

Eastward propagation mechanism and asymmetric horizontal structure of super cloud clusters in the equatorial region

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Contents:

- (1) Strategy,
- (2) Observations and numerical results of super cloud clusters (SCCs),
- (3) Simple linear or nonlinear models (32 constant vertical layers; N , $C_s = \text{constant}$),
- (4) Full nonlinear model
(grid location same as NICAM, and environmental fields, such as zonal wind, N , C_s etc, are given by NICAM outputs),
* Role of the equatorial wave on EP mechanism of SCCs,
- (5) Expansion of our results.

(1) Our strategy

- To get **coherent organized structures**, “**instability**” is essential, while “**neutral waves**” are only needed to make response patterns around forcing. However, tropical (large-scale) researchers tend to consider that “**waves**” are primary and “**instability (=convection)**” is secondary. However, “**CONVECTION**” is **PRIMARY** and “**equatorial waves**” are secondary.

- **Main doubts:** “Is the coupling of instability and neutral waves, such as convectively coupled Kelvin waves, available ?”

At least, **NO** in the linear theory. Then, what about nonlinear case?

>> (I want to conclude) **NO !**

Assumptions (to get deep physical understanding as simply as possible)

- **Positive-only wave CISK** is included to describe the isolated features of diabatic heating due to precipitation.
- **Large horizontal viscosity/diffusivity** is included to attain preferred disturbances of a horizontal scale $O(1,000 \text{ km})$. These large-scale disturbances are needed for the equatorial beta effect to work.

(2) Observations

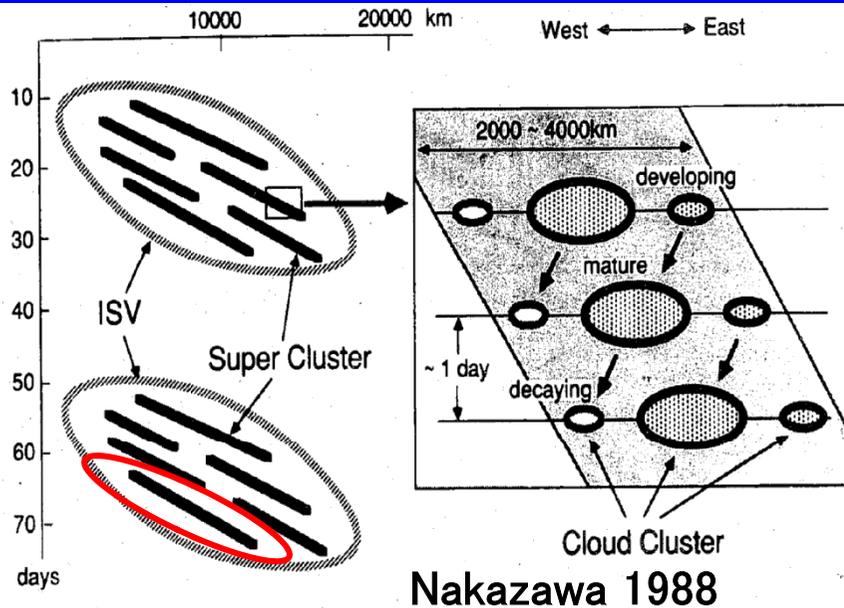


Fig. 5. Schematic diagram for the hierarchy of ISV.

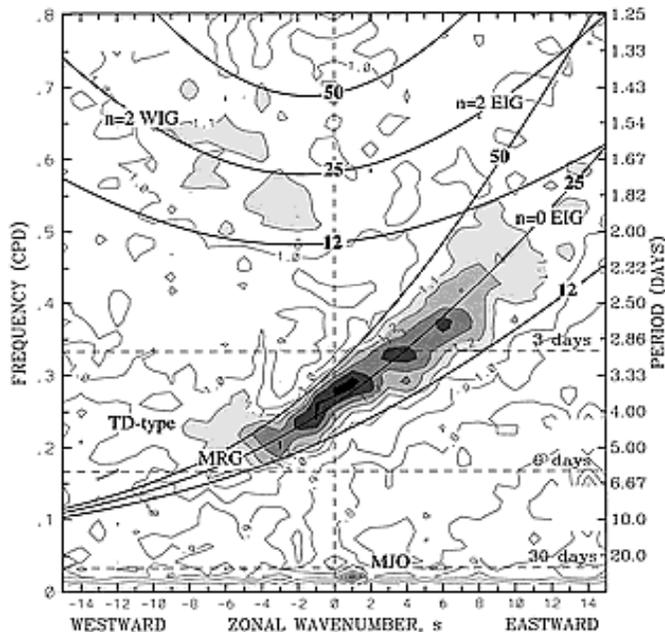
Super cloud clusters (SCCs)

* horizontal scale: $O(1,000 \text{ km})$

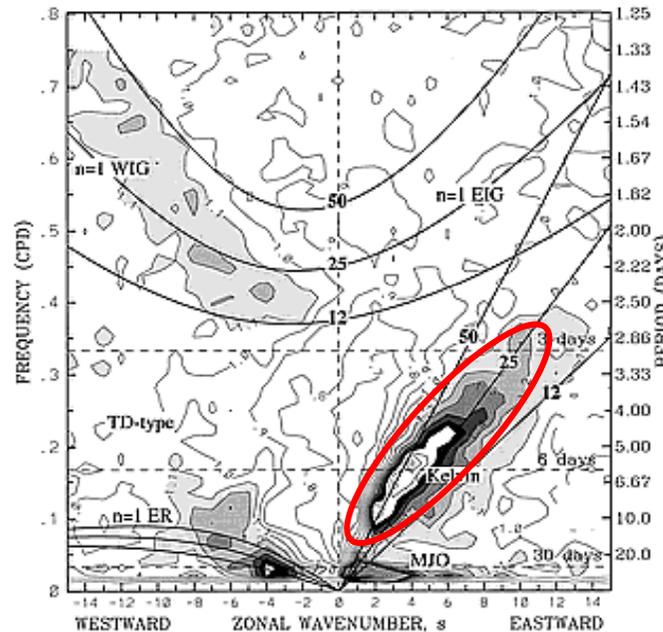
* time scale: $O(10 \text{ days})$

SCCs; fundamental components of MJO with horizontal scale of $O(10,000 \text{ km})$, and MJO may be an envelope of SCC-ensembles

a) $\left\{ \sum_{15S}^{15N} \text{POWER(OLR A)} \right\} / \text{BACKGROUND}$



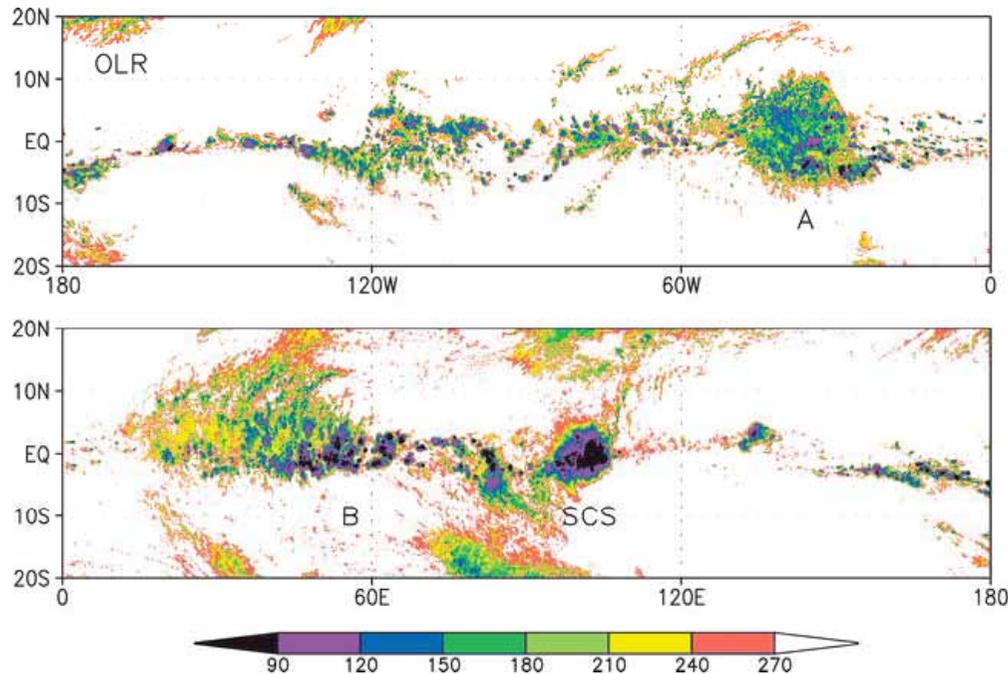
b) $\left\{ \sum_{15S}^{15N} \text{POWER(OLR S)} \right\} / \text{BACKGROUND}$



