emission of fossil fuel CO₂ in China has been increasing significantly in recent years [Marland *et al.*, 2005] due to rapid economic

Table 1. Average Increase Rates, Average Values and Seasonal Cycles of the CO₂ Concentration, and Standard Deviations of the Observed Data From the Best-Fit Curve at 7 Sites in China

Site	Observation Periods	Increase Rate, ppm yr ⁻¹	Average Conc., ppm	Seasonal Maximum, ppm	Appearance of Seasonal Maximum, day	Seasonal Minimum, ppm	Appearance of Seasonal Minimum, day	Peak-to-Peak Amplitude, ppm	Standard Deviation, ppm
Fukang	Apr. 2003–Apr. 2006	2.4	385.3	12.8	18	−12.0	222	24.8	3.4
Huangshan	Mar. 2003–Mar. 2006	2.3	380.3	6.2	57	−7.1	236	13.3	3.2
Lin-an	Jun. 2003–Jun. 2006	3.0	390.3	6.0	22	−5.5	235	11.5	8.1
Longfengshan	Jun. 2003–Jun. 2006	3.6	383.8	8.0	25	−11.3	207	19.3	4.1
Shangdianzi	Jul. 2003–May 2006	2.7	383.8	6.1	62	−9.9	214	16.0	6.4
Taishan	Apr. 2003–Apr. 2006	1.7	382.0	8.0	42	−10.3	226	18.3	5.3
Waliguan	Jun. 2003–Apr. 2006	1.9	378.4	5.1	86	−5.8	227	10.9	1.7

growth, resulting in the high CO₂ increase rates observed at our surface sites especially in eastern China.

[9] Characteristics of the average seasonal CO₂ cycle at individual stations are given in Table 1. The CO₂ concentration usually reaches a maximum in February and a minimum in August. However, the timing of the occurrences of seasonal maximum and minimum varies significantly from station to station, ranging from a maximum in January at Fukang to a maximum at the end of March or in early April at Waliguan, while the seasonal minimum at Longfengshan is found in July and at Shangdianzi at the end of July or in early August. These differences are produced not only by regionally different terrestrial biospheric and human activities but also by local meteorological conditions. The high CO₂ concentrations observed at Fukang in the wintertime is likely due to a local strong surface inversion that traps CO₂ emitted by human and terrestrial biospheric activities [Li, 1991]. A long-range air transport is important for the appearance time of the maximum CO₂ concentration at Waliguan, since its timing agrees well with those observed at many background monitoring sites at similar latitudes [e.g., Nakazawa *et al.*, 1993]. As seen in Figure 2, no summertime low CO₂ concentrations were observed at Lin-an in 2003. We speculate that this resulted from high CO₂ emitted by local electric power stations operating at maximum capacity during the hot/dry summer of 2003 to match the power shortage caused by the suspending of originally dominant hydro-power stations in Zhejiang Province where Lin-an station is located. It is interesting to note that the same region was struck by less intense heat waves in the summer of 2005, resulting in higher CO₂ concentration than that observed during the summer of 2004.

[10] The largest seasonal CO₂ cycle with a peak-to-peak amplitude of 24.8 ppm is found at Fukang, mainly due to the wintertime CO₂ accumulation in the surface inversion layer. On the other hand, Waliguan shows the smallest amplitude of 10.9 ppm, since it is a high mountain site. Lin-an also shows a small seasonal amplitude of 11.5 ppm, mainly due to the high summertime CO₂ concentrations especially in 2003 and 2005. The small seasonal cycle observed at Lin-an is also due to the fact that the station is located in the southern part of our network where seasonal variation in the biospheric flux is smaller. The seasonal amplitudes at the remaining stations range between 13.3 and 19.3 ppm, which are in agreement with those at other continental sites at similar latitudes. For example, the amplitude of 19.3 ppm at Longfengshan (44.7°N) is similar to those at Kasproy Wierch, Poland (19.1°E, 49.2°N,

1987 m) [Necki *et al.*, 2003] and Fraserdale, Canada (81.6°W, 49.8°N) [Higuchi *et al.*, 2003]. However, the seasonal CO₂ cycles at Longfengshan and Taishan (36.2°N, 1534 m) are larger than those observed at similar latitudes in Japan located downwind of China; the respective amplitudes are larger by 1.6 and 4.9 ppm than those near Wakkanai (141.5°E, 45.5°N) (T. Nakazawa, Tohoku University, unpublished data, 2007) and at 0–2 km over Sendai (141.8°E, 37.7°N) [Nakazawa *et al.*, 1993] for this analysis period.

