

# A comparative experiment of warm rain bin schemes using a kinetic driver

## 暖かい雨に関する ビン法雲微物理モデルの比較実験

RICO

RF-09

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Fields of trade wind congestus  
typical cloud base 600 m  
typical cloud top 2000-3000 m

# I . Objectives: To improve bulk parameterization scheme

## Target : bulk scheme for warm rain (boundary layer clouds)

Boundary layer clouds are important for climate study.

Bulk parameterization schemes are used in large-scale models.

However, there are many ambiguous parameters in the schemes.

➔ To improve the bulk schemes, we will use the results of bin scheme models.

## Model : CReSS with Kuba-Fujiyoshi bin model

Case : [RICO](#) (Bermuda area, trade wind cumulus) model intercomparison case  
↑ GCSS (GEWEX cloud system study)

Results : layered structure (average potential temperature etc.), is well simulated, however, ...

## Problems and what to do :

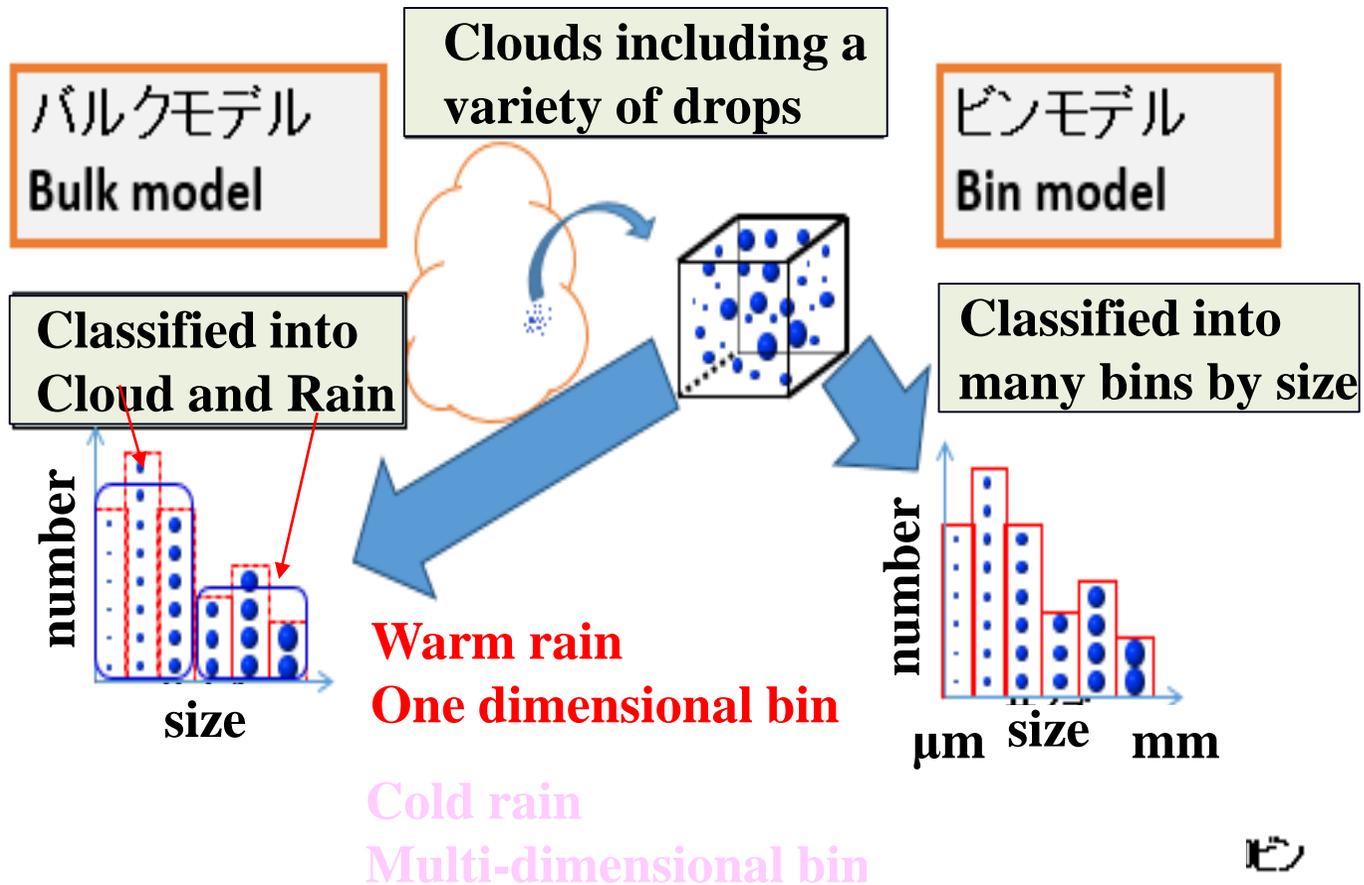
There are many differences between the results of bin models.

Also in bin models, there are many ambiguities.

Improve the bin model by intercomparison experiments using **KiD (Kinematic Driver for microphysics intercomparison)** and comparison of the results with observations.



# What occurs in clouds and what we are modeling



## Processes included

- 1. Activation of aerosols**  
maximum value of super saturation determines the number of activated aerosols
- 2. Deposition growth**  
the difference of the basic equations is whether the integration of time is used or not.
- 3. Collision-coalescence**  
whether the kernel of the collision includes the effect of turbulence.

We used 71 bins whose ratio of mass between adjacent bins is  $2^{1/2}$ . However, for comparison, the results with 34 bins will be presented.

The details of the differences of the bin models will be omitted.

# II . Inter comparison experiments ( LES )

**RICO : Bermuda area, trade wind cumulus  
Case of GEWEX Cloud System Study**

<http://www.knmi.nl/samenw/rico/>

## Grid

$\Delta x = \Delta y = 100\text{m}$ ,  $\Delta z = 40\text{m}$

**Domain Size** 12.8km×12.8km× 4.0km

**Number of Grids** 128 x 128 x 100

**Geostrophic Wind ( cyclic lateral B.C. )**

**u** : constant vertical shear  $2 \times 10^{-3} \text{ s}^{-1}$

**v** : uniform

pot. temp., water vapor

**Bottom surface B.C.**

SST: 299.8K  $\Delta T = 0.6^\circ\text{C}$

**Large scale forcings**

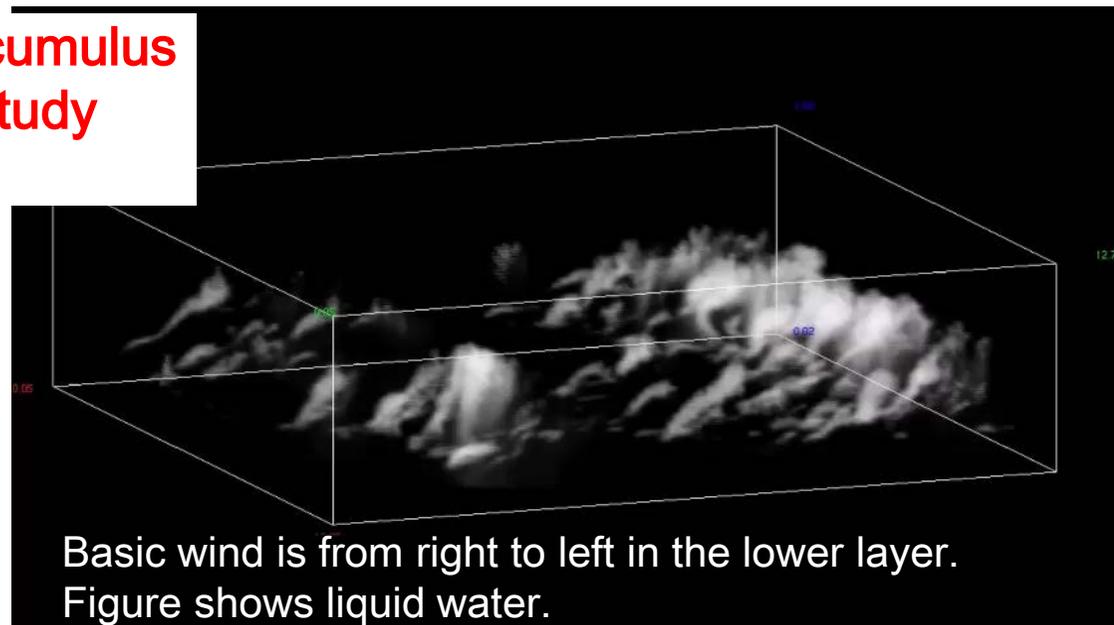
Subsidence

$z > 2260\text{m}$  :  $w = -0.005[\text{m/s}]$

$z < 2260\text{m}$  : Constant Divergence

Horizontal advection (depending on  $z$ )

