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Equatorial Upwelling in the Central Indian Ocean Estimated from Moored ADCP Array

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Outline

- **Part I:**

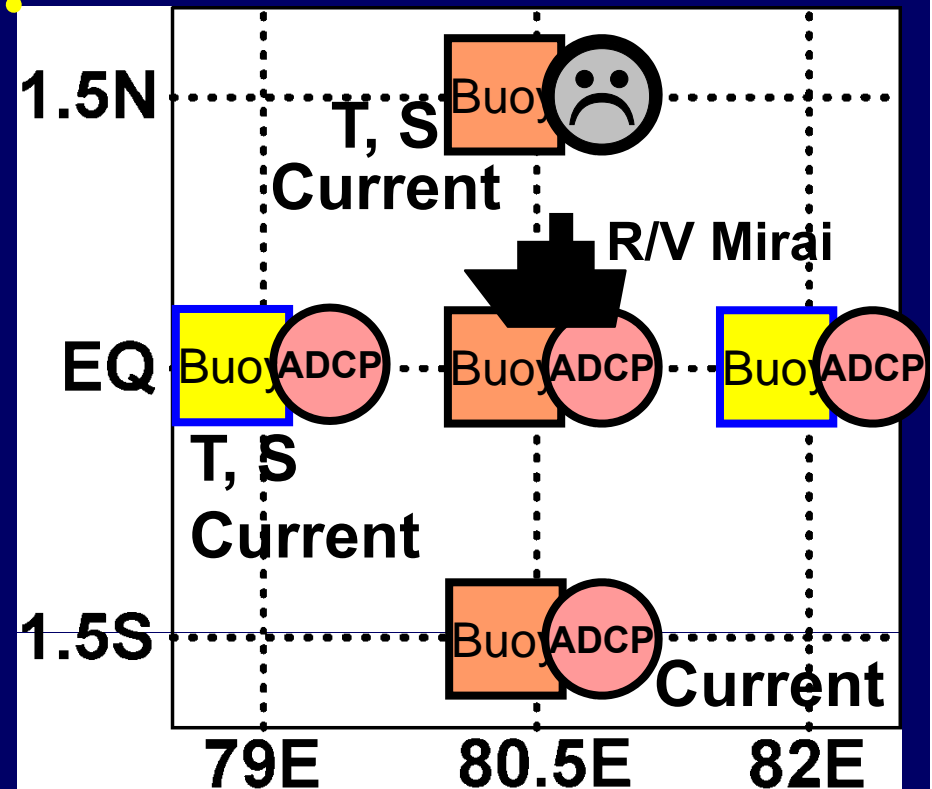
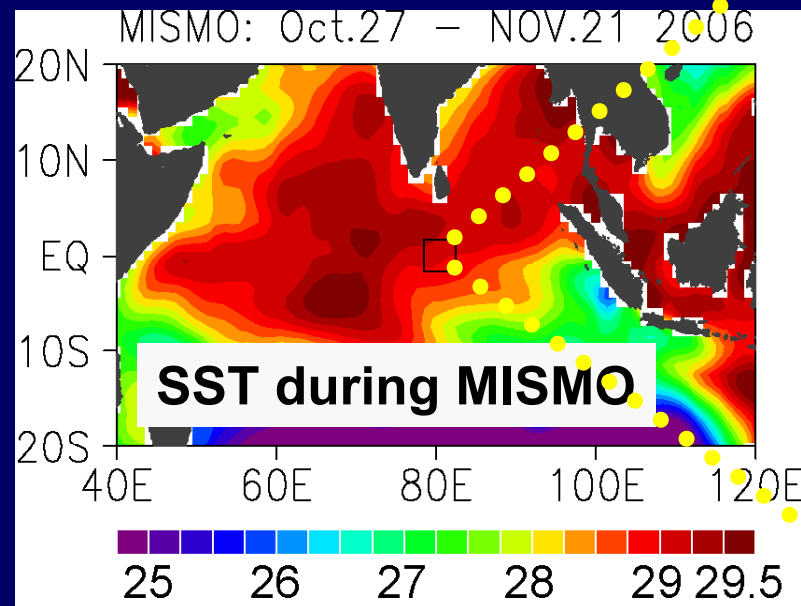
- **Introduction of Ocean Observation during MISMO**
- **A Review of Ocean Condition Observed in MISMO in Oct-Nov 2006**

- **Part II:**

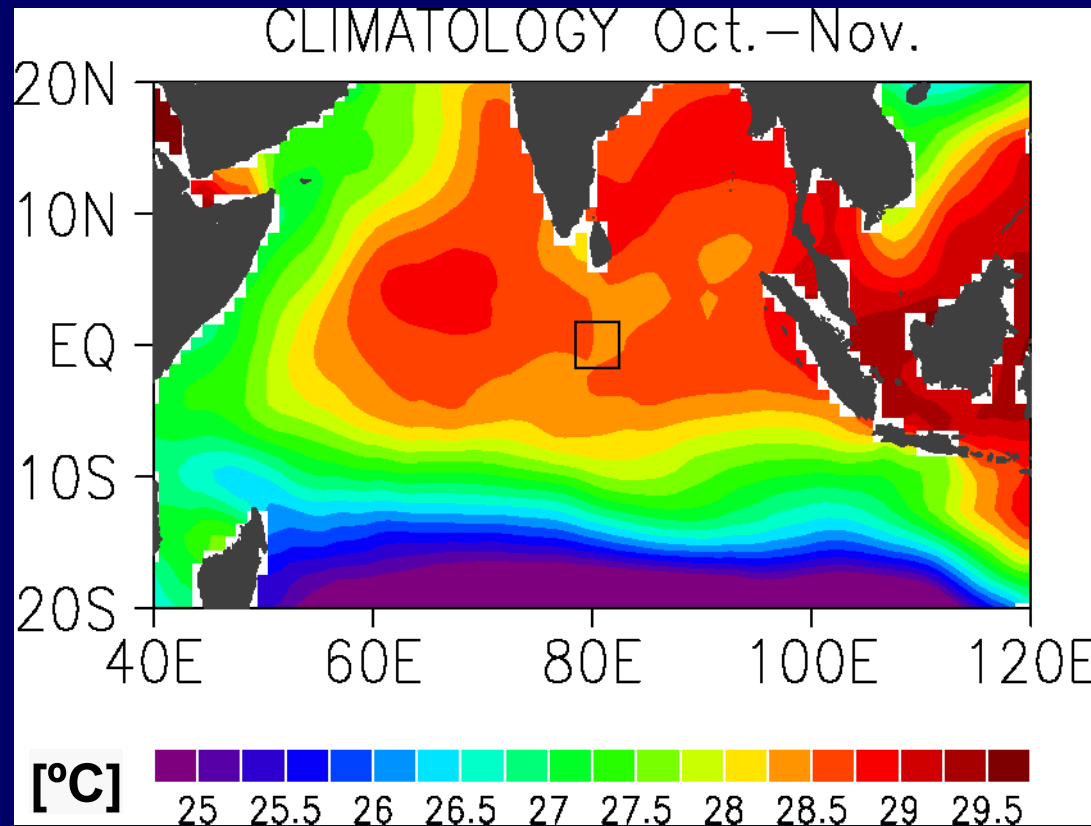
- **Equatorial Upwelling Estimated by the Ocean Current Data (ADCPs)**

Ocean Observation

- Ocean Intensive Observation during “**MISMO**”
 - Temperature, Salinity, Current, and Flux (Masumoto et al. 2008 GRL)
 - **m-TRITON Buoy & ADCP Array**
 - Useful data for Short-term Ocean-Atmosphere Interaction