Wheeler · Kiladis
1999

(2)' Numerical results

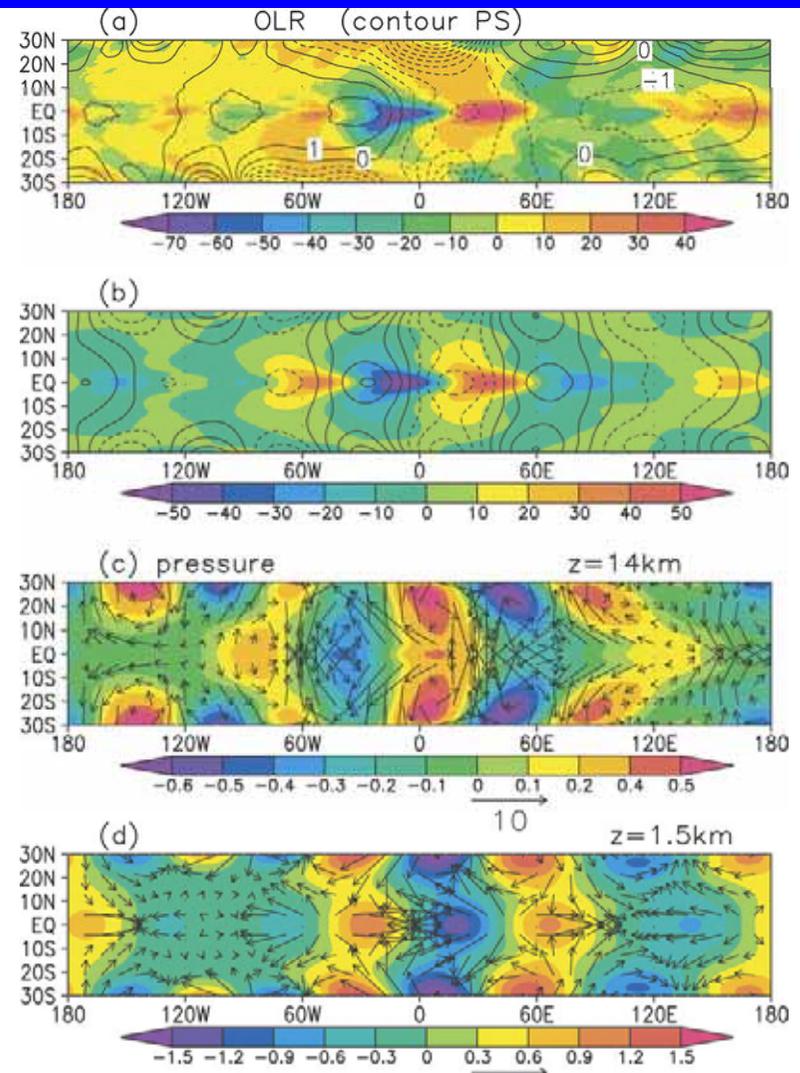
NICAM aquaplanet simulation



Nasuno et al. 2007

In $O(1000\text{km})$ scale of SCCs,

- $O(100\text{km})$ mesoscale convective system (MCS),
 - $O(10\text{km})$ scale cumulonimbus cloud
- **Multi-scale (hierarchical) structures** of cloud systems

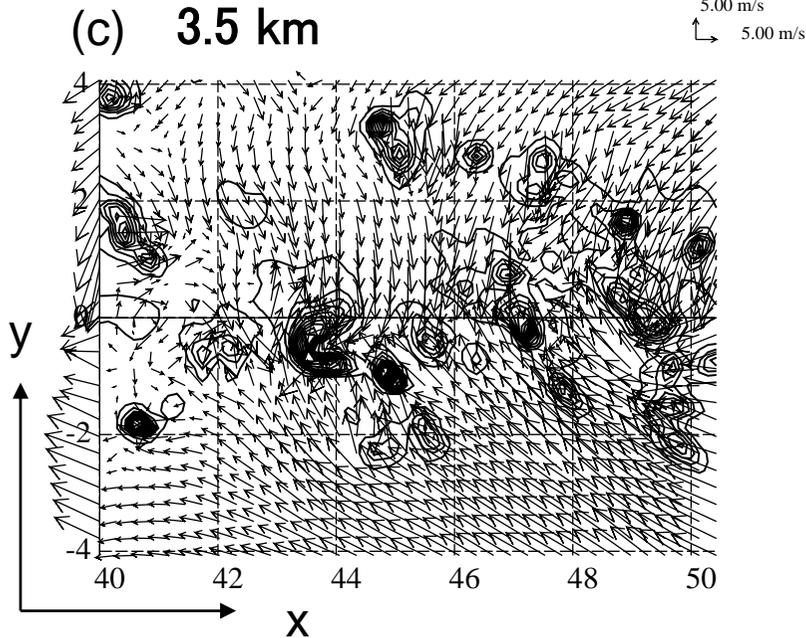
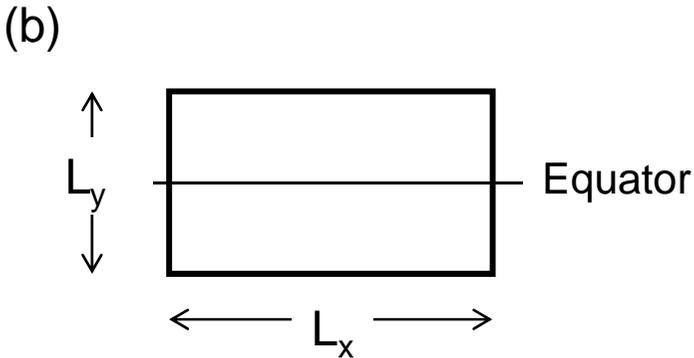
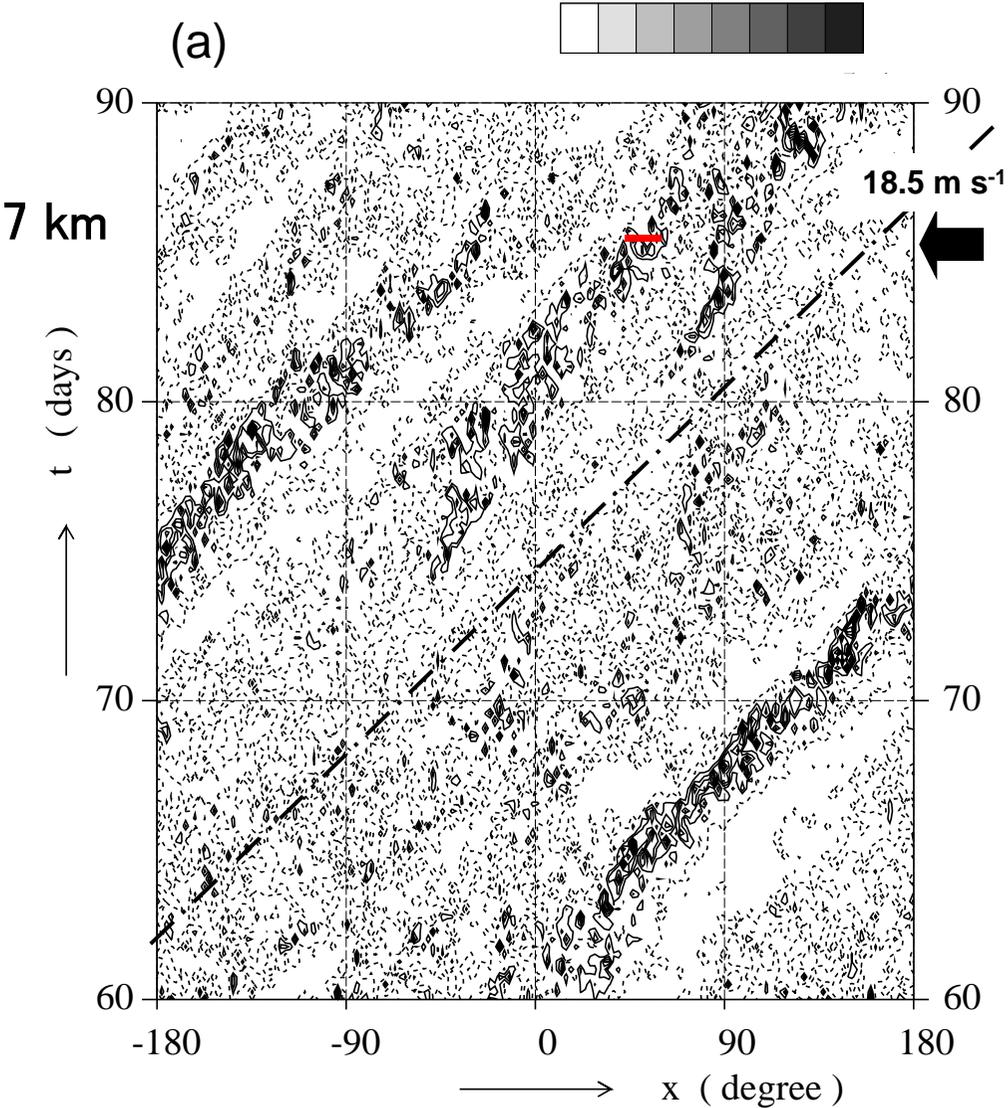