[11] The average CO₂ concentrations at the 7 sites are also listed in Table 1. Lin-an shows the highest value of 390.3 ppm, reflecting the fact that the summertime CO₂ concentrations were extremely high especially in 2003 and 2005 due to local contamination. The value at Fukang is also high, i.e. 385.3 ppm, mainly due to the wintertime accumulation of anthropogenic CO₂ near the ground surface. The average CO₂ concentrations at the mountain stations are lower than those at the surface stations, since the mountain sites are remote from areas with strong human activities. For example, Waliguan station shows the lowest CO₂ concentration of 378.4 ppm and the value at Huangshan is found to be 380.3 ppm. However, the average CO₂ value at Taishan is higher by 2–4 ppm than those at the other 2 mountain stations. In this connection, it may be noted that Mt. Taishan is situated in the North China Plain with a high population and intense industrialization. The NOAA/ESRL has also measured the CO₂ concentration at Waliguan (ftp://ftp.cmdl.noaa.gov/ccg/co2/flask/month/wlg_01D0_mm.co2). By comparing with their monthly mean values, we found that our results agree fairly well with theirs, except for summer months when our values are clearly lower by a few ppm. As a result, our average CO₂ concentration over the present observation period is lower by 0.7 ppm than theirs. The summer difference with the NOAA/ESRL could be due to sampling time; their samples were collected at around 8 a.m., while our sampling was made in the early afternoon. Continuous CO₂ measurements at Mt. Waliguan show a clear diurnal cycle in summer [Zhou *et al.*, 2003], probably due to diurnal effects of local biospheric activities and/or upslope/downslope winds.

[12] To understand how CO₂ sources and sinks contribute to the spatially different variations of the CO₂ concentration in China, we employ an atmospheric general circulation model (AGCM [e.g., Takigawa *et al.*, 2005]) and one-year monthly-mean CO₂ fluxes derived from a 22-region time-dependent inversion. This set of flux is estimated following Gurney *et al.* [2004], but using atmospheric CO₂ data [GLOBALVIEW-CO₂, 2005] from a network of 87 stations