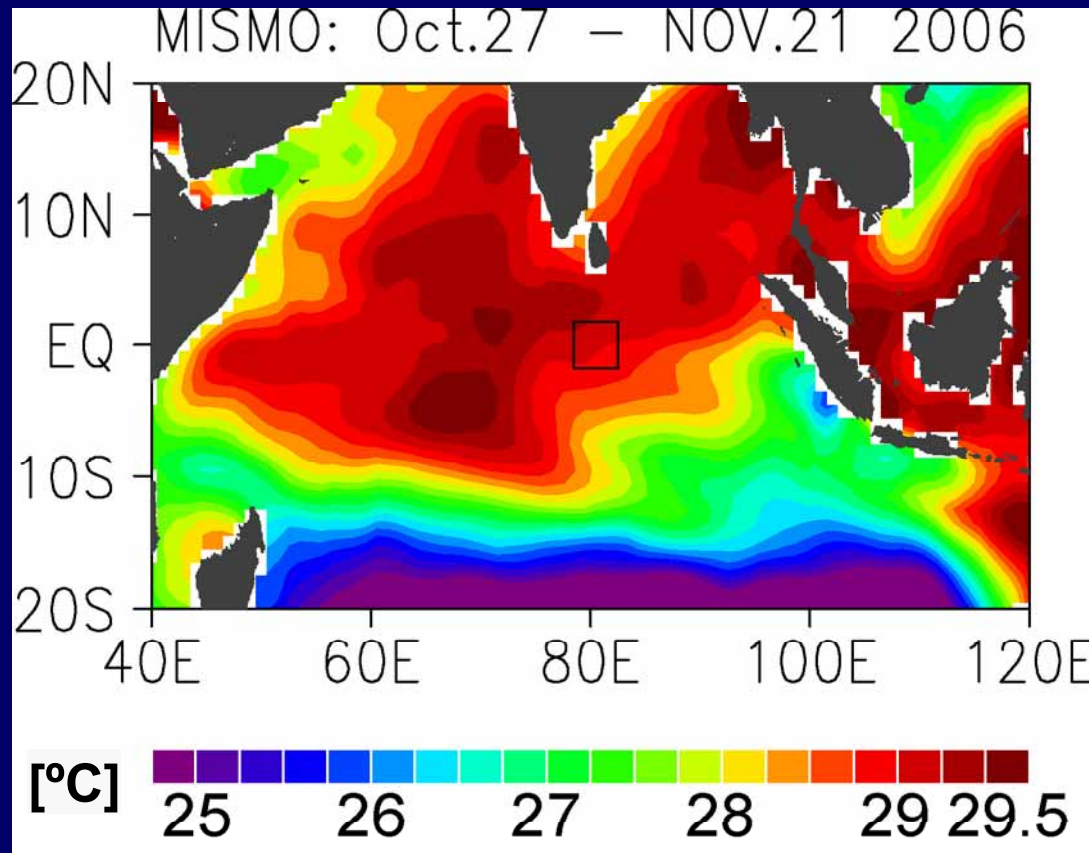


Background: Indian Ocean Dipole

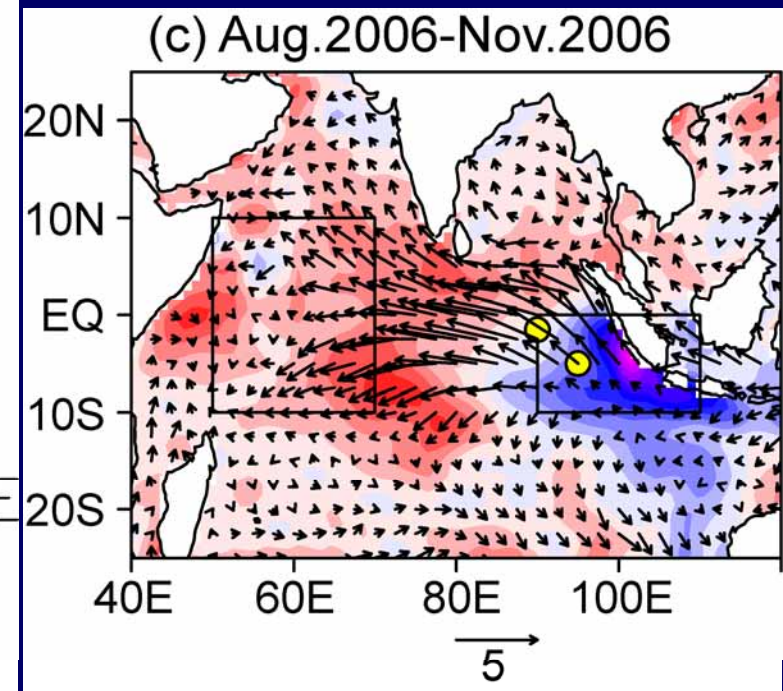


- Climatology in Oct-Nov
- No SST Gradient
- EQ Westerly Wind

Background: Indian Ocean Dipole



● MISMO Period

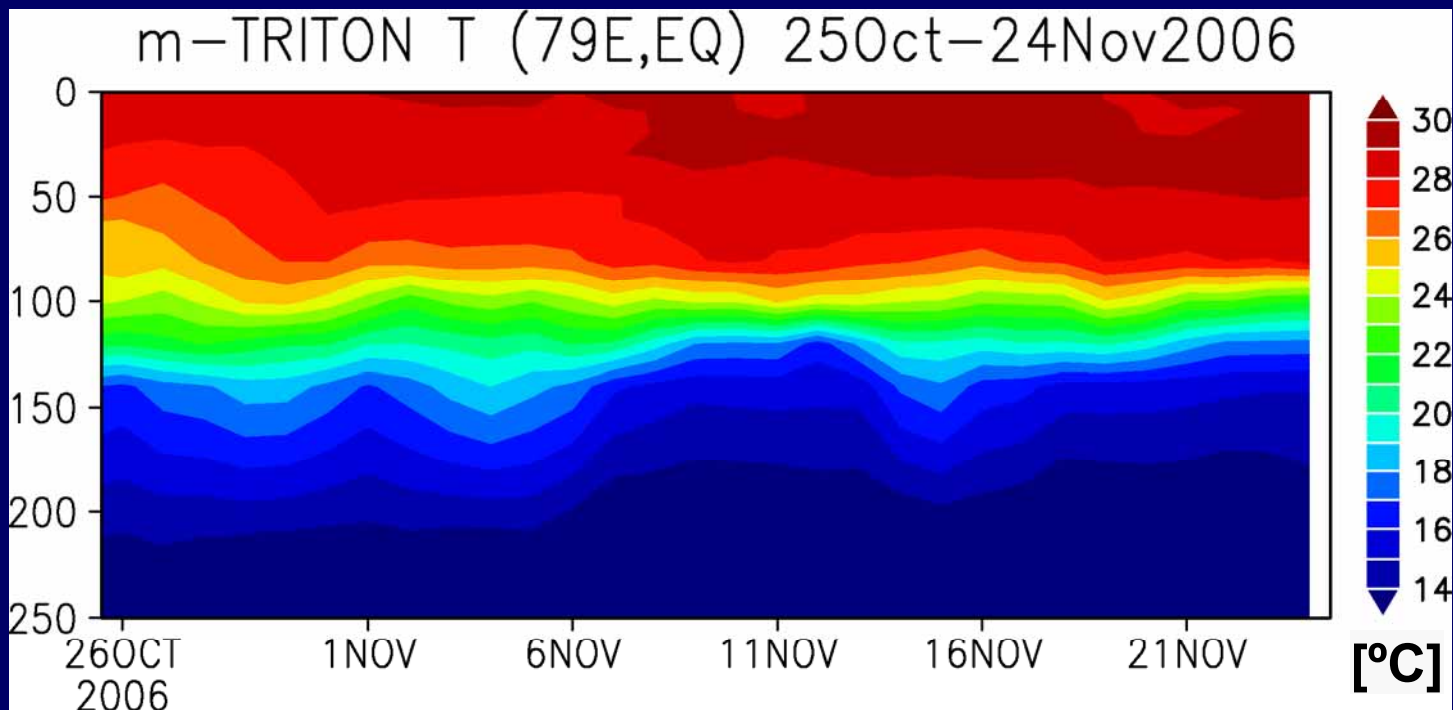


- Cooler SST
- Large SST Gradient in the East
- Reversal of Climatological Wind (Westerly → Easterly Wind)