Nasuno et al. 2008

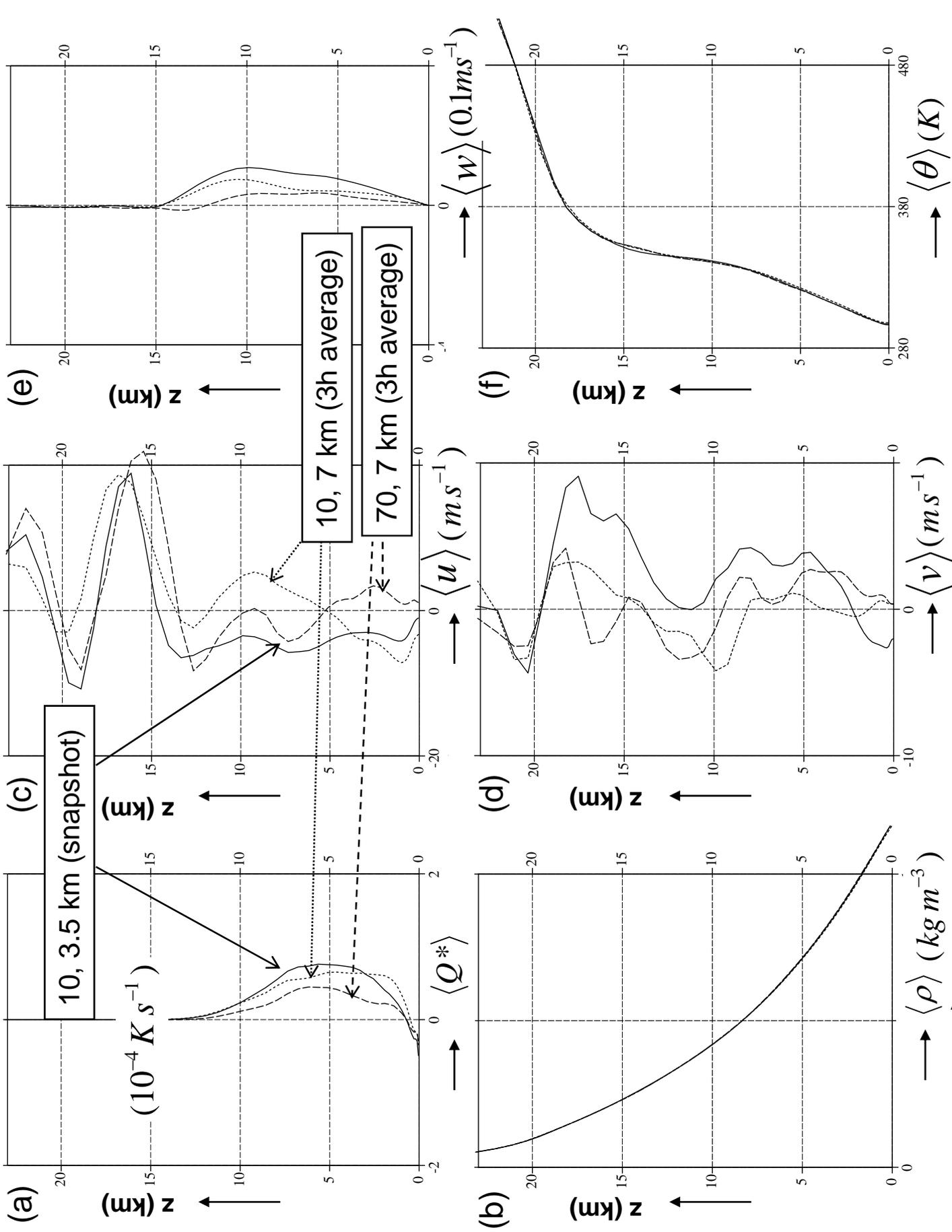
Structure of SCC mode

- a pair of off-equatorial gyre
- ⇒ **Gill pattern**

NICAM aquaplanet simulation



$\langle \rangle$: (L_x, L_y) averaging
 $()'$: deviations



Numerical simulation with a dry model

$$U' = \langle \rho \rangle u', V' = \langle \rho \rangle v', W' = \langle \rho \rangle w', B' = \langle \rho \rangle g \frac{\theta'}{\langle \theta \rangle}$$

$$N^2 = \frac{g}{\langle \theta \rangle} \frac{\partial \langle \theta \rangle}{\partial z}, C_s = \text{speed of sound wave}:$$

$$\frac{\partial U'}{\partial t} = \varepsilon_1 NL(U') - \varepsilon_2 \langle u \rangle \frac{\partial U'}{\partial x} - \varepsilon_2 W' \frac{\partial \langle u \rangle}{\partial z} - \frac{\partial p'}{\partial x} + \beta y V' - \varepsilon_3 r U',$$

$$\frac{\partial V'}{\partial t} = \varepsilon_1 NL(V') - \varepsilon_2 \langle u \rangle \frac{\partial V'}{\partial x} - \frac{\partial p'}{\partial y} - \beta y U' + \nu_H \Delta_H V' - \varepsilon_3 r V',$$

$$0 = -\frac{\partial p'}{\partial z} - \frac{g}{C_s^2} p' + B',$$

$$\frac{\partial B'}{\partial t} = \varepsilon_1 NL(B') - \varepsilon_2 \langle u \rangle \frac{\partial B'}{\partial x} - W' N^2 + Q + \nu_H \Delta_H B' - \varepsilon_3 r B',$$

$$\frac{\partial U'}{\partial x} + \frac{\partial V'}{\partial y} + \frac{\partial W'}{\partial z} = 0.$$

- Horizontal grid size: 1 degree (360 x 60)

- Vertical grid size : uniform or variable

Large viscosity

Large diffusivity

Diabatic heating Q:

* Positive-only wave-CISK

$$Q = \begin{cases} w_B \cdot F(z) & \text{for } w_B > 0 \\ 0 & \text{for } w_B \leq 0 \end{cases}$$

Cloud base
Height ≈ 1km

* Parameters ε : 0 / 1

ε_1 : Nonlinear (NL) / Linear

ε_2 : Zonal wind / No zonal wind

ε_3 : Rayleigh damping / No damping

N^2 : Variable / Uniform

ν_H : Large values

to attain O(1,000km)-scale disturbances

(3) Simple linear model (32 layers; N, C_s=constant)

$$N^2 = \frac{g}{\langle \theta \rangle} \frac{\partial \langle \theta \rangle}{\partial z}, \quad C_s = \text{speed of sound wave:}$$

$$\frac{\partial U'}{\partial t} = -\frac{\partial p'}{\partial x} + \beta y V' + v_H \Delta_H U', \quad \frac{\partial V'}{\partial t} = -\frac{\partial p'}{\partial y} - \beta y U' + v_H \Delta_H V',$$

$$0 = -\frac{\partial p'}{\partial z} - \frac{g}{C_s^2} p' + B', \quad \frac{\partial B'}{\partial t} = -W' N^2 + Q + v_H \Delta_H B',$$