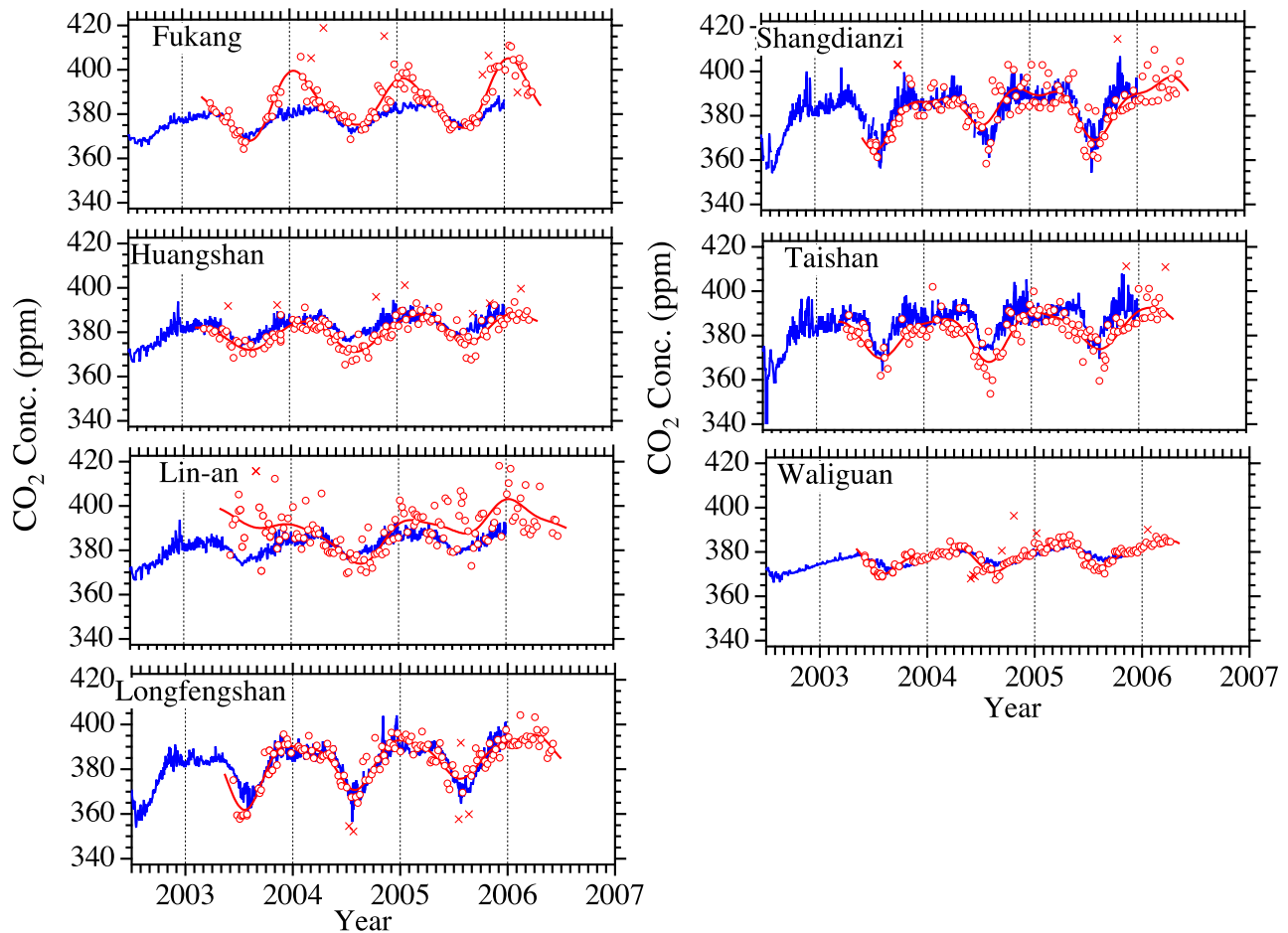


Figure 2. Temporal variations of the CO₂ concentration at 7 sites in China. Circles, solid lines, and crosses show the observed data, their best-fit curve, and outliers rejected in the curve fitting, respectively. The CO₂ concentration variations simulated using the atmospheric general circulation model and 22-region time-dependent inversion CO₂ fluxes are also shown by blue lines.

for the period 1999–2001 (see *Patra et al.* [2005a] for site locations). An average of inverse fluxes calculated by 12 atmospheric transport models is used in our AGCM simulation. The CO₂ concentrations thus calculated are also plotted in Figure 2 for each of the stations. In Figure 2, all the calculated concentration values are shifted so that the annual average of the calculated CO₂ concentration for 2005 at Waliguan agrees with the observed value. Because the model spin-up run was performed for a short period (5 years), an offset is required to match the observed concentration, and we selected Waliguan for the offset determination as CO₂ data from this site was utilized in the inverse calculation. It is seen that the CO₂ concentration variations calculated for Waliguan, Shangdianzi, Longfengshan and Huangshan are fairly consistent with the observed results, while the disagreement between the calculated and the observed CO₂ variations is clearly found at Fukang, Lin-an and Taishan. Especially at Fukang in the winter season, the calculated CO₂ concentrations are much lower than the observed values, suggesting that the model cannot reproduce the local effect of the anthropogenic CO₂ accumulation in the strong surface inversion layer in very cold

season. The CO₂ concentrations calculated for Lin-an are also lower than the observed values, except for 2004; this is mainly due to neglecting the effect of additional CO₂ emissions that occurred in the summer of 2003 and 2005. On the other hand, the calculated CO₂ concentrations at Taishan are clearly higher than the observed values, especially in the warm season. A similar phenomenon is also seen at Huangshan, although the concentration differences are smaller. This discrepancy may suggest that the summertime upward air transport in the model is weaker than what actually takes place in the real atmosphere. Note here that the inverse model flux regions are too coarse to capture regional emission distributions, no interannual variability in fluxes is considered and in 22-region inverse model, China, Mongolia and two Koreas form one region.