SST & Wind Anomaly

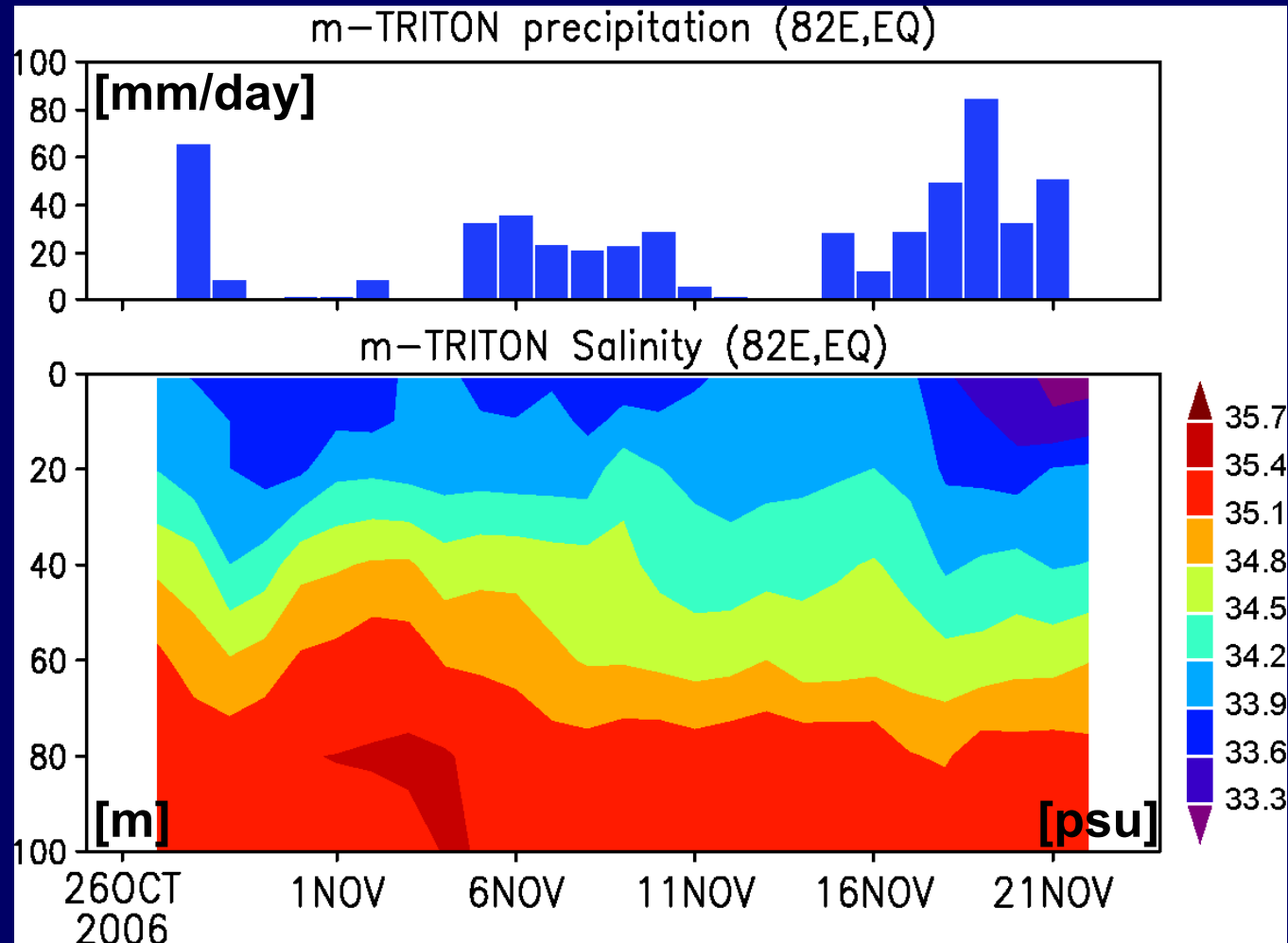
Temperature (79E, EQ)

- **Surface Layer: Warming up to $> 29^{\circ}\text{C}$**
- **Deeper Layer: Cooling in the bottom of Thermocline**
 - **Existence of Upwelling?**
- **Tightening of the Thermocline**

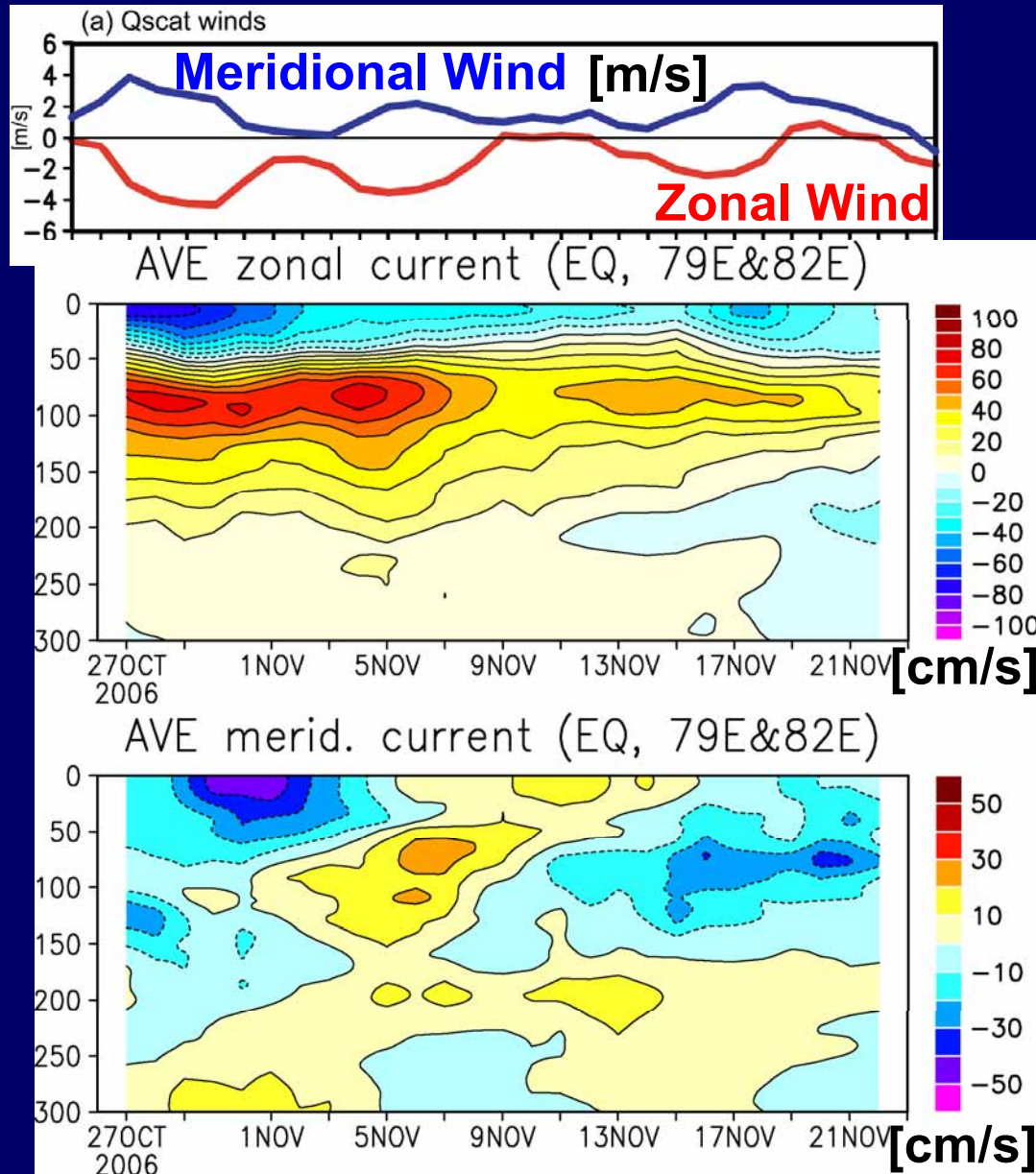


Salinity & Precipitation (82E, EQ)

- Low Salinity Water in the Upper Layer
 - Consistent with Precipitation
- High Salinity Water in the Thermocline Depth