**Duration** : 24hours (last 4 hrs are analyzed)



## Cloud physics models

**1 moment bulk**

**2 moment bulk**

**bin**

MESO-NH

DALES

RAMS

SAM

UCLA

@NOAA

JAMSTEC

WVU

SAMEX

Utah

COAMPS

DHARMA

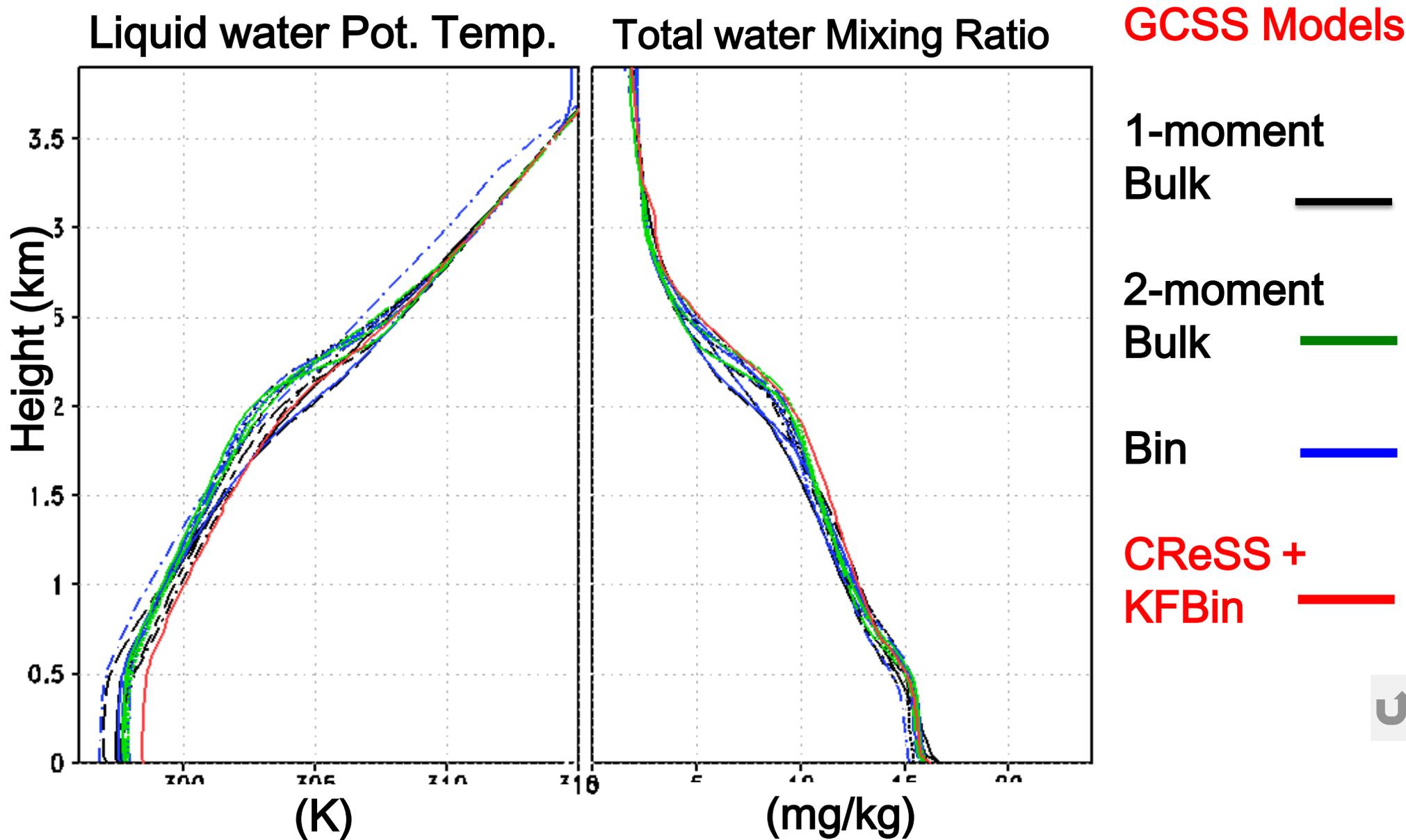
EULAG

UKMO

2DSAM

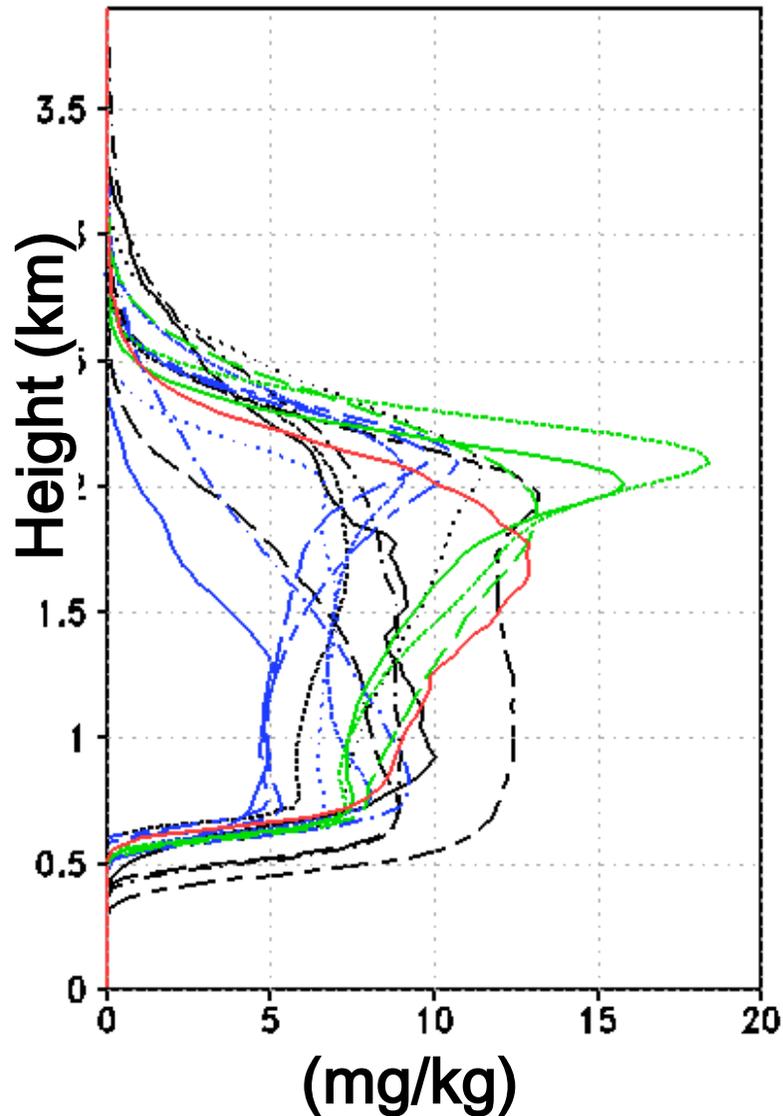
RAMS

# II-2. Averaged vertical profiles (20-24 hr)

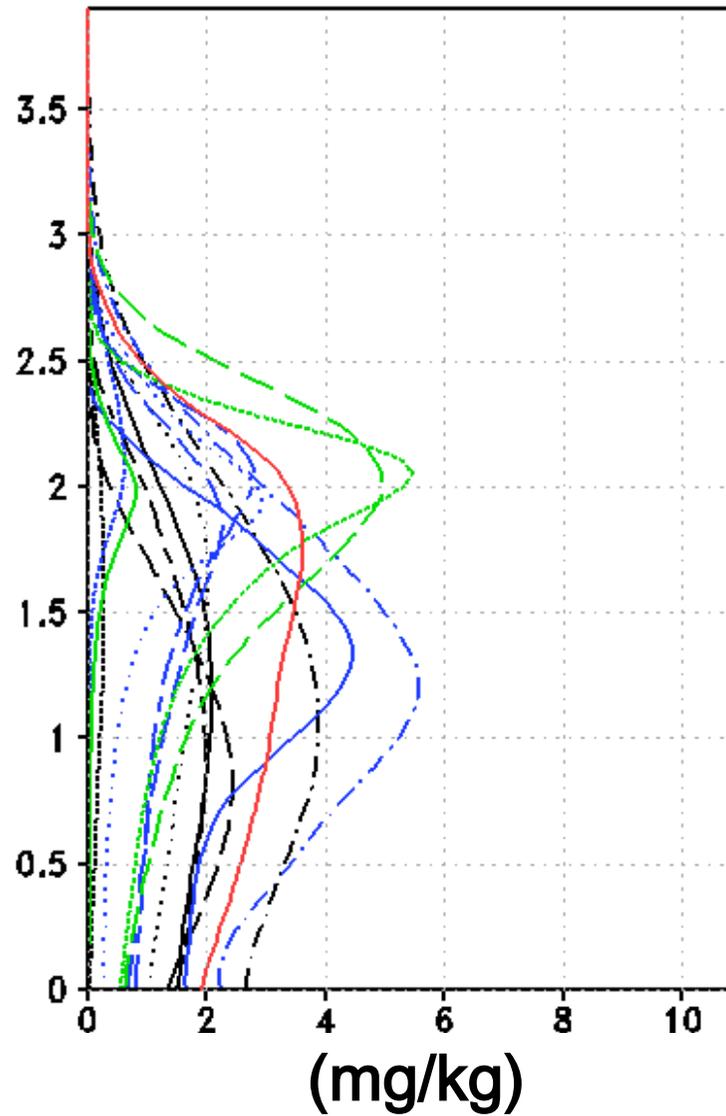


# II-3. Averaged vertical profiles (20-24 hr)

## Cloud water Mixing ratio



## Rain water Mixing Ratio



**GCSS Models**

1-moment

Bulk



2-moment

Bulk



Bin



**CReSS +**

**KFBin**



Cloud :

$r \leq 47.9 \mu\text{m}$



# II-4. Cause of the difference and KiD

## Cause of the difference

### Dynamical model

basic equation

advection scheme

diffusion scheme

subgrid-scale scheme

### Cloud physical model

classification ( cloud, rain etc. )

number of variables (moments)