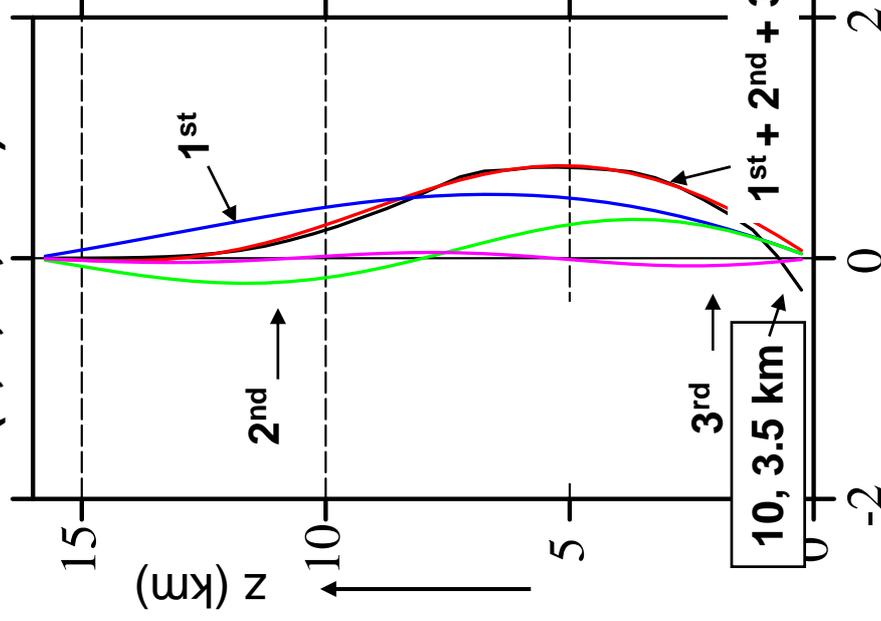
$$\frac{\partial U'}{\partial x} + \frac{\partial V'}{\partial y} + \frac{\partial W'}{\partial z} = 0.$$

- $\Delta z = 500\text{m}$, 32 layers, $N = 10^{-2} \text{ s}^{-1}$, $C_s = 300 \text{ m s}^{-1}$

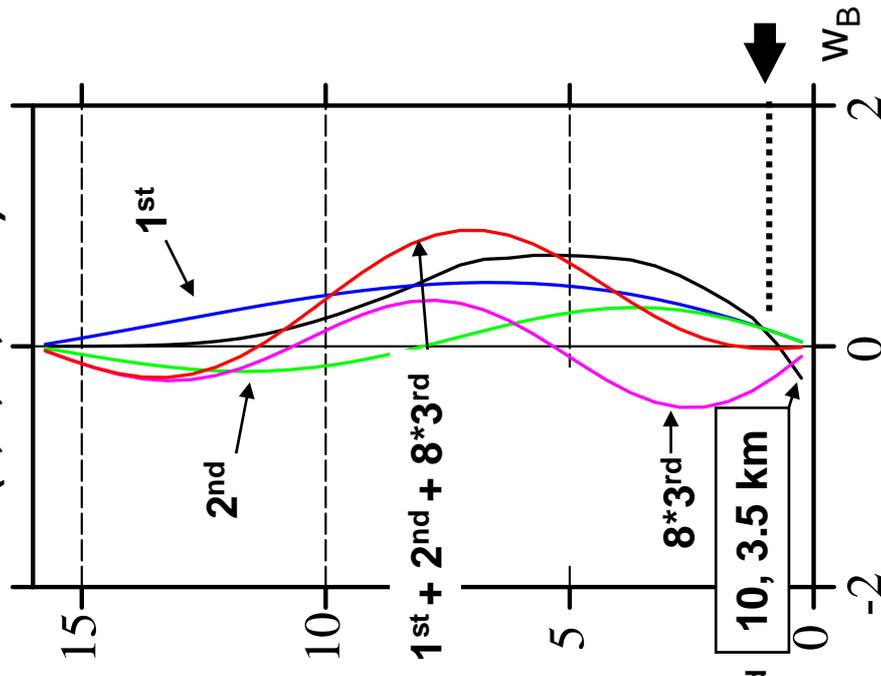
- vertical-mode expansion

$$\frac{d^2 h_n(z)}{dz^2} + \frac{g}{\langle C_s^2 \rangle} \frac{d h_n(z)}{dz} + n\pi h_n(z) = 0$$

(a) (1, 1, -1, 0...)



(b) (1, 1, -8, 0...)



- (Q, W', B') vs. (U', V', p')

- Vertical-mode

expansion of Q

$$= 0.559 h_1(z) + 0.282 h_2(z) - 0.053 h_3(z) - 0.040 h_4(z) \dots$$

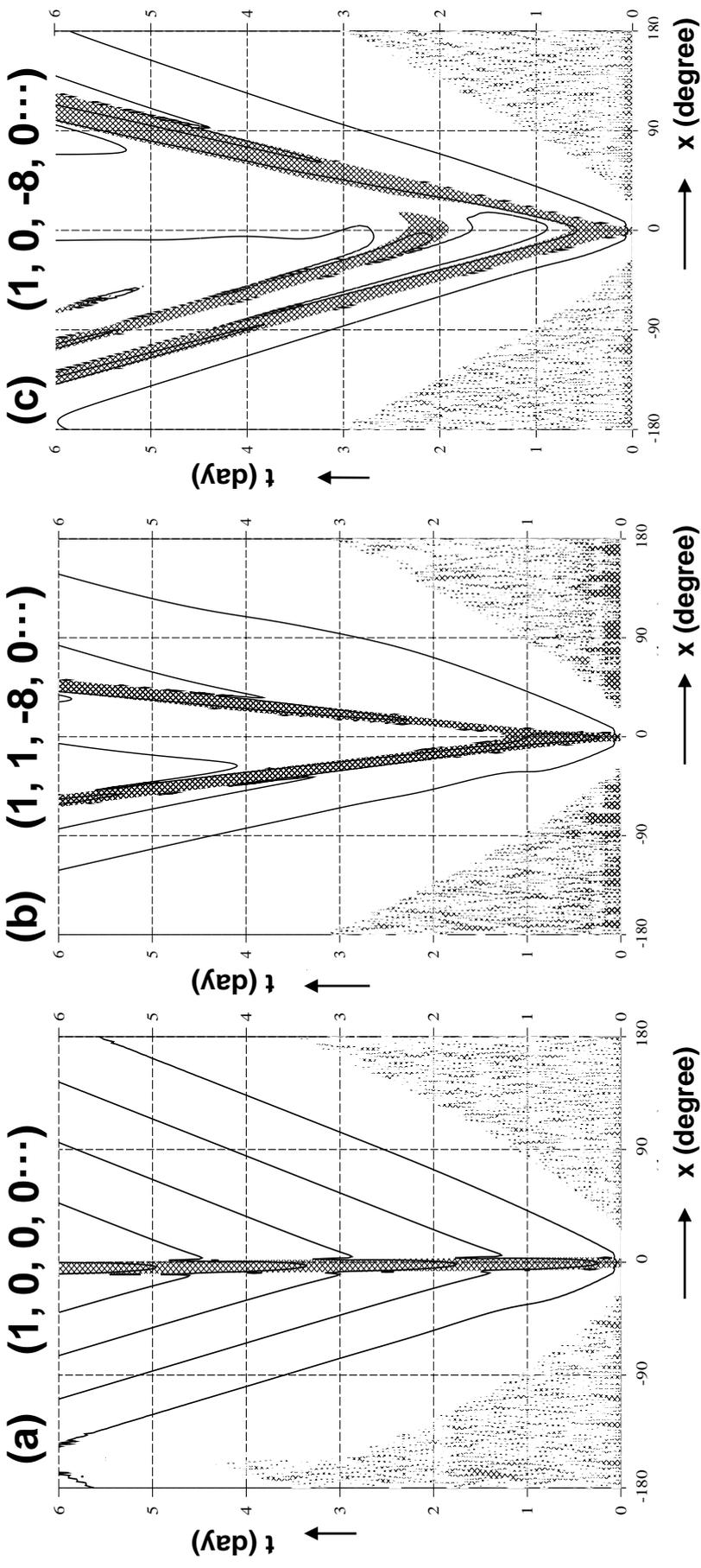
as (1, 1, -1, -1...)