Zonal & Meridional Current



ADCP & m-TRITON data

- Akima Spline Method

Zonal Current

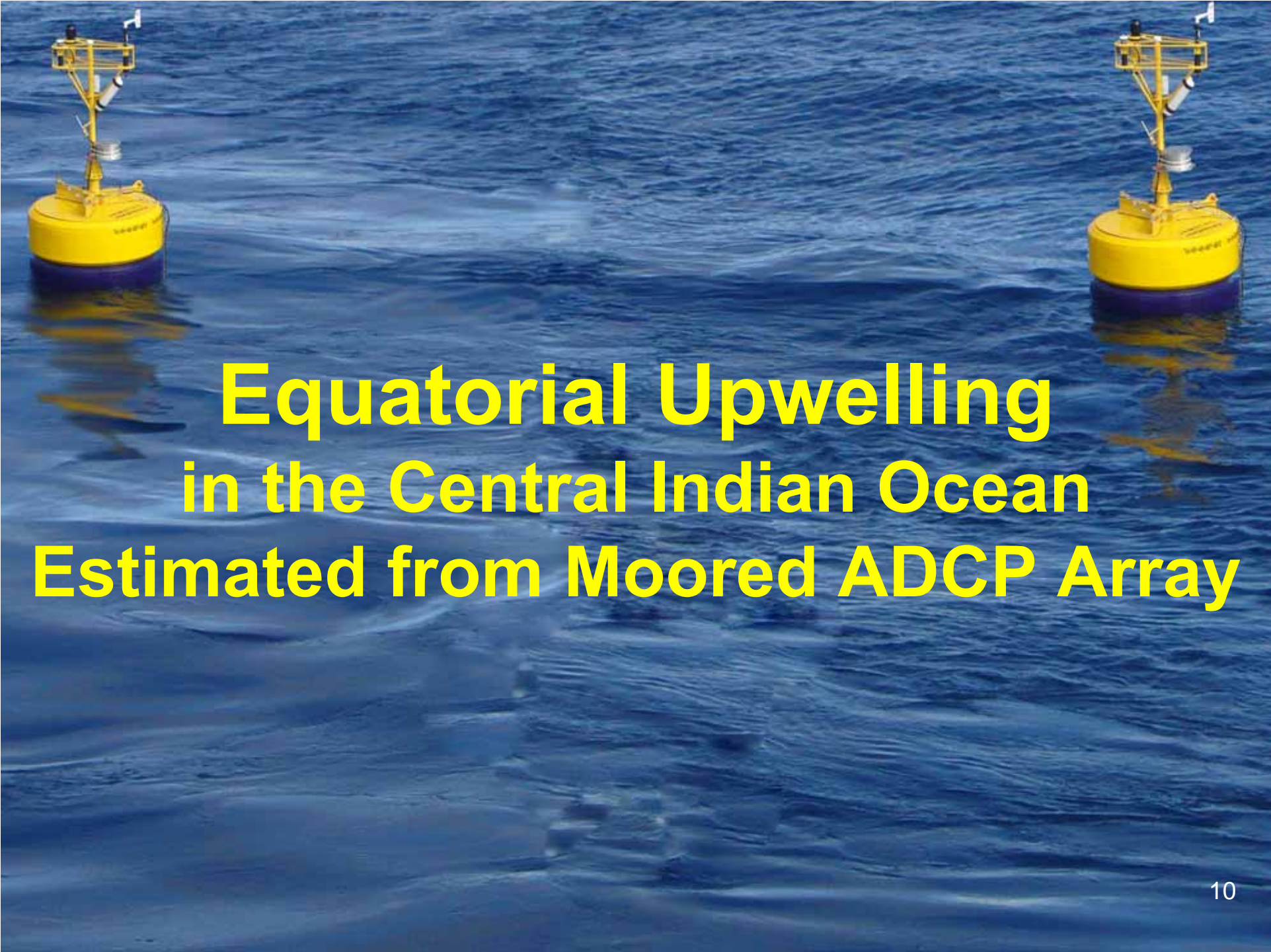
- Upper: Westward
- Subsurface: Eastward
(≠ Indian Ocean Clim.)
- Indian Ocean Dipole

Meridional Current

- 14-20 day period
→
- EQ Mixed Rossby Gravity Wave?
(Masumoto et al. 2005)⁸

Summary of Part I

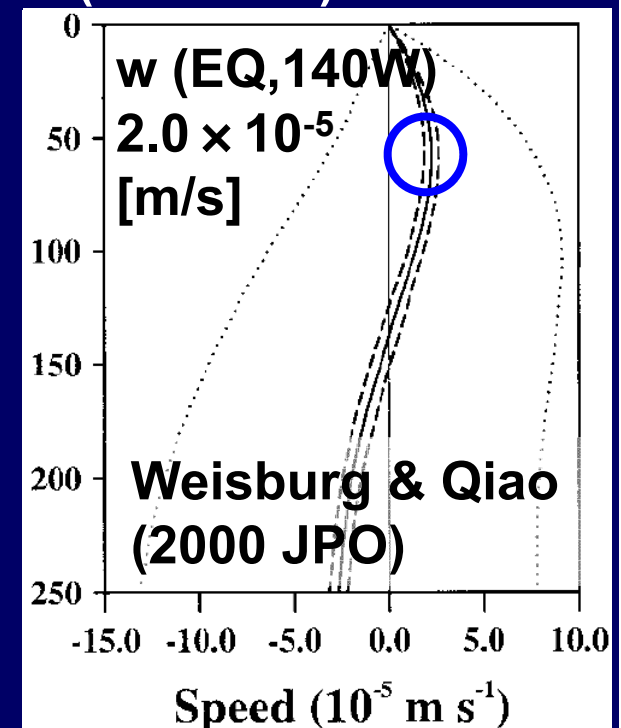
- **MISMO ocean observations reveal the unique conditions in the central equatorial Indian Ocean during the 2006 Indian Ocean Dipole.**
- **New findings are**
 - **Tightening of the thermocline with warming of the surface layer and cooling of the layer below the thermocline;**
 - **Large vertical salinity gradient, with association with the precipitation at the surface;**
 - **Surface westward and subsurface eastward current, and large vertical shear of the zonal currents above the thermocline; and**
 - **Short-term variability in the meridional current in the upper 150m depth.**

The image shows two identical moored ADCP (Acoustic Doppler Current Profiler) buoys floating in the deep blue ocean. Each buoy has a yellow upper section and a purple lower section. A metal frame on top of each buoy supports various instruments, including what appears to be a GPS antenna and a sensor. The water surface is slightly rippled, and the buoys are positioned symmetrically on either side of the center of the frame.

**Equatorial Upwelling
in the Central Indian Ocean
Estimated from Moored ADCP Array**

Background: EQ Vertical Velocity

- **EQ upwelling is an important element in the EQ ocean circulation**
 - Influence SST variation (→ Climate)
- **Order of Vertical Velocity (W) is $\sim 10^{-5}$ [m/s]**
 - Estimate of “ w ” in the Pacific (Ocean Current Array)
 - East: Halpern et al. (1989JGR); Meinen et al. (2001JPO)
 - West: Helber & Weiseberg (2001JGR)
 - Mean Upwelling: $\sim 2.0 \times 10^{-5}$ [m/s]
- **No Observation in the IO**
- **MISMO Observation**
 - Observation of Horizontal Current by Three or More ADCPs
 - Estimate of “ w ” by Horizontal Conv./Div.

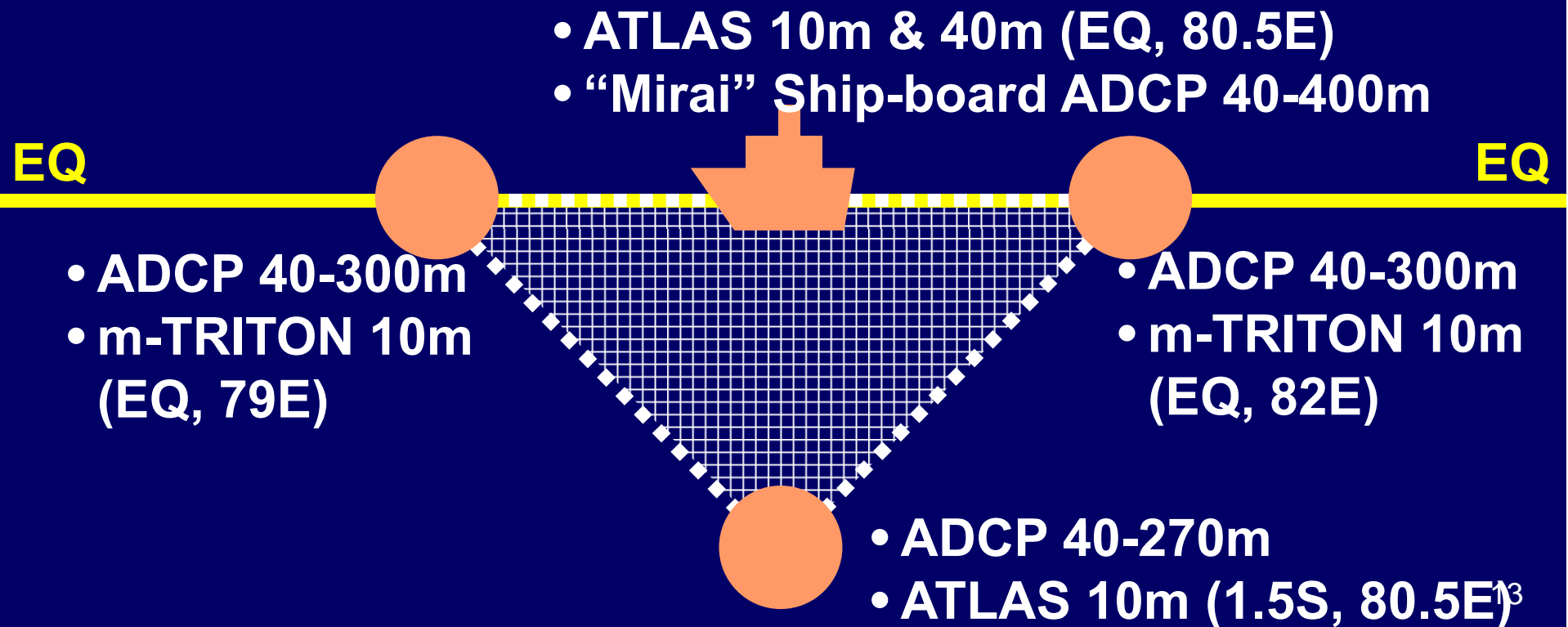


Objective

- **Big Goal:**
**Understanding 3-Dimensional Circulation
in the Equatorial Indian Ocean**
- **In this Study:**
**Estimate of Vertical Velocity “w”
Using the ADCP array during MISMO**
- **Content**
 - **Data & Method for Estimate of “w”**
 - **Result and Discussion**