equation for each process,

activation, ( initial cloud distribution

deposition, conversion,  $v_T$  etc

⇒ To examine what produces the difference, we use KiD.

**KiD** Kinematic Driver for  
microphysics intercomparison  
( Shipway, Met. Office )

Common setting for dynamic  
process(velocity fields, initial T,  $q_v$   
are given) advection scheme : ULTIMATE

### Cloud Physical model

We can put our own model.

Following examples can be used.

Thompson07, 09, Morrison(2B),

**Tel-Aviv University(TAU, bin)**

### CASES

**1D(GCSS-Warm**, Mix, Deep),

2D(Cu, St, squall line)

WMO-workshop cases.

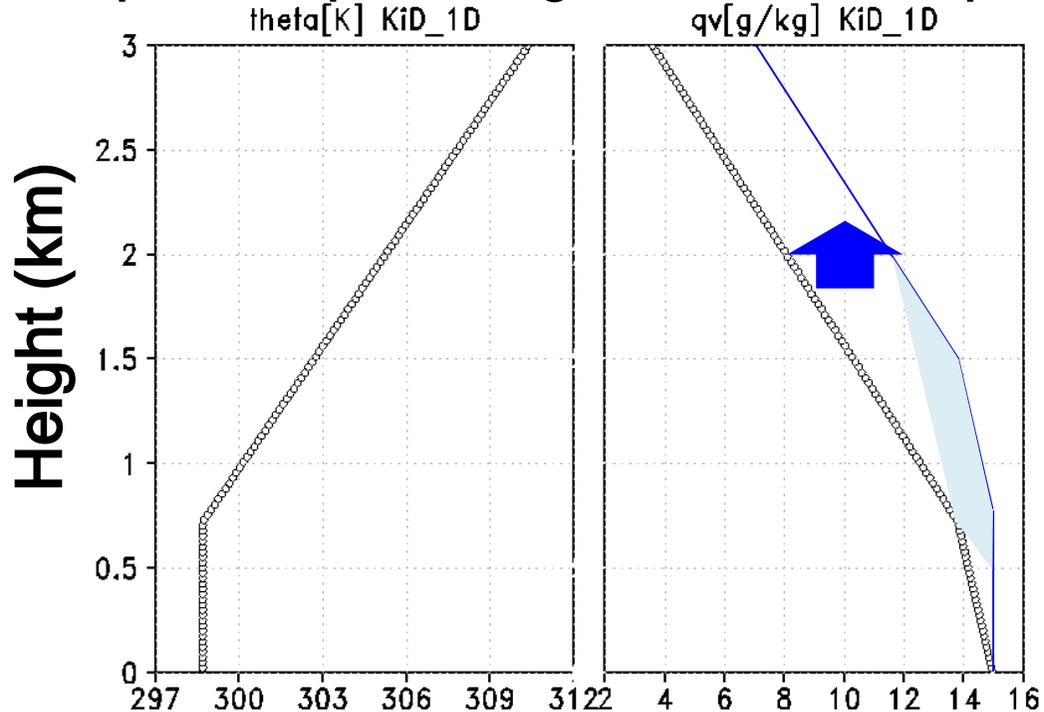
+ **2DResults of CReSS for RICO**

# III. KiD(Kinetic Driver)

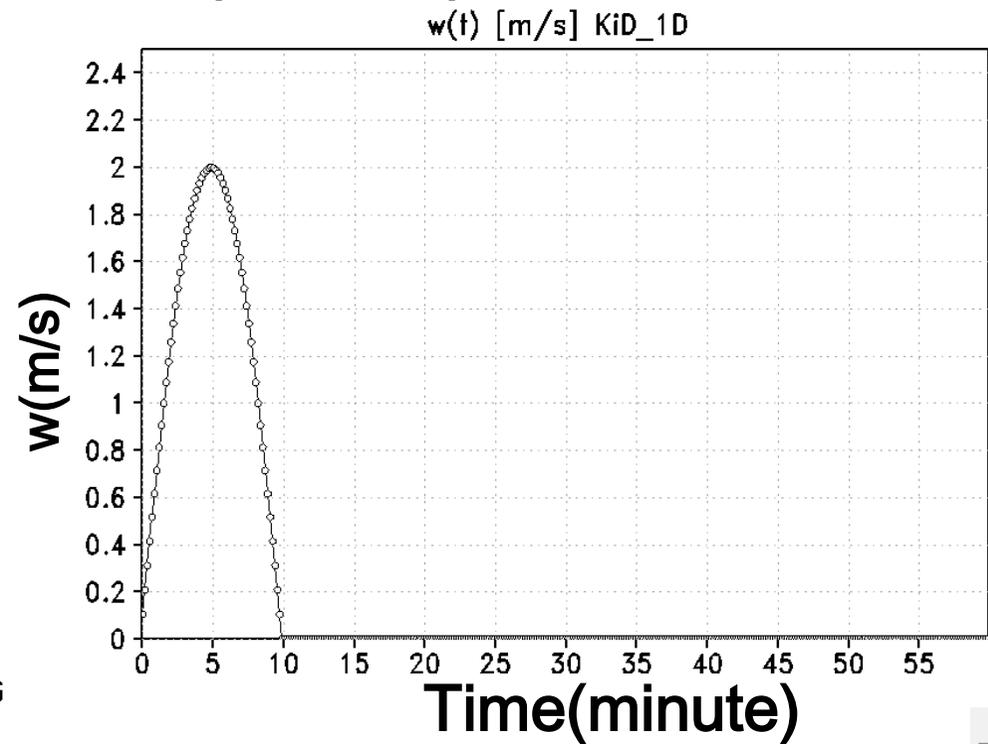
## III-1. 1 dimension (1D、warm1)

Initial profiles

pot. Temp. mixing ratio of water vapour



Time change of  $w$ , which does not depend on position.