- Q can be nearly expressed by three vertical modes.

→ (1, 1, -8, 0...)

for comparison

Linear case ($\log_{10}|w_{3.5\text{km}}|$)



A slowly WP unstable D.

A pair of WP and EP unstable Ds.

A pair of WP and EP unstable Ds.

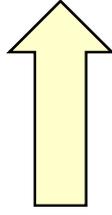
Propagation property:

(0,1,-8) is essential

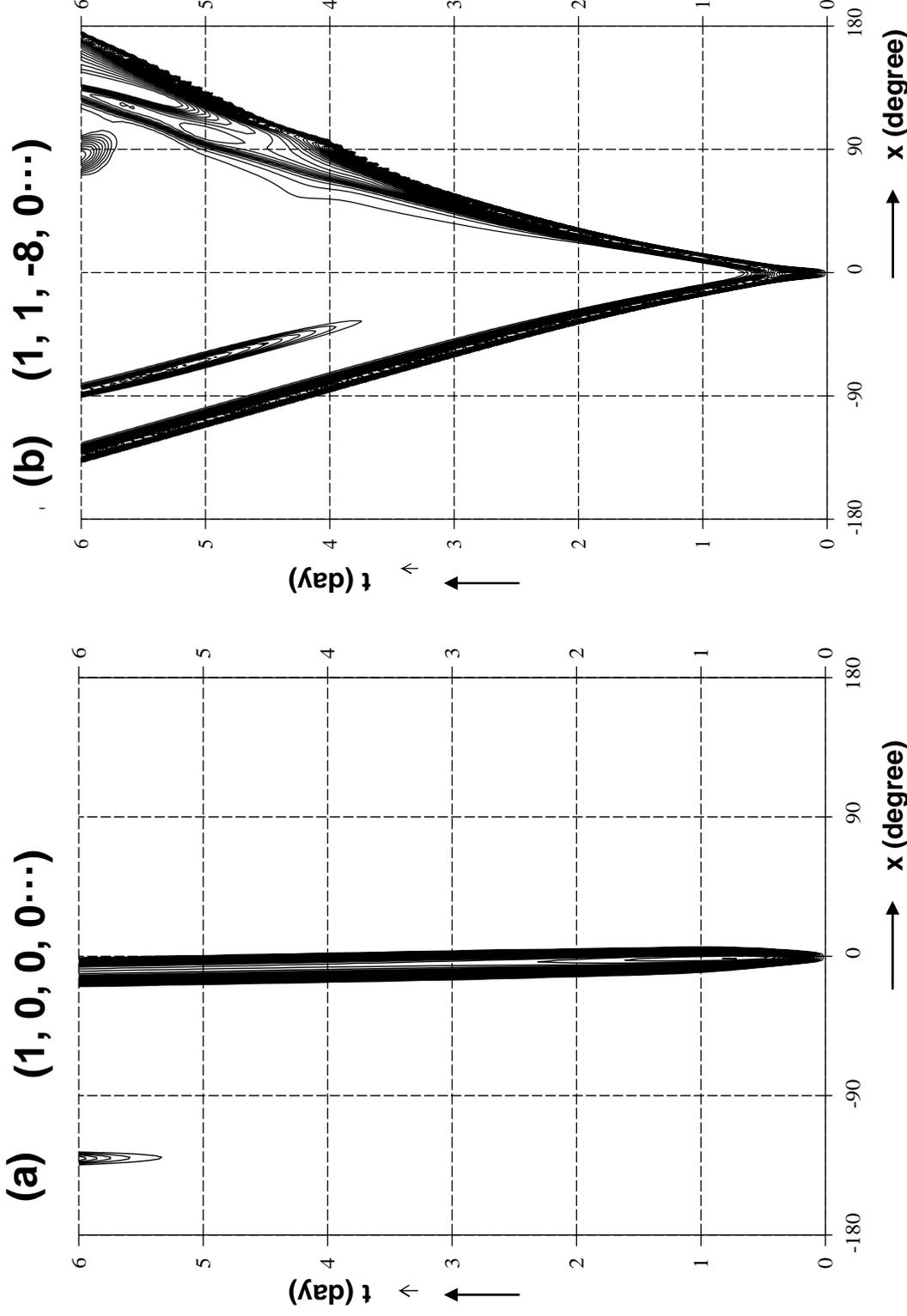
(0,1,0) is **NOT** essential

Two propagation properties of growing disturbances

< A slowly WP D. and oppositely-propagating (WP and EP) Ds.>

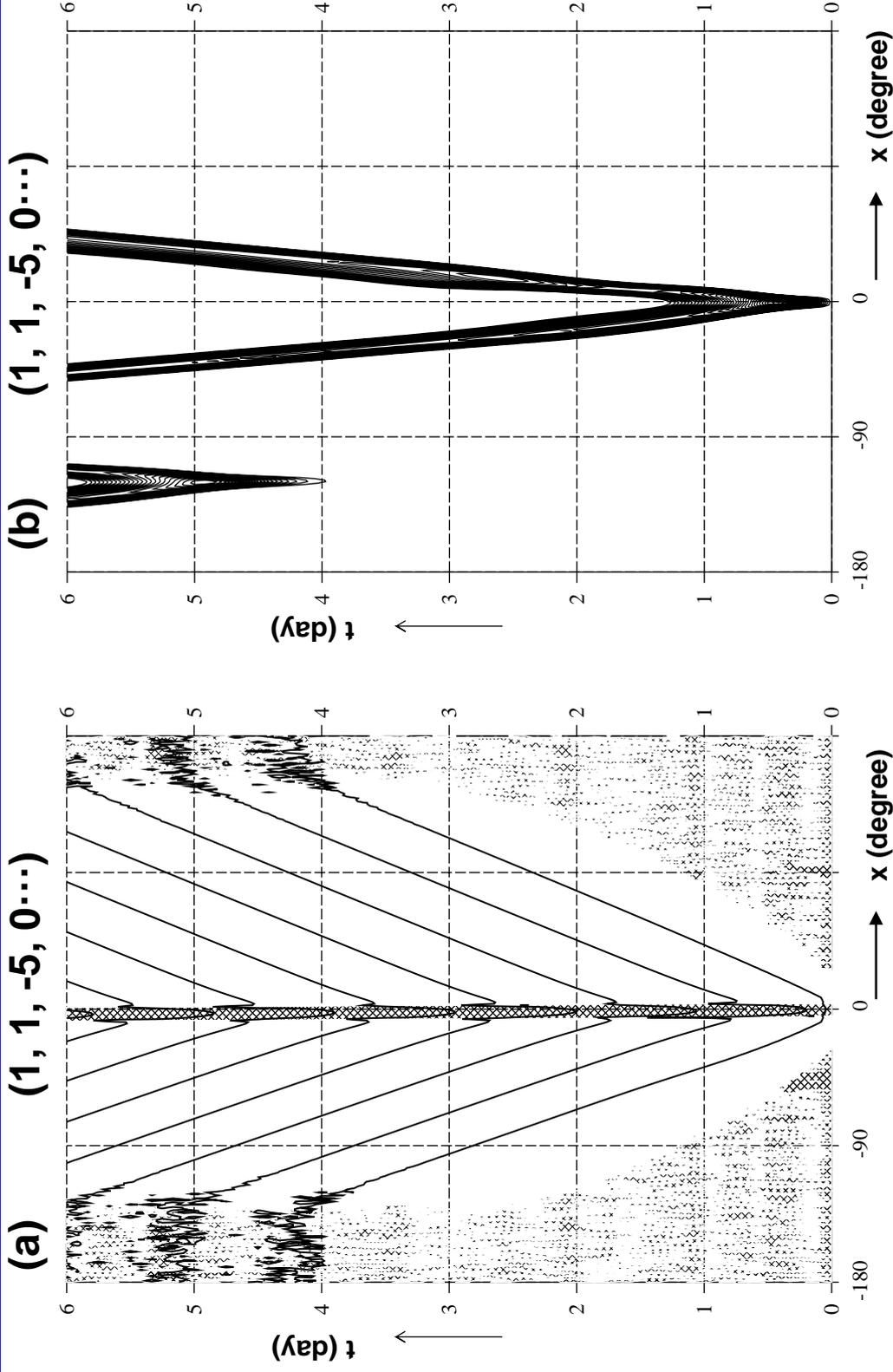


(3)' Simple nonlinear model (32 layers; $N, C_s = \text{constant}$) ($W_{3.5\text{km}}$)