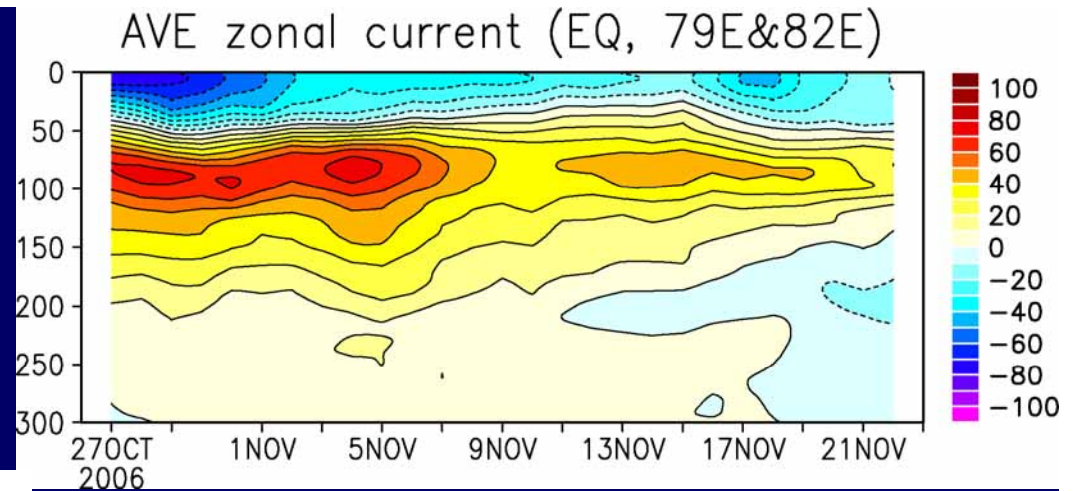
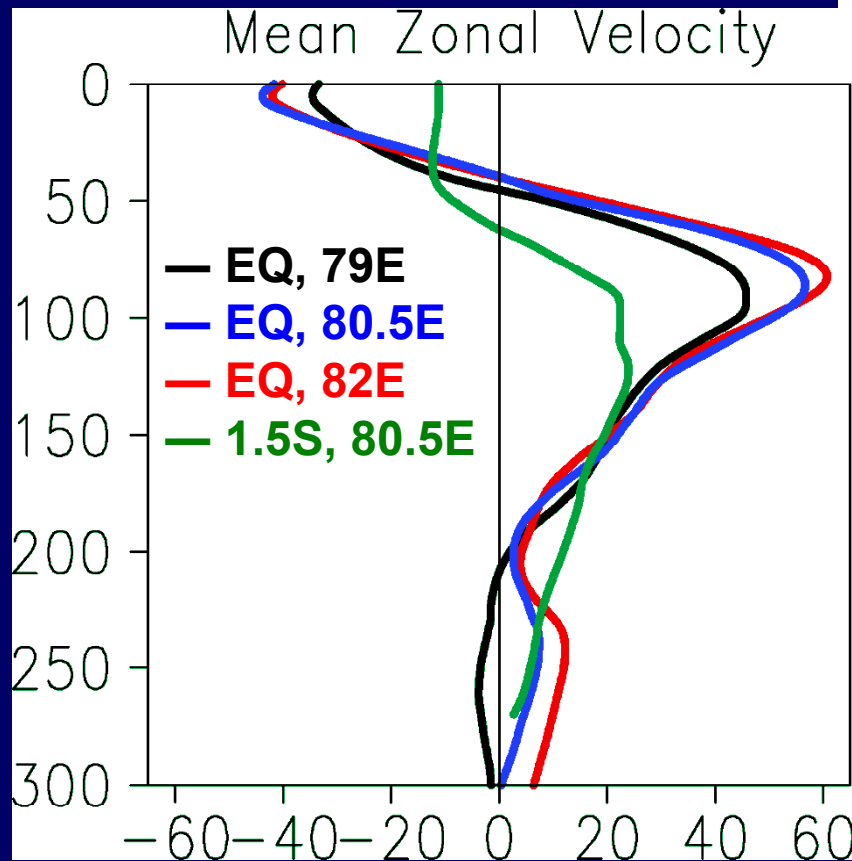
Data & Method

1. Merge “Buoy 10m data” & “ADCP 40-270m” Together
2. Interpolate Every 1m (0m~) by the Akima Spline Method
3. Calculate Conv./Div. in the Triangle Region,
and Integrate Vertically Applying the Continuity Equation



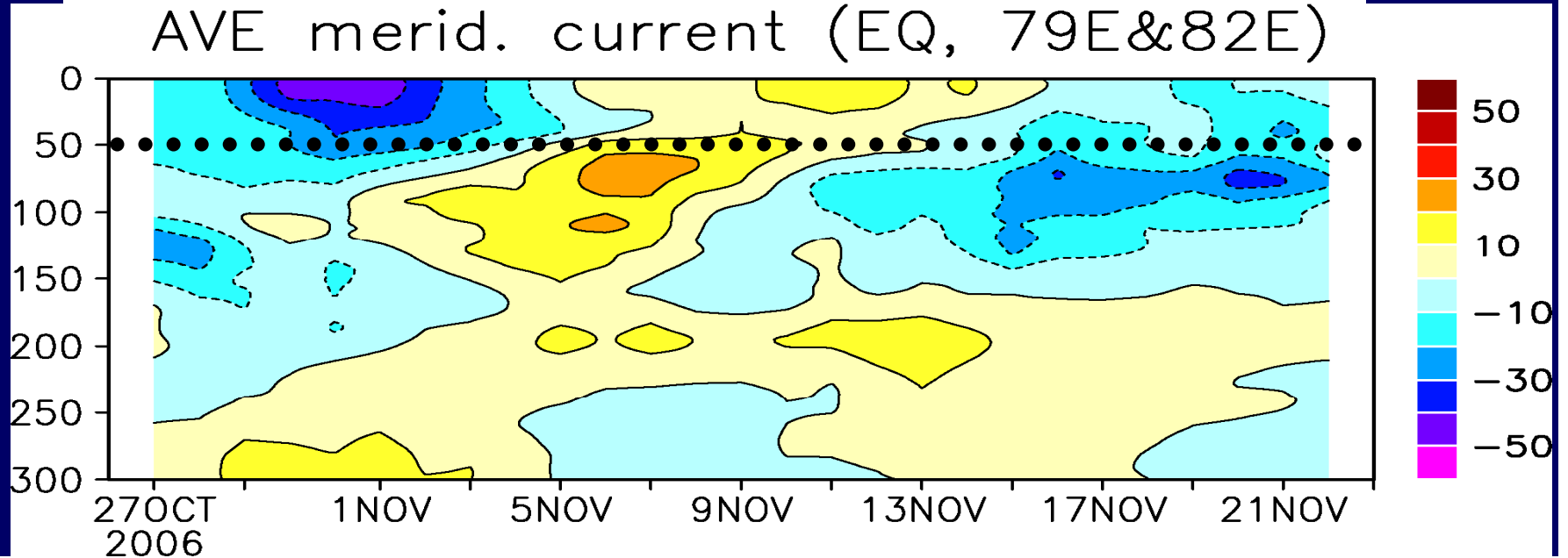
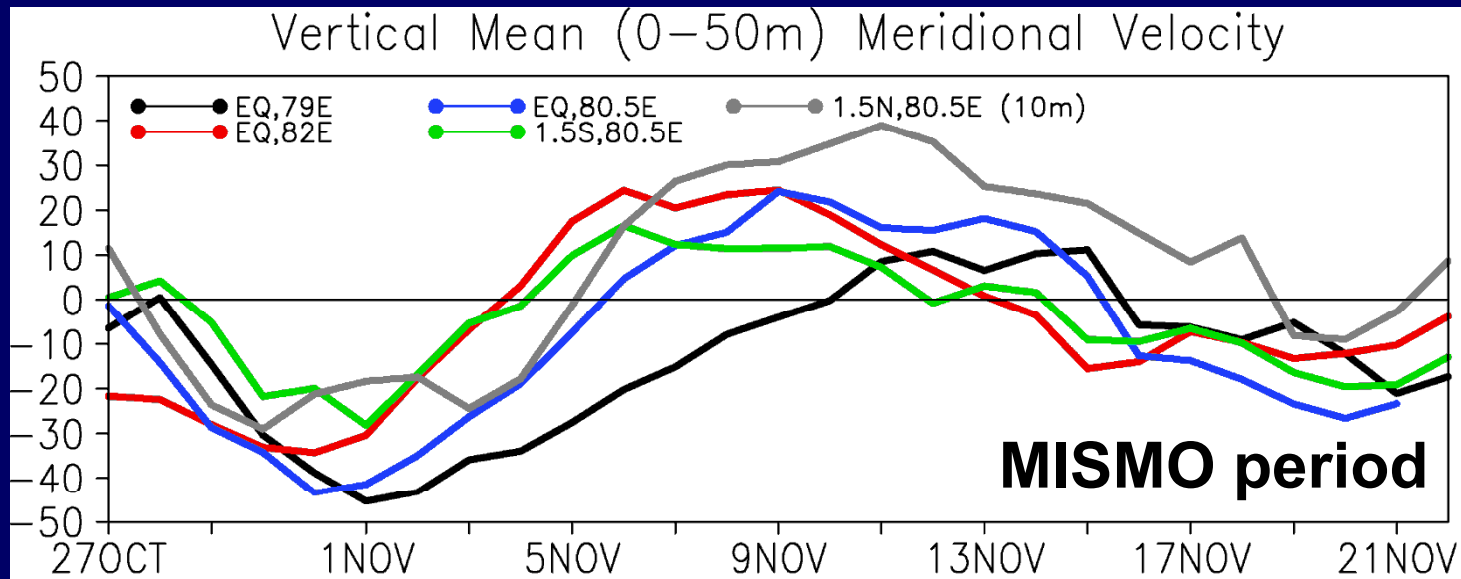
Zonal Current