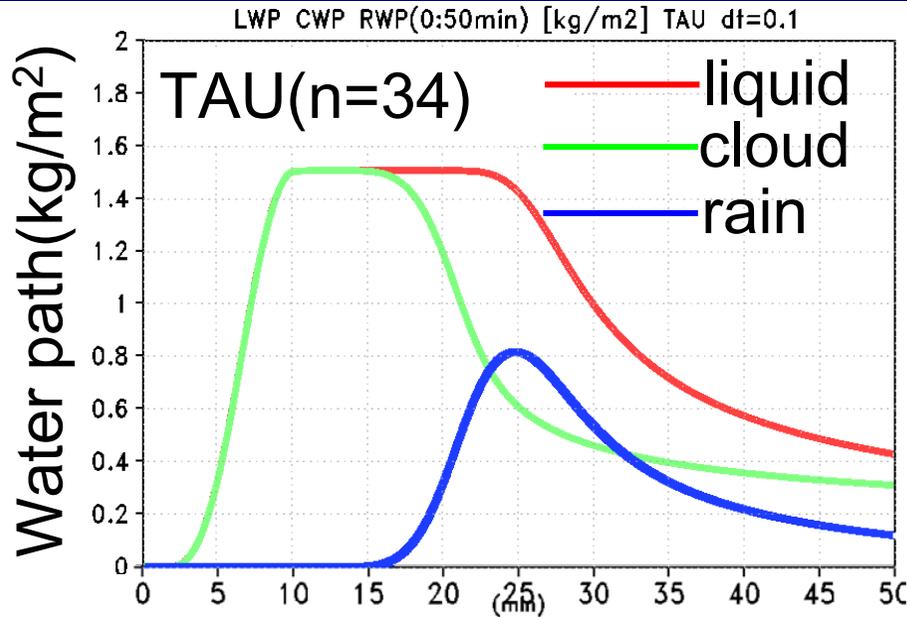


Because of rising, the air is saturated.

In this case, the temperature is kept fixed in spite of the condensation, for simplicity.



# III-1 ' KiD-1D Result. Time change of water path



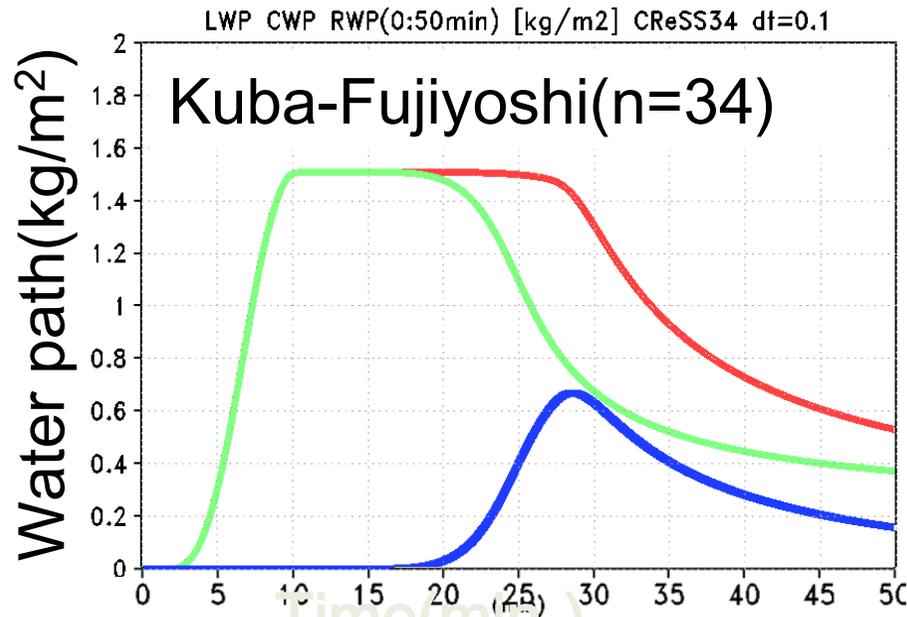
**Difference between TAU and KF**

KF: originally n=70.

n=34 is used for comparison.

(The ratio of the masses of the adjacent bins is 2.)

**KF: conversion from cloud to rain is slow, especially with n=70.**



# III— 2 2D experiment (flow pattern resulted in CReSS)

Several 2D experiments are performed using simple flow pattern.  
The results are reasonable.

Then, we used the **flow pattern resulted in 2D CReSS simulation.**

**Grid:**  $\Delta x = 100\text{m}$ ,  $\Delta z = 40\text{m}$

**Time interval:**  $\Delta t = 1\text{sec.}$ ,  $\Delta t \text{ bin} = 0.1\text{sec.}$

**Domain Size**  $12.8\text{km} \times 4.0\text{km}$

**Number of Grids**  $128 \times 100$

**Geostrophic Wind ( cyclic lateral B.C. )**

$u$  : uniform     $v$  : uniform

pot. temp., water vapor : RICO case

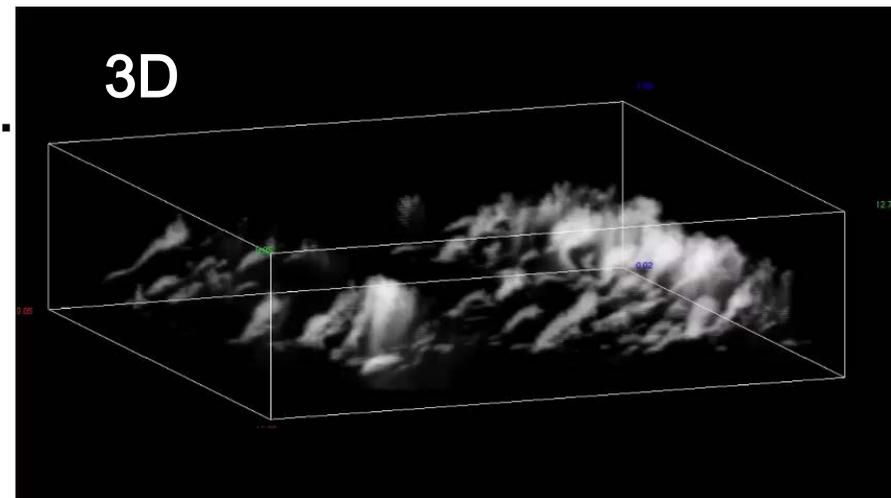
**Bottom surface B.C.**

Fix surface flux.

**Large scale forcings: same as GCSS**

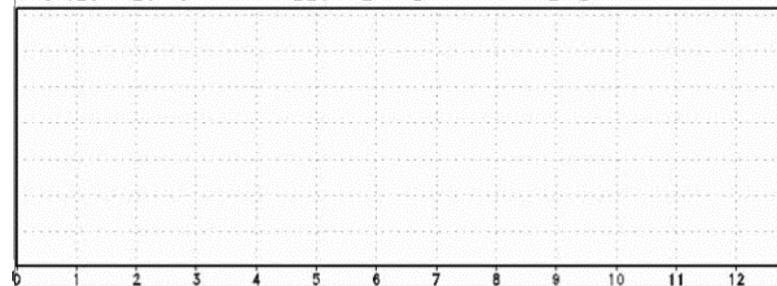
**Duration : 24hours**

**We stored flow fields every 1 sec.**



**2D**

$q_l(\text{g/kg}), q_c = 0.01[\text{g/kg}, W] \text{ rh} = 100[\text{B}] \text{ t} = 0:00:00$