4. Conclusions

[13] By systematically measuring the CO₂ concentration at 7 sites in China for almost 3 years from March 2003, we have presented its temporal and spatial characteristics over the Chinese mainland. The observed CO₂ variations were

reproduced reasonably well using the AGCM with 22-region time-dependent inversion CO₂ fluxes, although significant differences between the calculated and the observed CO₂ concentrations were found at stations with strong local effects on the CO₂ concentration. The atmospheric CO₂ concentration data obtained in this study would be useful for constraining CO₂ sources and sinks in China at higher spatial resolutions. Given the fact that China is emerging as a significant contributor to future atmospheric CO₂ levels, it is important to expand the observation network, bind it interactively closer together and maintain it for a long time, which will benefit future/high-resolution inverse modeling efforts.

[14] **Acknowledgments.** We express our sincere thanks to all the staffs who contributed to the collection of air samples at Waliguan, Shangdianzi, Longfengshan, Lin-an, Taishan, Huangshang and Fukang stations. This work was supported by the China Ministry of Science and Technology project (2001DIA10009), the research project 2-1 of the Research Institute for Humanity and Nature, Japan, and the Research Revolution 2002 (project 3) and the Grants-in-Aid for Creative Scientific Research (2005/17GS0203) of the Ministry of Education, Science, Sports and Culture, Japan.