↓ Mean Zonal Velocity

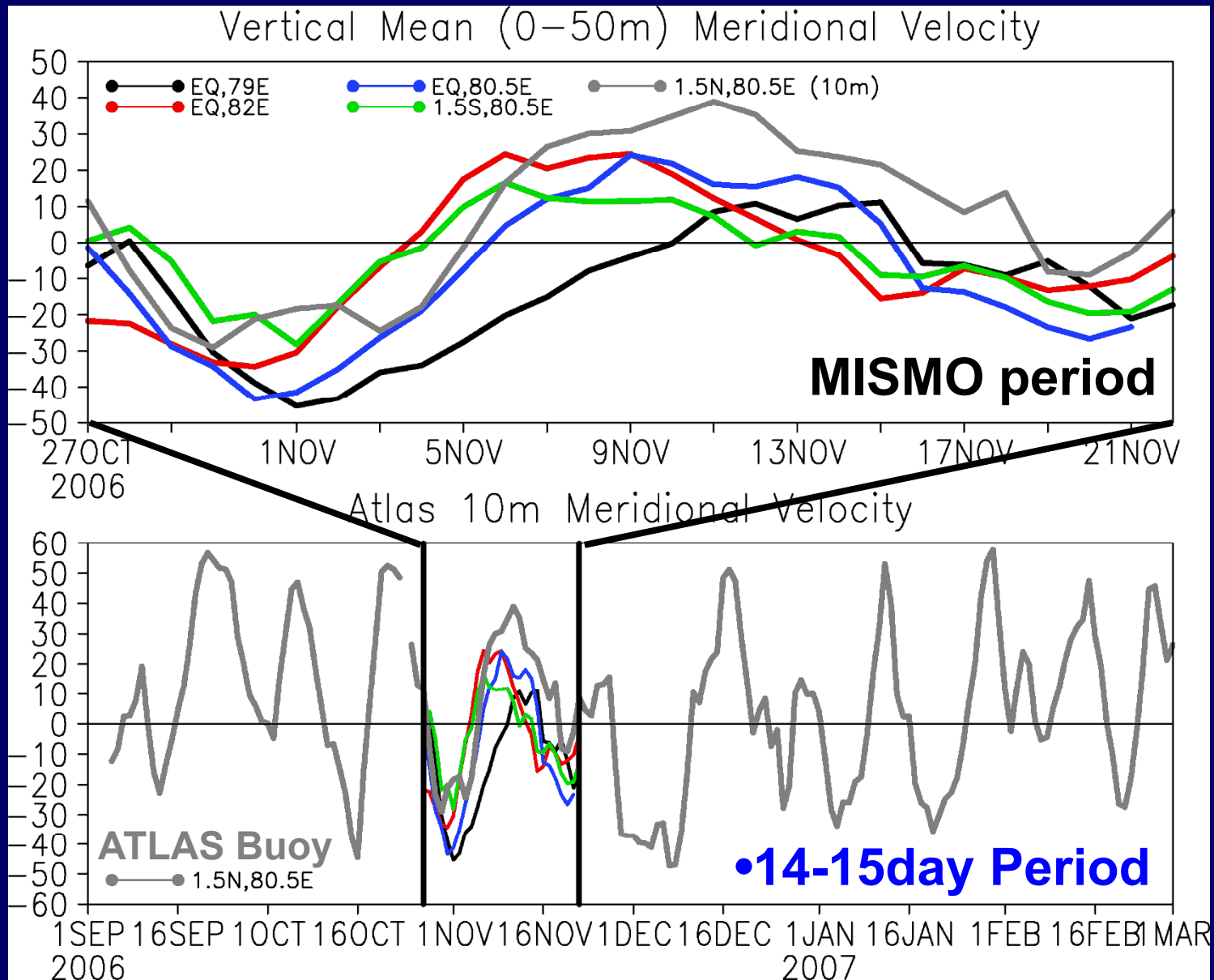


- Opposite Direction of Surface & Subsurface Current
- Acceleration of Eastward Current
- Weaker Current off of EQ

Meridional Current (0–50m)



Meridional Current (14-15day period)