# III— 2 2D experiment (flow pattern resulted in CReSS)

Several 2D exper  
The results are re  
Then, we used th

**Grid:**  $\Delta x = 100$

**Time interval:**

**Domain Size**

Number of (

**Geostrophic V**

**u** : uniform

pot. temp.,

**Bottom surface**

Fix surface

**Large scale fc**

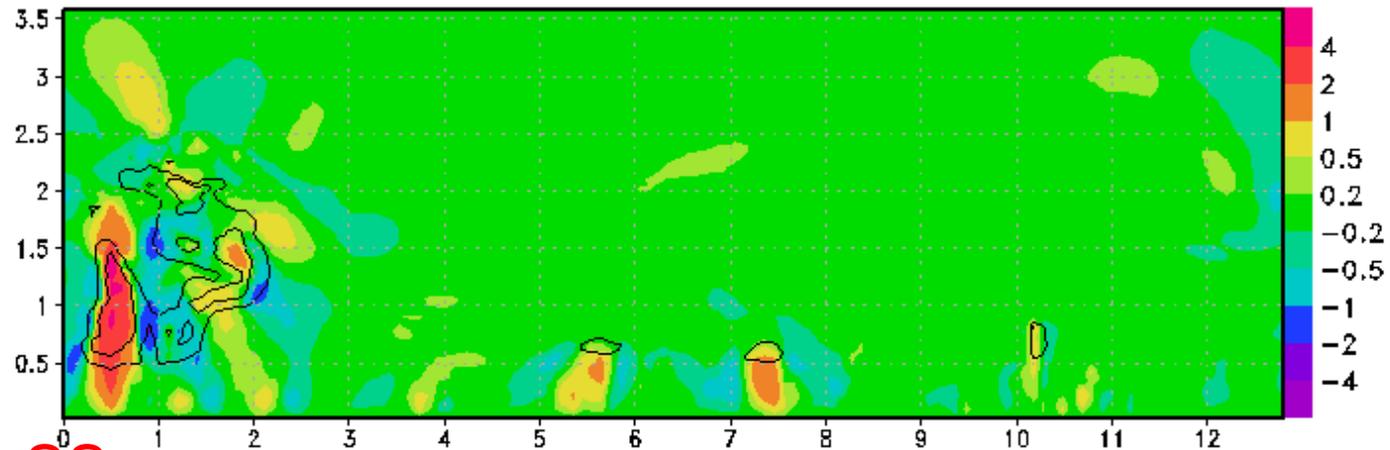
**Duration : 24h**

**We stored flow**

**CReSS**

**W**

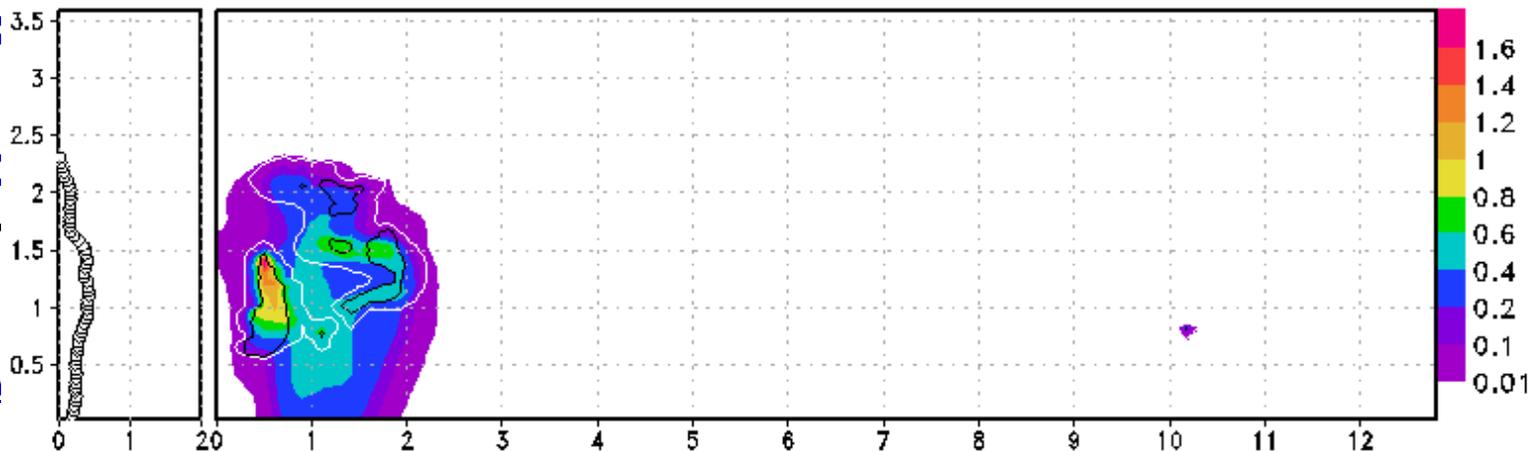
w(m/s) rh(95,100%) t=22:15:00



**CReSS**

**ql**

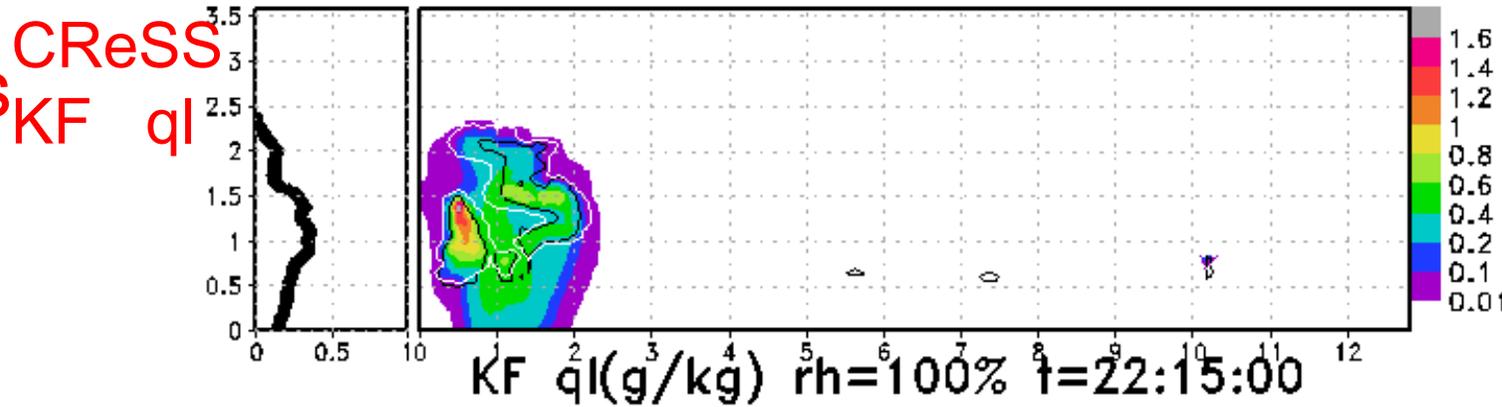
ql(g/kg), qc=0.01[g/kg,W] rh=100[B] t=22:15:00



# III— 2 2D experiment (flow pattern resulted in CReSS)

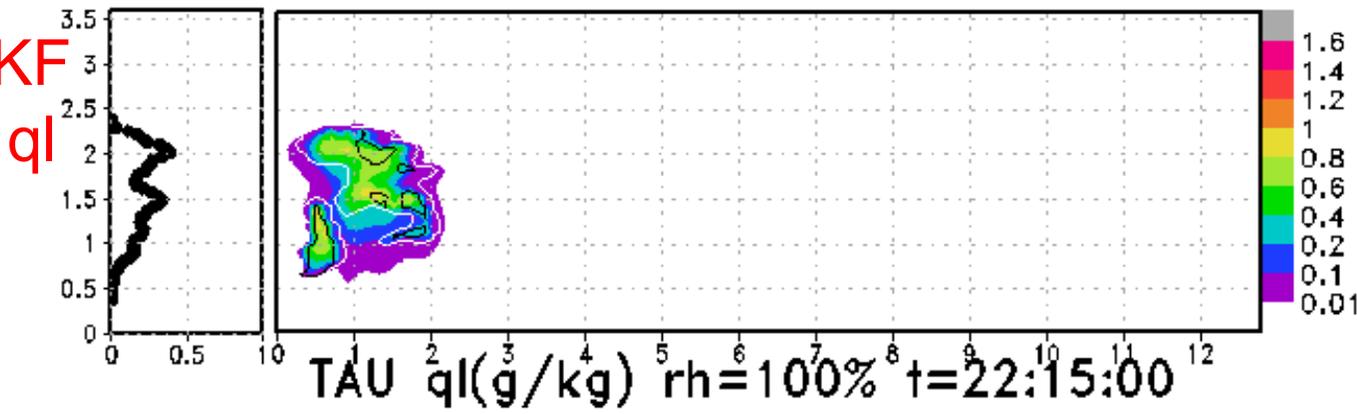
Flow fields in CReSS  
 advection  
 + cloud physics  
 KF and TAU scheme

CReSS ql(g/kg) rh=100% t=22:15:00



Liquid water mixing ratio  $\Rightarrow$   
 upper : CReSS+KF  
 center : KiD+KF  
 lower : KiD+TAU