- * Comparison of linear ($\epsilon_1 = 0$) and nonlinear ($\epsilon_1 = 1$) cases:
 - These two cases agree with those in the linear cases about the propagation.
 - **Amplitude saturation** and **additional developing disturbances** are found
- ← Not found in the linear cases

Linear and nonlinear cases of $(1, 1, -5, 0 \dots)$



- Linear : a slowly WP D.

- $\log_{10}|w|$

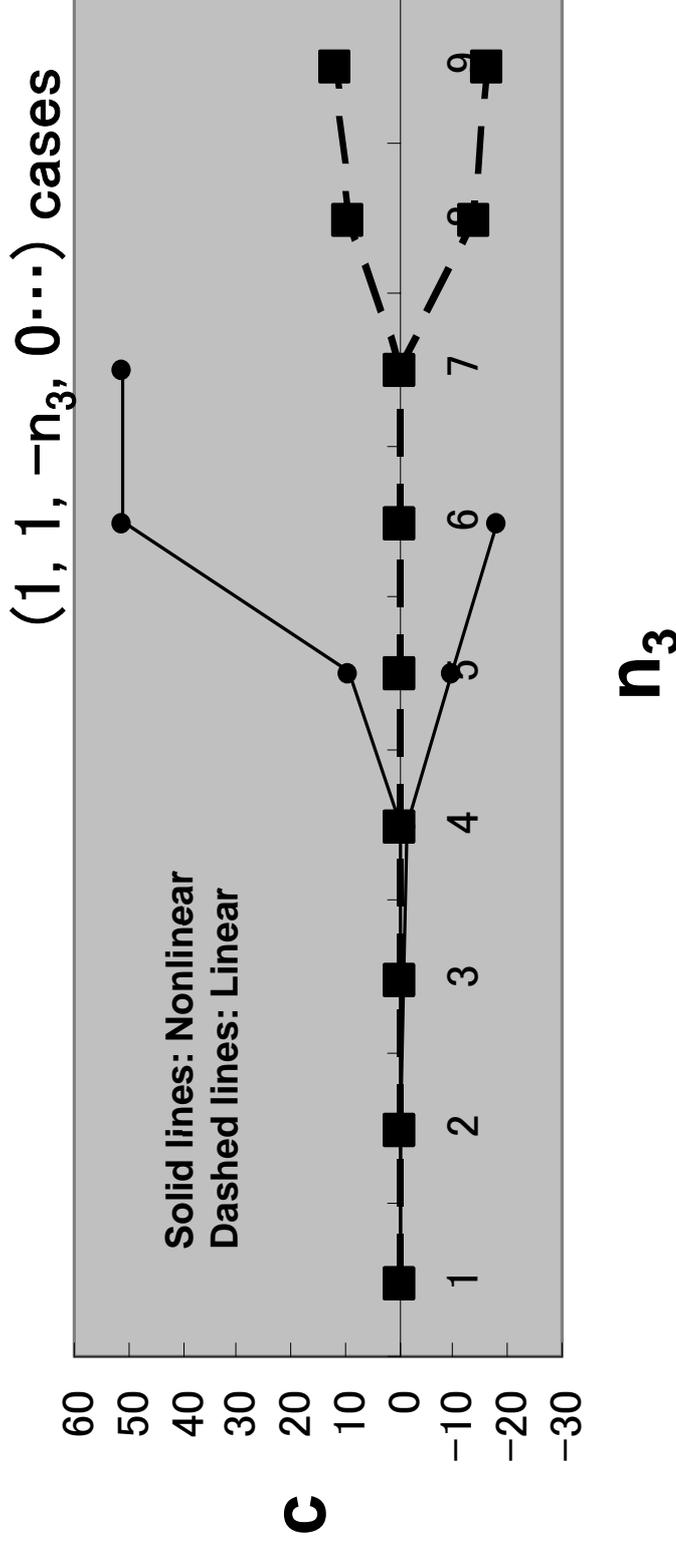
- Nonlinear : a pair of WP and EP Ds.

- w

Propagation property may change between linear and nonlinear cases.

Dependence of propagation property on n_3

Among $(n_1, n_2, -n_3, 0 \dots)$, the values of “ $-n_3$ ” are essential to determine the propagation property of disturbances.

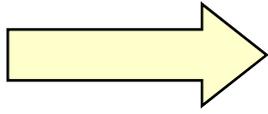


Threshold values between non-propagation and propagation are different between linear and nonlinear cases. The value in the linear case is larger.
→ Propagating disturbances are likely to occur in the nonlinear cases.

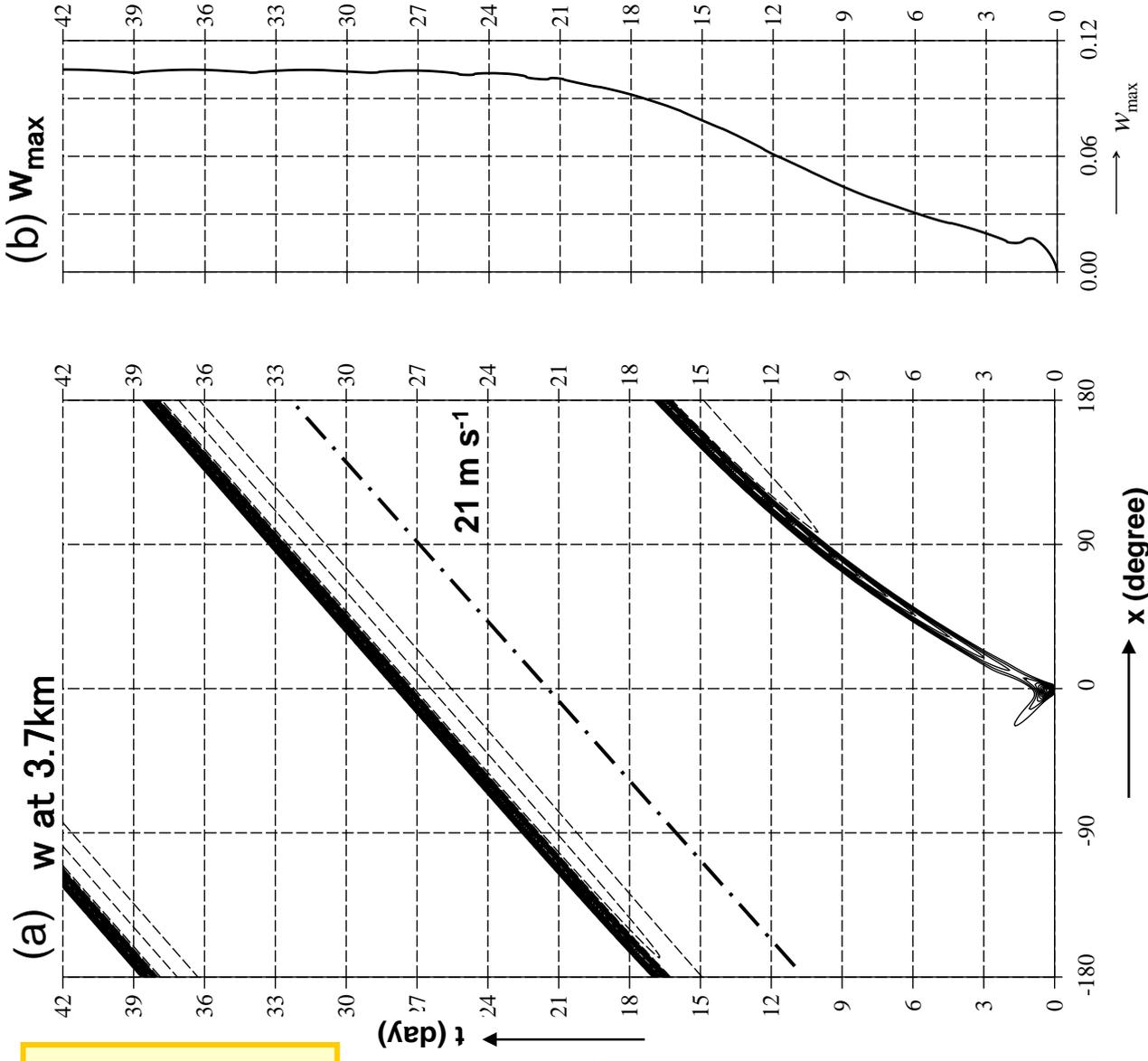
(4) Full nonlinear model

(grid location = NICAM, and NICAM outputs are given as environmental fields, i.e., N, Cs etc.)