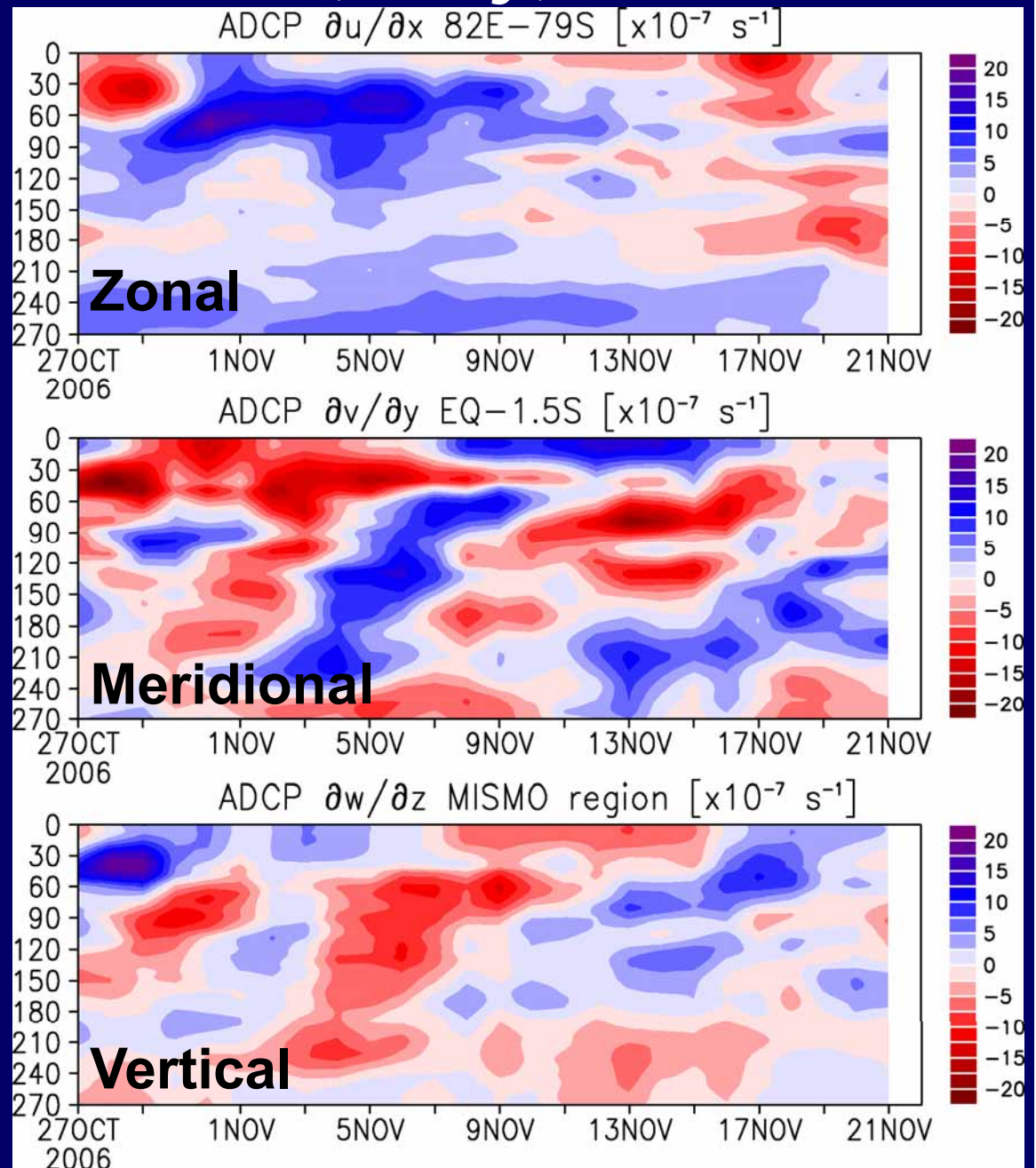


Conv/Div of Currents: $\partial u/\partial x, \partial v/\partial y, \partial w/\partial z$

- Positive (Blue): Div.
- Negative (Red): Conv.

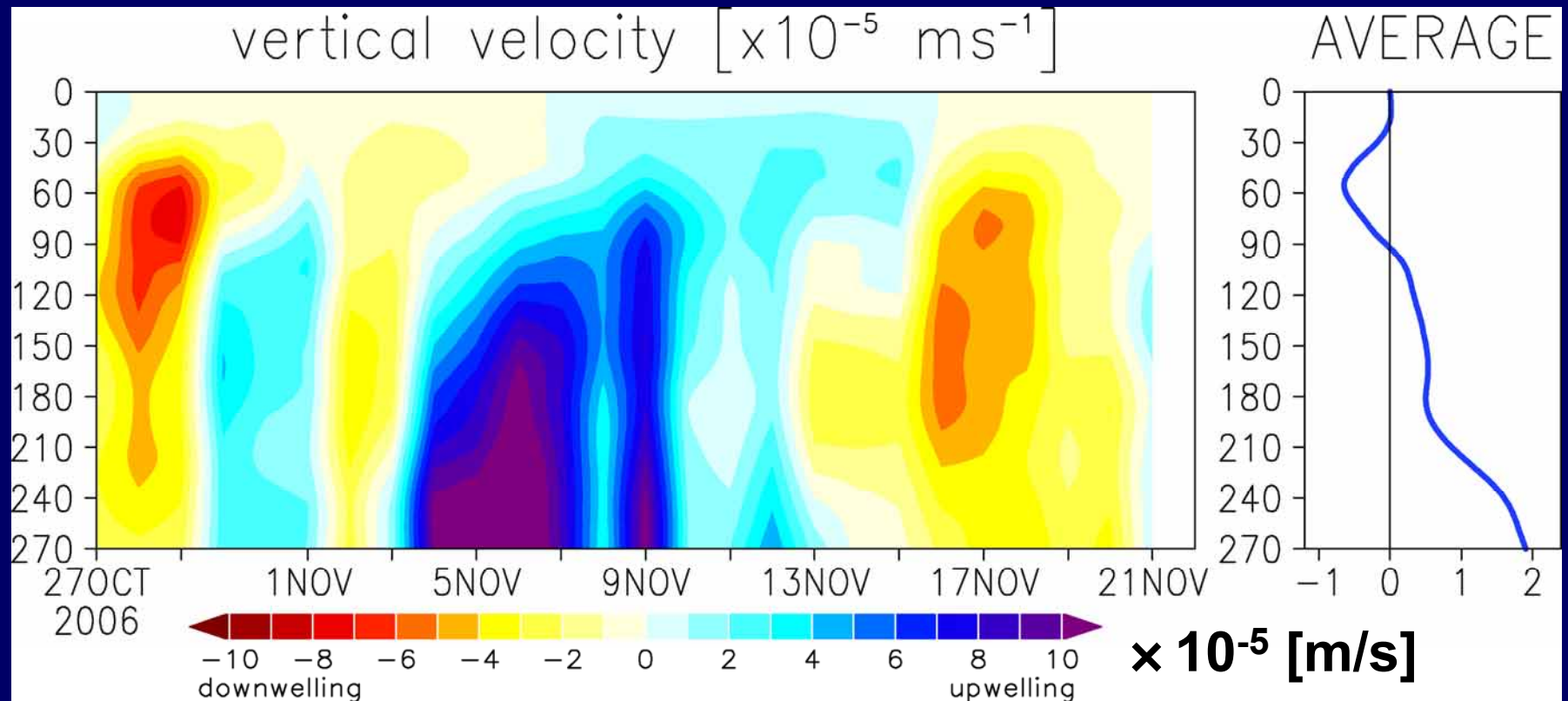
$$\frac{\partial w}{\partial z} = - \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) \quad (1)$$

- Continuity Eq. (1)
- Integrate Eq. (1)
($w=0$ at sea surface)
- Estimate
Vertical Velocity



Vertical Velocity

- Intraseasonal Strong Upwelling Nov.4~9
- Amplitude $> 10 \times 10^{-5}$ [m/s] (10 [m/day])

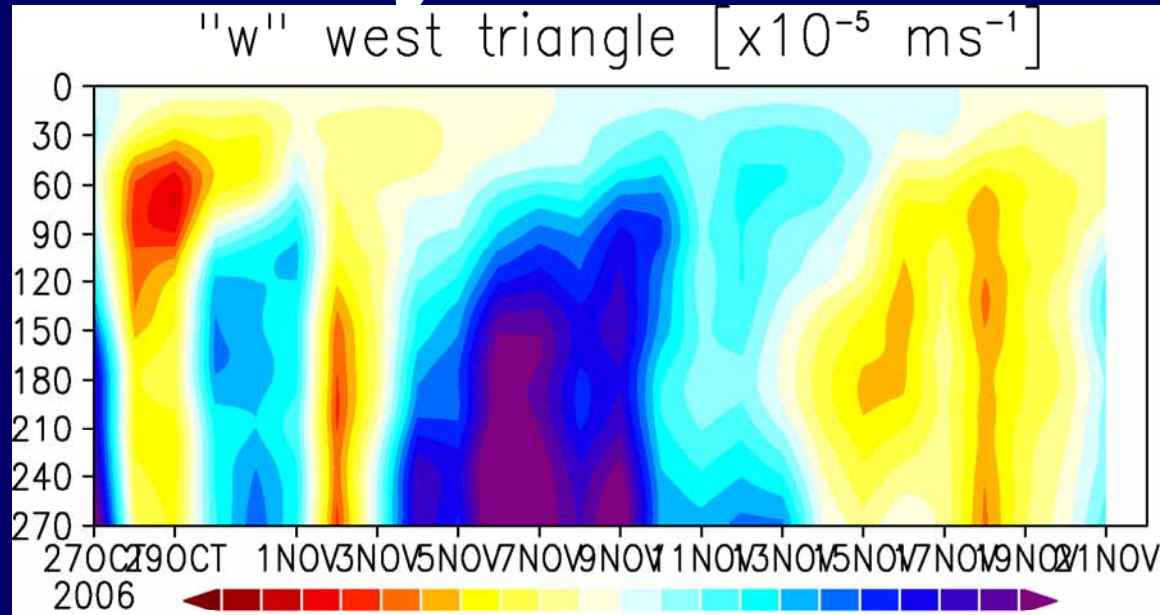
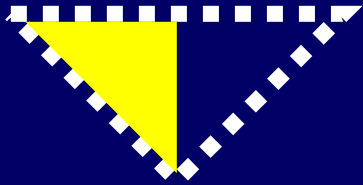


Depth-Time Section of “w”