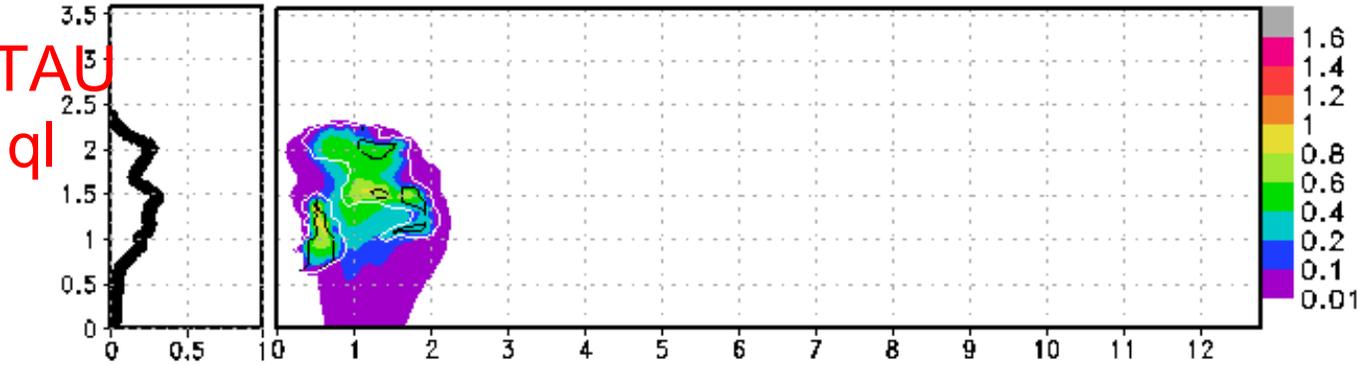
KiD KF



TAU ql(g/kg) rh=100% t=22:15:00

Left : average in cloud  
 Right : pattern

KiD TAU



# III— 2 2D experiment (flow pattern resulted in CReSS)

liquid water mixing ratio(left) and cloud amount(right)  
**KF(Red)** and **TAU(Green)**

Average from 20:00~24:00

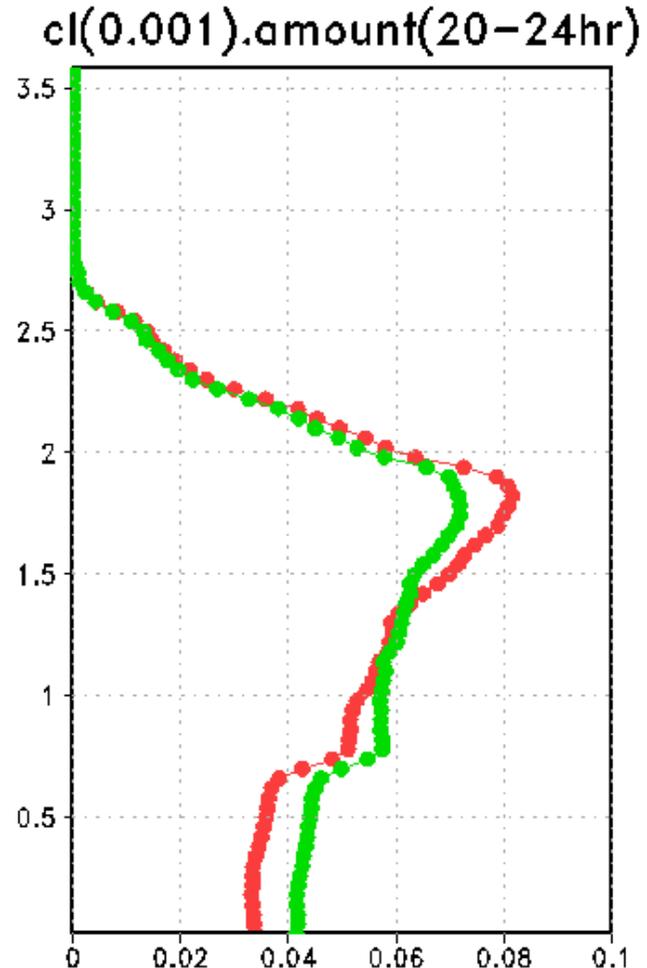
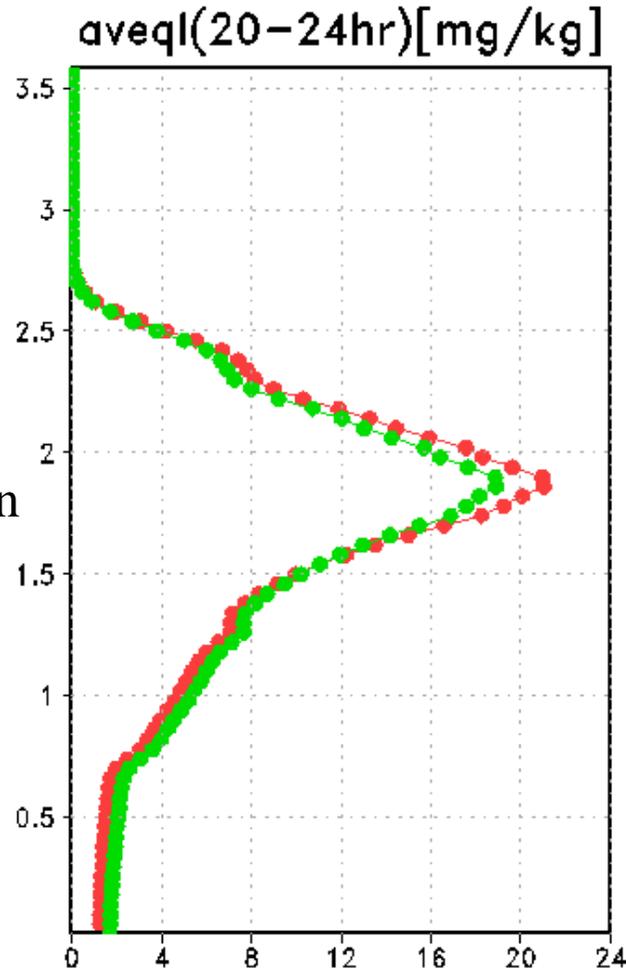
Difference with schemes  
 KF and TAU.

Cloud grid is the grid with  
 liquid water mixing ratio  $>$   
 $0.001$  [g/kg]

More liquid water is stored in  
 clouds with KF scheme,  
 than with TAU scheme.

The precipitation is a little  
 larger in TAU than in KF.

Conv. is slow in KF than in  
 TAU



# III— 3 2D. cloud physical process divided into 3 parts.

Vertical profiles of averaged liquid water mixing ratio (t=20-24hr)

Basic scheme is KF.

The cloud physical process is divided into [3 parts](#); activation, deposition, and collision.

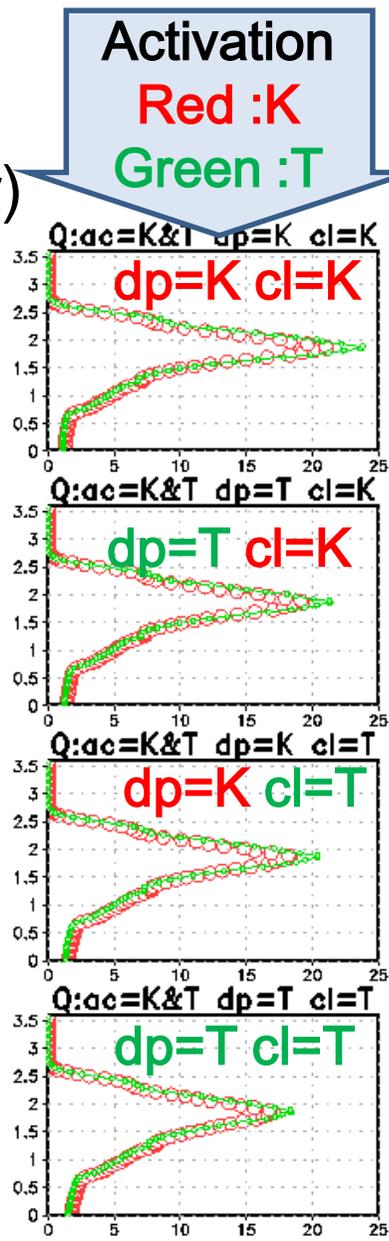
For each process, **K(Kuba-Fujiyoshi)** or **T(Tel-Aviv Univ.)** can be selected.

We performed 8 runs.

ac : activation

dp : depositional growth

cl : collision-coalescence.