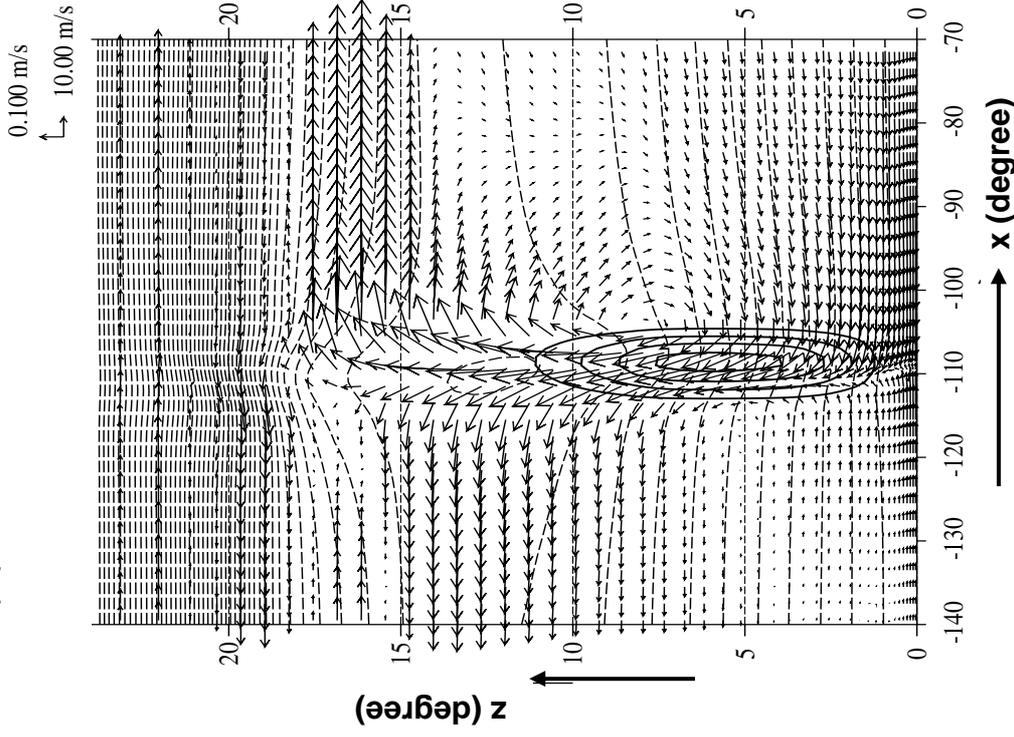
- Vertical grid: 54 layers, variable
- Environmental fields obtained by NICAM
- $\varepsilon_1 = \varepsilon_2 = \varepsilon_3 = 1$



- Propagation speed and amplitudes are synchronized, and attain constant values
- Propagation speed
 $\sim 21 \text{ms}^{-1}$
NICAM $\sim 18.5 \text{ms}^{-1}$

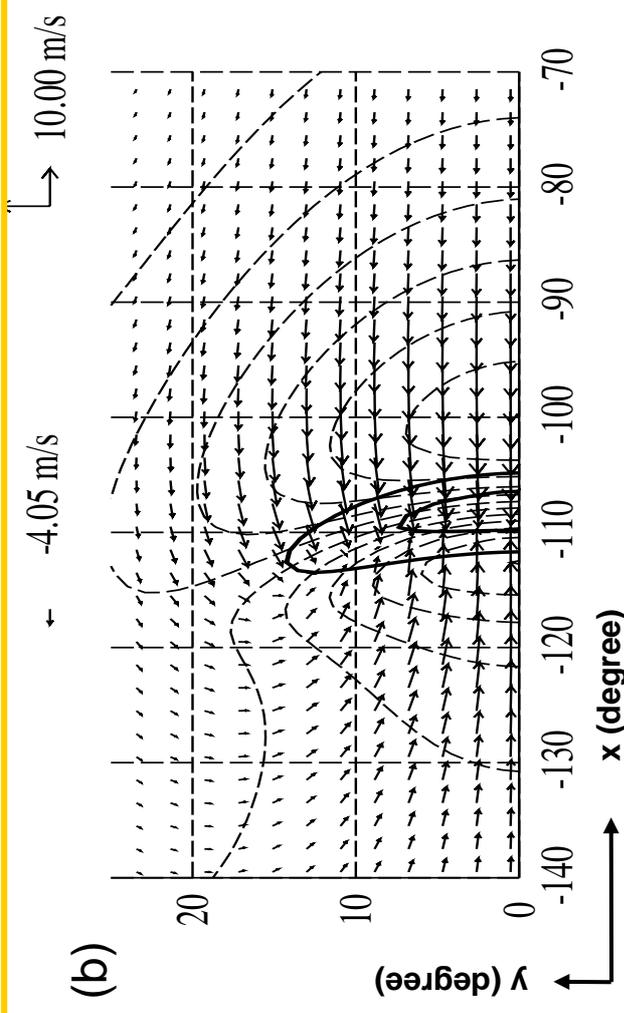


Structures of disturbances



- Vertical structure
 - Vertically erected structure
 - The tops of upward motions are different on the eastern and western sides of heating

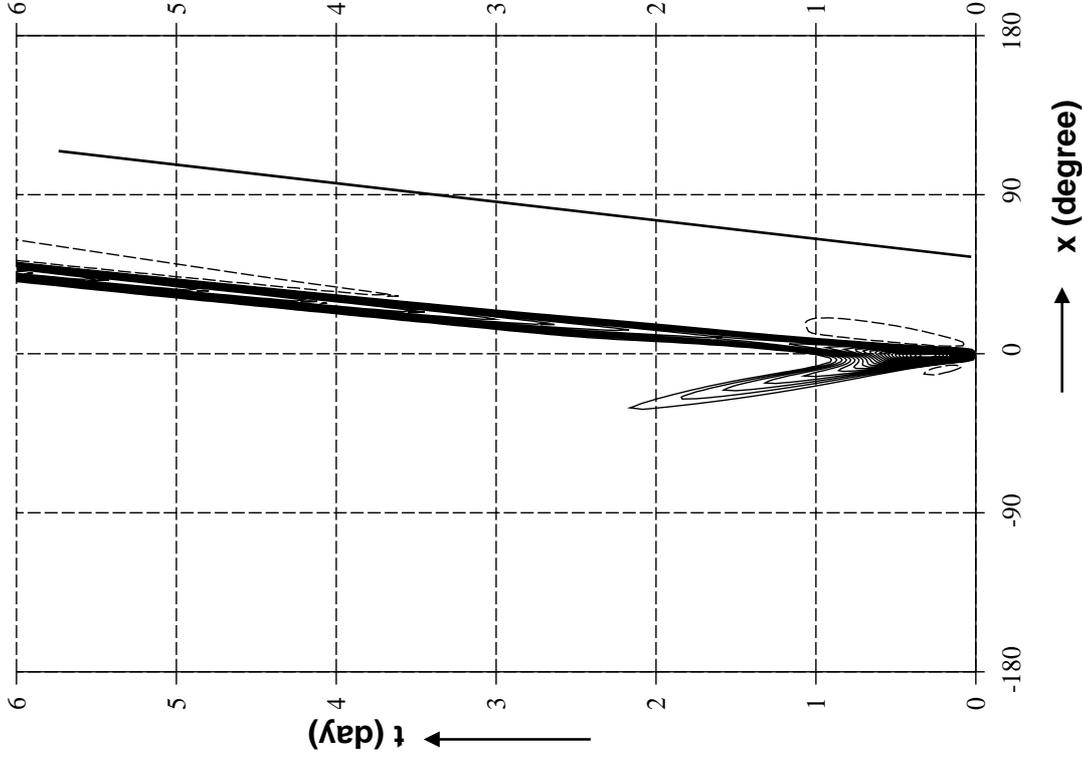
- Horizontal structure
 - East of heating ; Kelvin wave-like features
 - West of heating ; Rossby wave-like features
 - off-equatorial vortical core
 - **Gill response pattern**