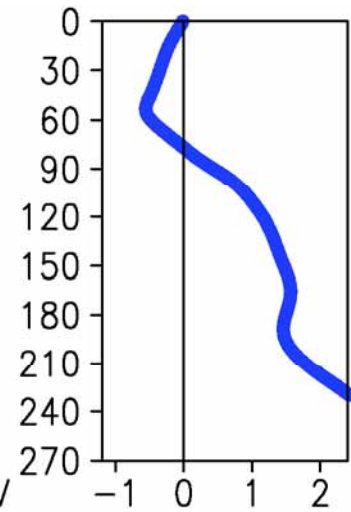
Mean of Period \uparrow 18

Vertical Velocity

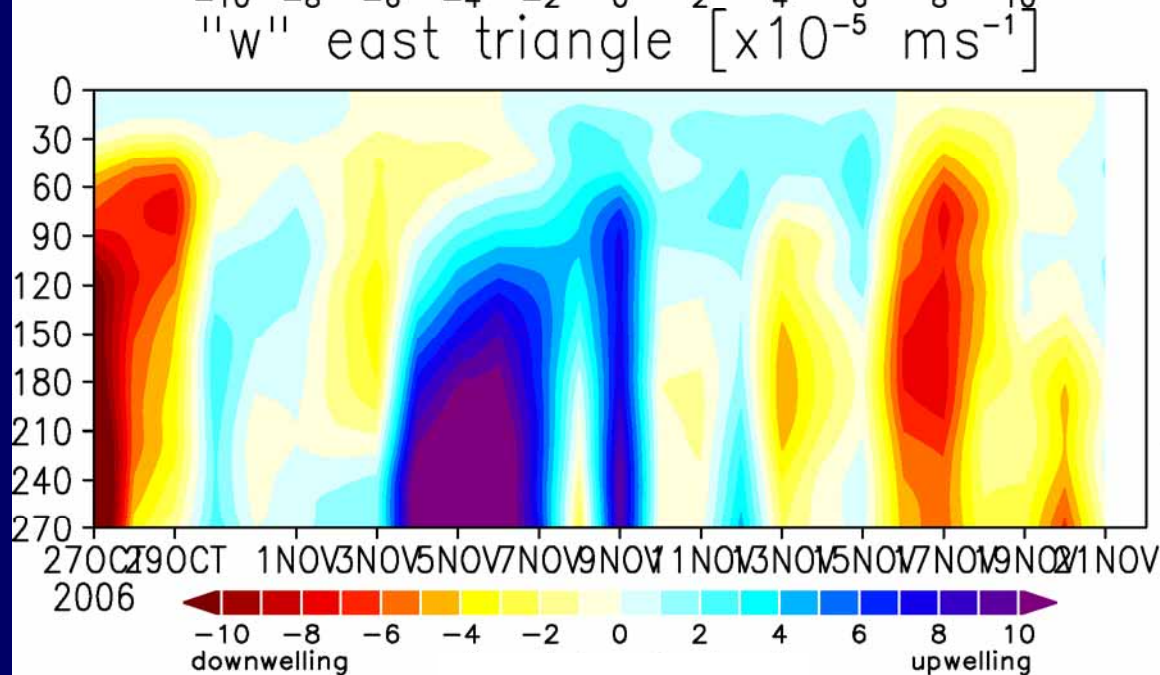
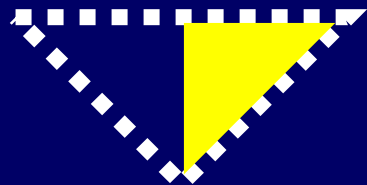
West Region



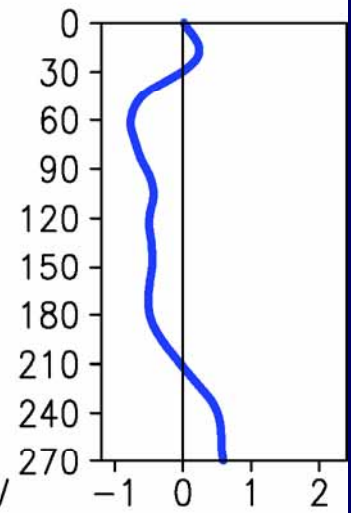
AVERAGE



East Region

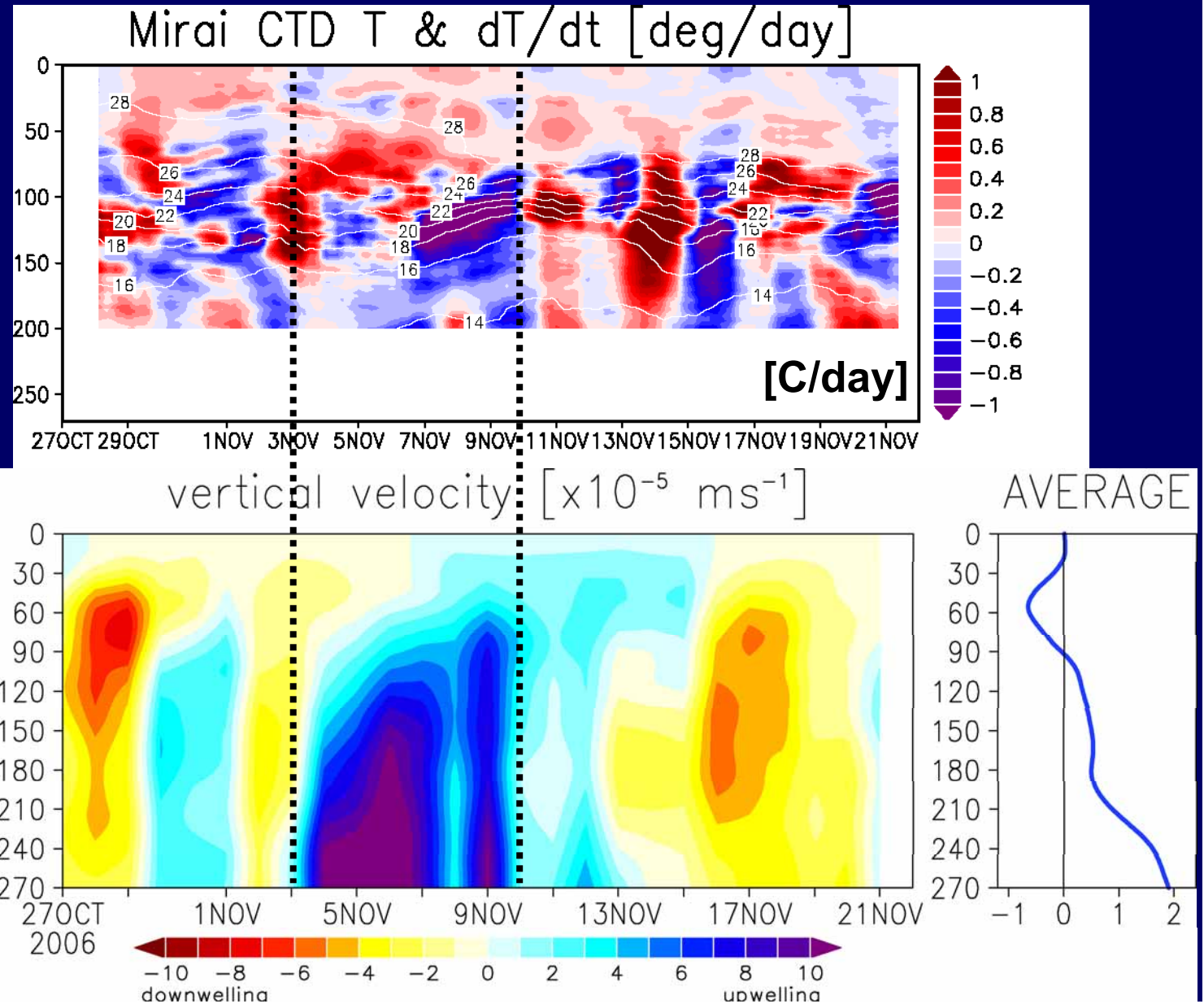


AVERAGE



T Change ($\partial T/\partial t$) & Vertical Velocity

R/V Mirai
CTD
Measurement
(Every 3 Hour)



Summary of Part II

- Using an array of the ADCPs, **Vertical Velocity** “w” is estimated by the continuity equation.
 - Intraseasonal strong upwelling event below 90m
 - Large amplitude of 10×10^{-5} [m/s] (~10 [m/day])
 - Lasted about a week in the MISMO period
- The upwelling was produced by the zonal & meridional divergence.
 - Zonal current: Subsurface strong eastward current accelerated eastward
 - Meridional current: 14-20day variation, suggesting mixed Rossby gravity wave
- Future:
 - Possible influences of the upwelling on the biological activity should be examined.