# Summary - 1

- A numerical experiment of cumulus convection observed during RICO was performed using a bin microphysical model developed by Kuba and Fujiyoshi.
- 久芳一藤吉が開発したビン法雲微物理モデルを使い、RICOで観測された積雲の数値実験を行った。
- We are improving the model by comparing the results with TAU model and observational results.
- テルアビブ大学のビンモデルや観測結果との比較を通じて、モデルの改良を行っている。
- Especially we compared the results of two bin models by dividing the models into three processes, i.e.,
  1. activation process of aerosols, 2. depositional growth process, and 3. Collision-coalescence process.
- ビンモデルの雲微物理過程を、1. 活性化、2. 凝結成長、3. 衝突併合成長の3つの部分に分けて、2つのビンモデルの比較をおこなった。

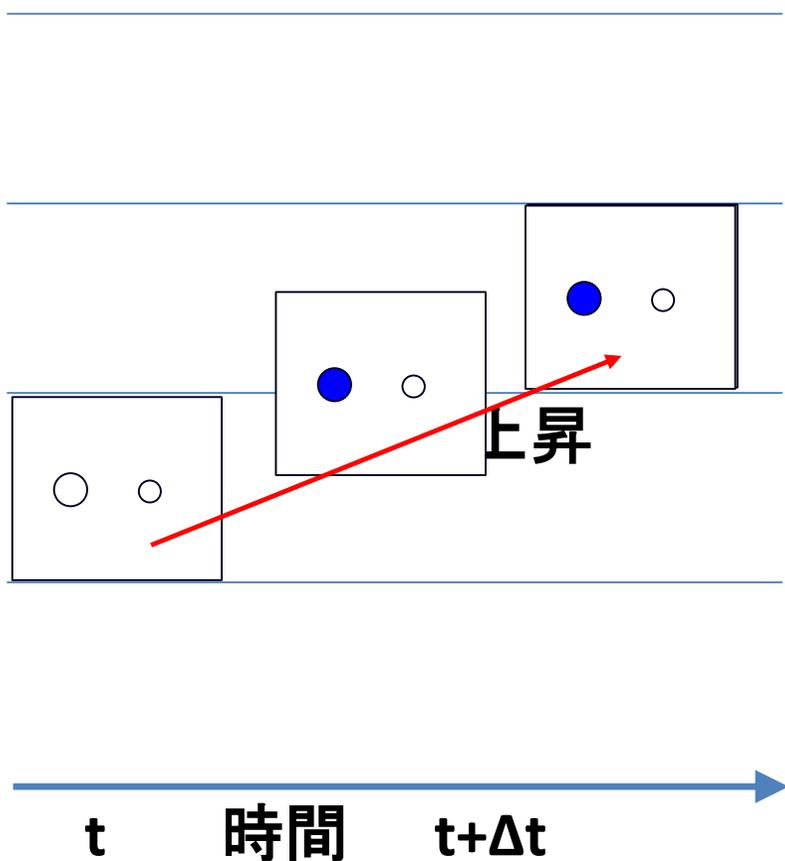


# Summary - 2

- The results show that
  - for the effect of activation process,  $q_l$  in cloud is larger in TAU than in KF.
  - for the effect of deposition and collision processes,  $q_l$  in cloud is larger in KF than in TAU.
  - The surface precipitation is large when  $q_l$  in cloud is small.
- 結果は、活性化では、KFモデルはTAUよりも雲内凝結水が少なく、凝結成長と衝突併合成長では、KFモデルの方がTAUよりも雲内凝結水を多くした。これらの結果、全体としては、KFモデルの方がTAUよりも雲内凝結水を多くした。地表面降水は、雲内凝結水が少ないほうが多くなった。
- We would like to compare the results with some observation, and ...
- 今後、観測結果（粒径分布、凝結水量と降水量の関係など）とも比較して、よりよいものにしていきたい。

# 雲微物理過程－1．初期粒径分布の決め方

## パーセルの上昇と活性化



最大過飽和度が雲粒数密度を決める

よく使われるが、できれば使いたくない式。

$$n_{\text{acti}} \propto \exp(a * (rh - 100))$$

↑ 格子のrh（相対湿度）を使っているか？

単純な方法としては、最小粒径に不足個数分を加える方法がある。が、

活性化される核の数は、十分に細かい格子と時間間隔で考えるべき

大きな粒子が活性化したら、相対湿度は低下する⇒小さな粒子は活性化しない。

⇒KFは、パーセルモデルの結果を使ったパラメタリゼーション（wの関数で総数を決めて、「分布させた。」）

# Time change of size distribution of $q_c$ in growing

Size distribution of  $q_c$  at the maximum  $q_c$  level

**during  $t=101-600\text{sec}$**

Left(TA) , Right(KF)

black :  $dt=0.1\text{sec}$ ,  $dz=25\text{m}$

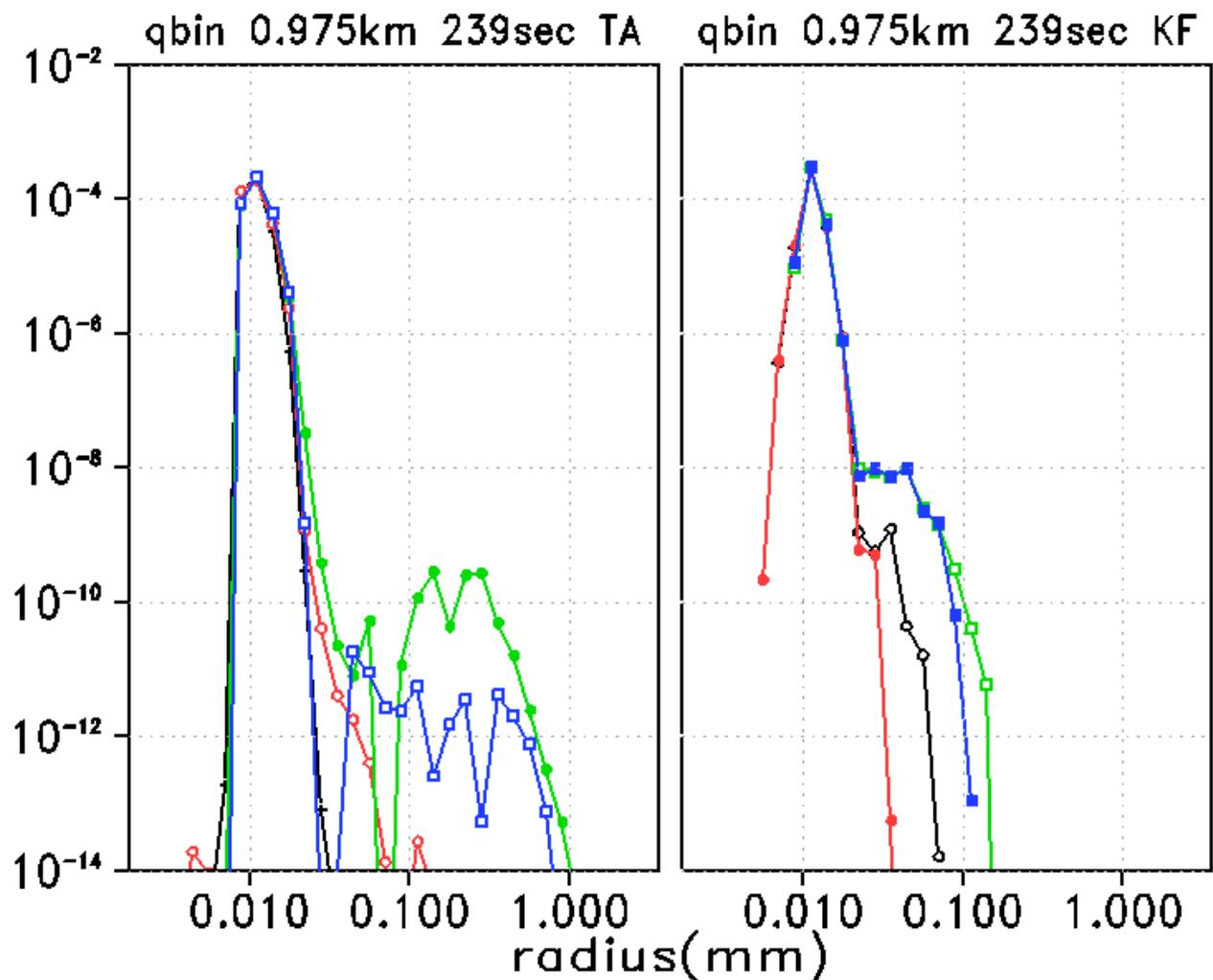
red :  $dt=0.01\text{sec}$ ,  $dz=25\text{m}$

green :  $dt=0.1\text{sec}$ ,  $dz=2.5\text{m}$

blue :  $dt=0.01\text{sec}$ ,  $dz=2.5\text{m}$

**KF**

Initial distribution of  $n_c$  is calculated by the use of parameterization scheme determined by a parcel model.