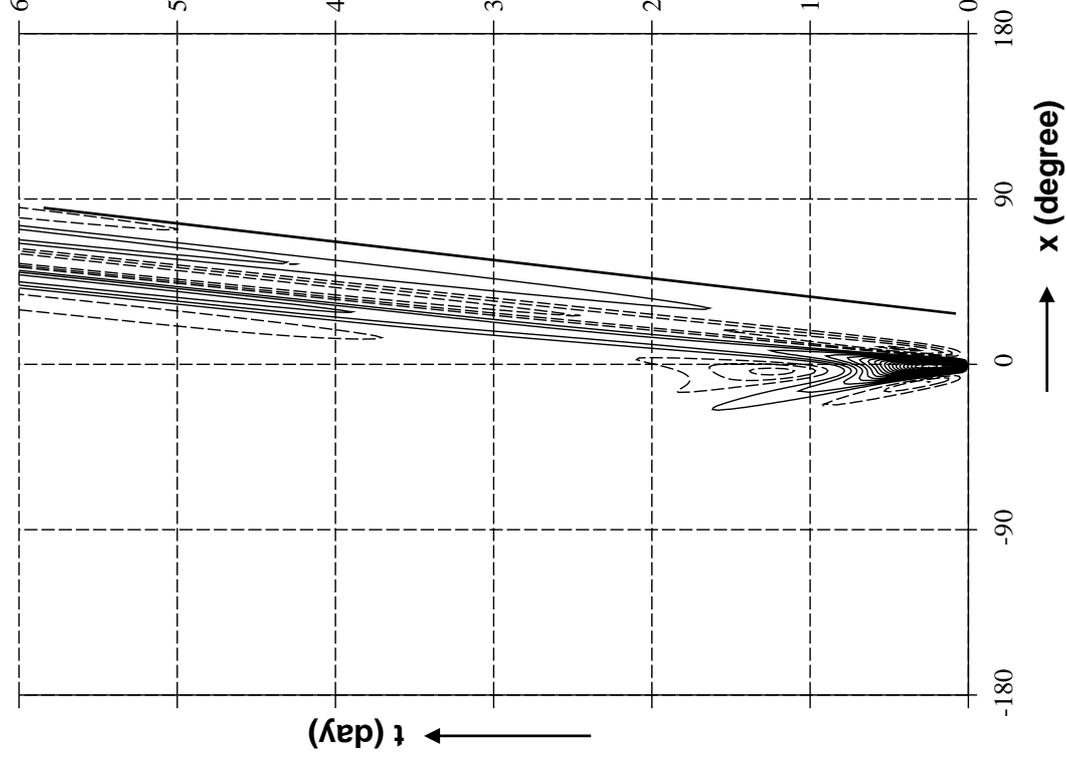
- Slanted structures are observed frequently. → preconditioning.
- Some techniques, such as time-lag between w and Q are needed.

Do equatorial waves play some roles on EP SCCs ?

(a) Positive-only wave CISK



(b) Ordinary wave CISK

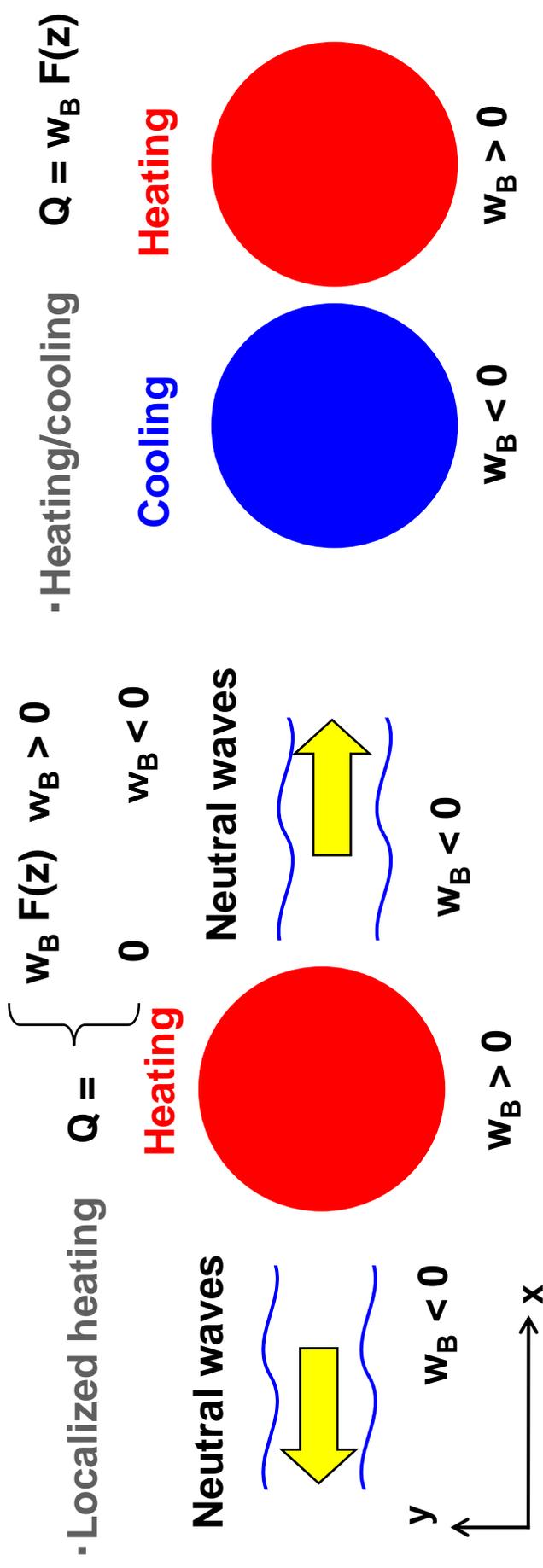


Solid lines : upward motion, dashed lines : downward motion



Propagation property between positive-only and ordinary wave CISK is same.

No role of equatorial waves for EP mechanism



Localized heating produces neutral waves and induces the response patterns outside of heating

In ordinary wave CISK, neutral waves are **NOT** excited.

>> Equatorial waves (= neutral waves) play **NO ROLE** on EP mechanism of SCCs.

>> **Convection** is **PRIMARY**, and neutral waves are secondary.

>> **Convection** always excites **neutral waves** outside it, while **neutral waves** sometimes induce convection.

Present summary

<Assumptions used in this study>

- Heating : positive-only wave CISK
- Large values of horizontal viscosity/diffusion

(1) In the linear cases, the propagation property of growing disturbances is determined by the sign of baroclinic modes of heating at the cloud base. Due to **positive-only wave CISK** and **equatorial beta effect, preferred EP disturbances** were attained.

(2) Using the positive-only wave CISK, **preferred EP disturbances have asymmetric horizontal structure**: Kelvin wave-like features on the eastern side of heating and Rossby wave-like ones on the western side, i.e., **Gill response pattern**.

* Using positive-only wave CISK, neutral equatorial waves are excited by the localized heating and the response pattern is formed around the heating.

(3) For the EP mechanism of disturbance induced by positive-only wave CISK, **neutral equatorial waves play no role**.

* EP mechanism produced by positive-only wave CISK is completely different from the propagation property in the neutral waves related to the dispersion relation.