References

- Bousquet, P., P. Peylin, P. Ciais, C. L. Quéré, P. Friedlingstein, and P. P. Tans (2000), Regional changes in carbon dioxide fluxes of land and ocean since 1980, *Science*, *290*, 1342–1346.
- Conway, T. J., P. P. Tans, L. S. Waterman, K. W. Thoning, D. R. Kitzis, K. A. Masarie, and N. Zhang (1994), Evidence for international variability of the carbon cycle from the NOAA/CMDL Global Air Sampling Network, *J. Geophys. Res.*, *99*, 22,831–22,855.
- Fan, S., M. Gloor, J. Mahlman, S. Pacala, J. Sarmiento, T. Takahashi, and P. Tans (1998), A large terrestrial carbon sink in North America implied by atmospheric and oceanic carbon dioxide data and models, *Science*, *282*, 442–446.
- GLOBALVIEW-CO₂ (2005), Cooperative Atmospheric Data Integration Project—Carbon Dioxide [CD-ROM], Global Monit. Div., Earth Syst. Res. Lab., NOAA, Boulder, Colo. (Available at ftp://ftp.cmdl.noaa.gov/ccg/co2/GLOBALVIEW/)
- Gurney, K. R., et al. (2004), Transcom 3 inversion intercomparison: Model mean results for the estimation of seasonal carbon sources and sinks, *Global Biogeochem. Cycles*, *18*, GB1010, doi:10.1029/2003GB002111.
- Higuchi, K., D. Worthy, D. Chan, and A. Shashkov (2003), Regional source/sink impact on the diurnal, seasonal and inter-annual variations in atmospheric CO₂ at a boreal forest site in Canada, *Tellus, Ser. B*, *55*, 115–125.
- Keeling, C. D., J. A. Adams, C. A. Ekdahl, and P. R. Guenther (1976a), Atmospheric carbon dioxide variations at the South Pole, *Tellus*, *28*, 552–564.
- Keeling, C. D., R. B. Bacastow, A. E. Bainbridge, C. A. Ekdahl Jr., P. R. Guenther, and L. S. Waterman (1976b), Atmospheric carbon dioxide variations at Mauna Loa Observatory, Hawaii, *Tellus*, *28*, 538–551.
- Li, J. F. (1991), *The Climate of Xinjiang*, Meteorol. Press, Beijing, China.
- Marland, G., T. A. Boden, and R. J. Andres (2005), Global, regional, and national CO₂ emissions, in *Trends: A Compendium of Data on Global Change* [electronic], Carbon Dioxide Inf. Anal. Cent., Oak Ridge Natl. Lab., U.S. Dep. of Energy, Oak Ridge, Tenn. (Available at http://cdiac.ornl.gov/trends/emis/meth_reg.htm)
- Nakazawa, T., S. Morimoto, S. Aoki, and M. Tanaka (1993), Time and space variations of the carbon isotopic ratio of tropospheric carbon dioxide over Japan, *Tellus, Ser. B*, *45*, 258–274.
- Nakazawa, T., M. Ishizawa, K. Higuchi, and B. A. N. Trivett (1997), Two curve fitting methods applied to CO₂ flask data, *Environmetrics*, *8*, 197–218.
- Neeki, J., M. Schmidt, K. Rozanski, M. Zimnoch, A. Korus, J. Laso, R. Graul, and I. Levin (2003), Six-year record of atmospheric carbon dioxide and methane at a high-altitude mountain site in Poland, *Tellus, Ser. B*, *55*, 94–104.
- Patra, P. K., S. Maksyutov, M. Ishizawa, T. Nakazawa, T. Takahashi, and J. Ukita (2005a), Interannual and decadal changes in the sea-air CO₂ flux from atmospheric CO₂ inverse modeling, *Global Biogeochem. Cycles*, *19*, GB4013, doi:10.1029/2004GB002257.
- Patra, P. K., M. Ishizawa, S. Maksyutov, T. Nakazawa, and G. Inoue (2005b), Role of biomass burning and climate anomalies for land-atmosphere carbon fluxes based on inverse modeling of atmospheric CO₂, *Global Biogeochem. Cycles*, *19*, GB3005, doi:10.1029/2004GB002258.
- Tagigawa, M., K. Sudo, H. Akimoto, K. Kita, N. Takegawa, Y. Kondo, and M. Takahashi (2005), Estimation of the contribution of intercontinental transport during the PEACE campaign by using a global model, *J. Geophys. Res.*, *110*, D21313, doi:10.1029/2005JD006226.
- Tang, J., Y. P. Wen, X. B. Xu, X. D. Zheng, X. Guo, and Y. C. Zhao (1995), China global atmosphere watch baseline observatory and its measurement program, in *Chinese Academy of Meteorological Sciences Annual Report 1994–95*, pp. 56–65, Chin. Meteorol. Press, Beijing.
- Tans, P. P., T. J. Conway, and T. Nakazawa (1989), Latitudinal distribution of the sources and sinks of atmospheric carbon dioxide from surface observations and an atmospheric transport model, *J. Geophys. Res.*, *94*, 5151–5172.
- Tans, P. P., I. Y. Fung, and T. Takahashi (1990), Observational constraints on the global atmospheric CO₂ budget, *Science*, *247*, 1431–1438.
- Thoning, K. W., T. J. Conway, N. Zhang, and D. Kitzis (1995), Analysis system for measurement of CO₂ mixing ratios in flask air samples, *J. Atmos. Oceanic Technol.*, *12*, 1349–1356.
- Zhou, L. X., J. Tang, Y. P. Wen, J. L. Li, P. Yan, and X. C. Zhang (2003), The impact of local winds and long-range transport on the continuous carbon dioxide record at Mount Waliguan, China, *Tellus, Ser. B*, *55*, 145–158.
- S. Aoki and T. Nakazawa, Center for Atmospheric and Oceanic Studies, Tohoku University, Sendai 980-8578, Japan.
- T. Hayasaka and T. Saeki, Research Institute for Humanity and Nature, Kyoto 603-8047, Japan.
- S. Morimoto, National Institute of Polar Research, Tokyo 173-8515, Japan.
- P. K. Patra, Frontier Research Center for Global Change, Yokohama 236-0001, Japan.
- G. Shi, Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing 100029, China.
- S. Sugawara, Institute of Earth Science, Miyagi University of Education, Sendai 980-0845, Japan.
- J. Tang, M. Wen, and D. Zhang, Key Laboratory for Atmospheric Chemistry, Centre for Atmosphere Watch and Services, Chinese Academy of Meteorological Sciences, Beijing 100081, China. (tangj@cams.cma.gov.cn)