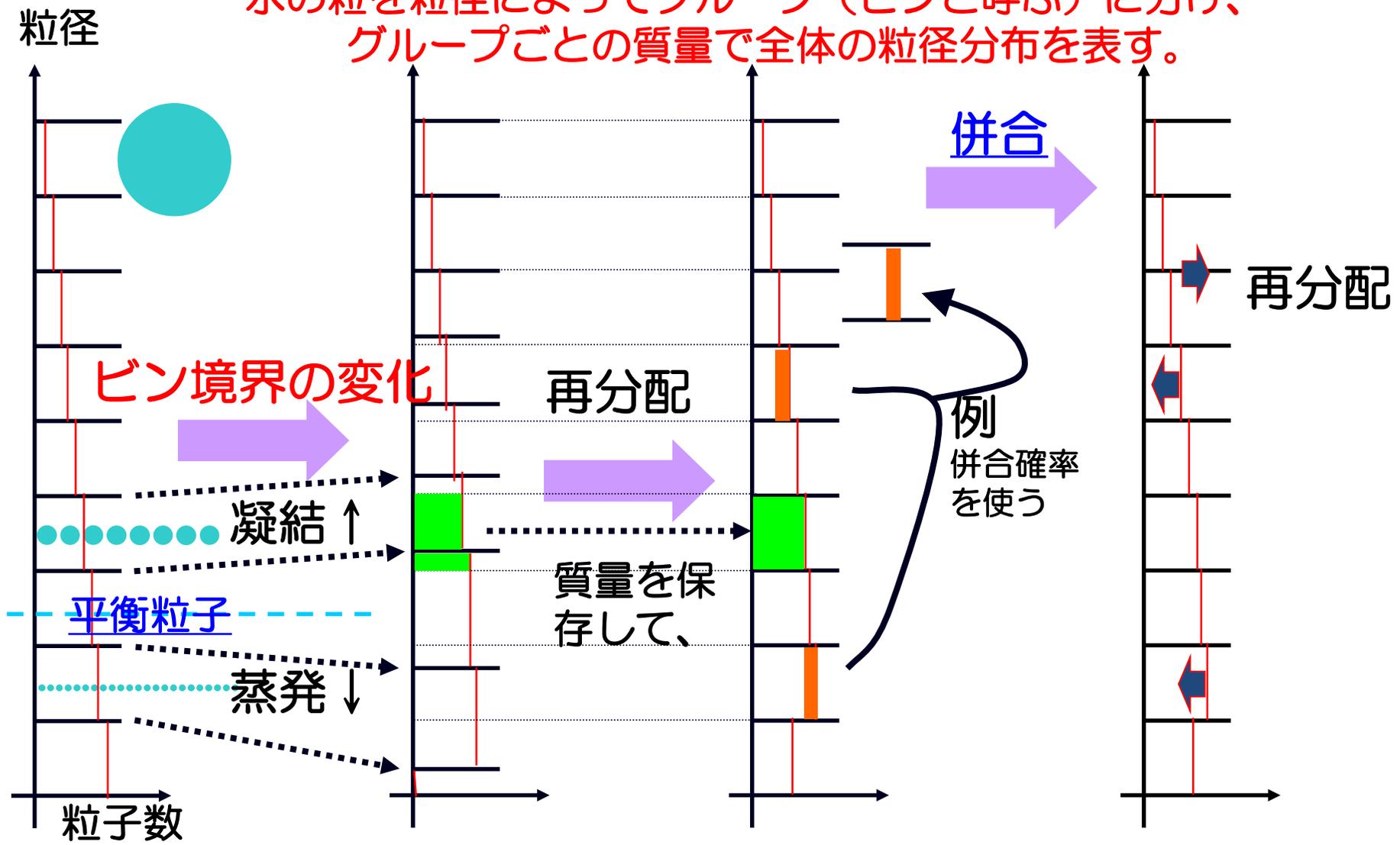


KF: dependency is small.

KFの特徴

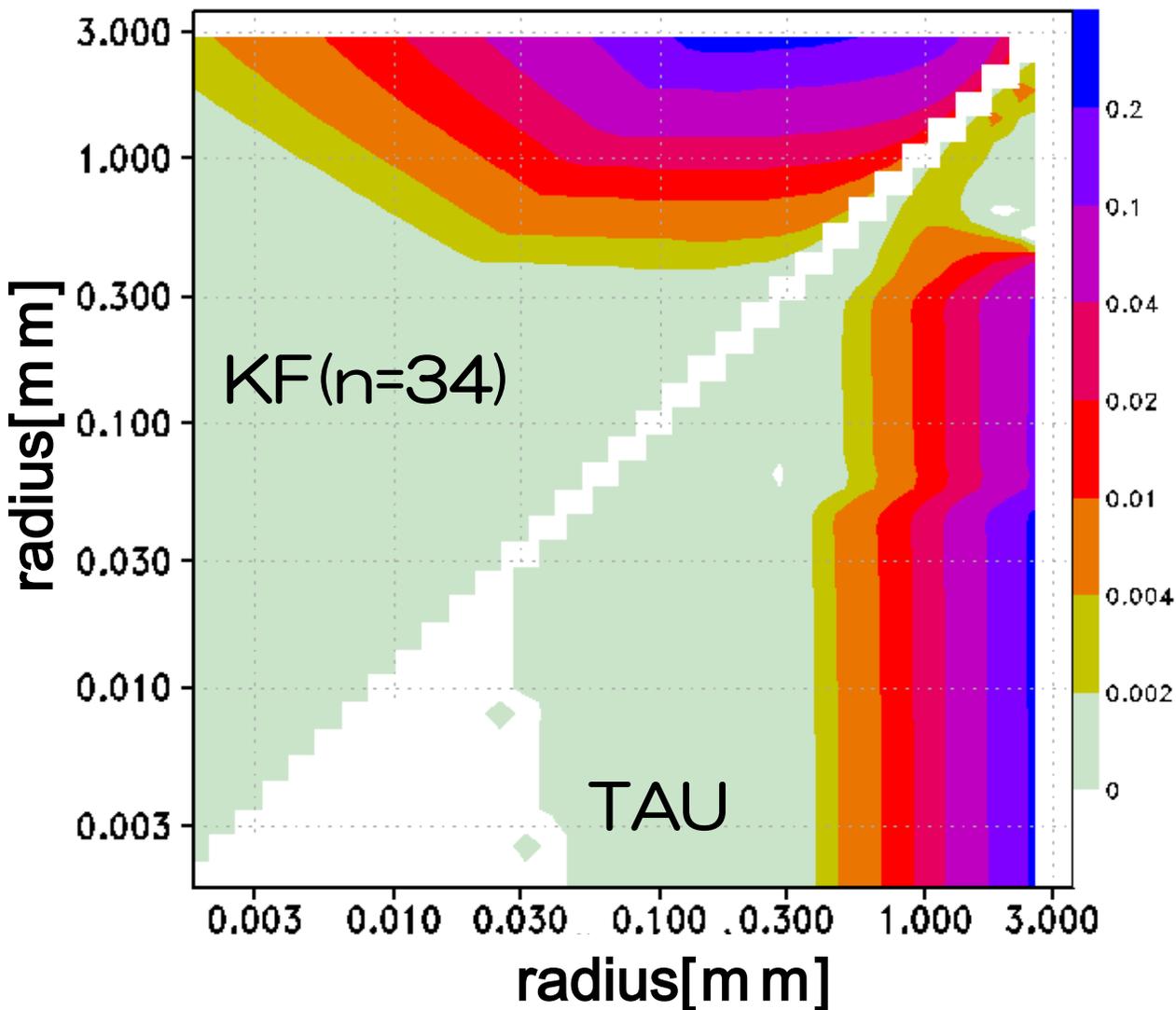
# 雲微物理過程－2. 凝結水の雲物理過程での変化

水の粒を粒径によってグループ（ビンと呼ぶ）に分け、グループごとの質量で全体の粒径分布を表す。



**単純な方法（2変数）**：各ビンの平均粒径を出して、それで丸ごと移動させる。

# Collision efficiency used in KF and TAU



**Situation:** only two kinds of drops exist. Number densities are  $10^2/\text{kg}$  and  $10^4/\text{kg}$ .

**Coalescence probability** during 0.1 sec.

The origins of the difference?

KF34. based on Hall(1980). linear interpolation in each bin. The effects of velocity perturbation is not included.

TAU. Assumption of  $K(x,y)=x+y$  is used. The effects of change in each bin is estimated by assuming a size distribution.

**This does not seem to be related to the rapid decrease in  $n_c$  with height.**

→ **We should use more complicated situation.**