Extended weather and intraseasonal oscillations in the SCS during winter and summer monsoon 2017-2018

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- I. Climate background of the SCSTIMX 2017-2018
- II. The Influences of the MJO & ER on the SCS monsoon onset in June 2018
- III. The Influences of Scale-Interactions Involving the MJO (KWs) on Rainfall Variability over SCS and MC in December 2016 (2017)

**4th International Workshop of the Years of the Maritime Continent, Feb 26-28, 2019,** Institute of Environmental Science and Meteorology (IESM), University of the Philippines, Manila

I. Climate background of the SCS Two-Island Monsoon eXperiment (SCSTIMX) 2017-2018

# Climate background of the SCS Two-Island Monsoon eXperiment (SCSTIMX) 2017-2018



#### Nino3.4 & Southern Oscillation



### Global warming hiatus since 2000

- -IPO
- +AMO
- Hemispheric asymmetric SST distribution
- Frequent CP type El Nino
- Strong extratropical influence

### 2014-2018 ENSO event

- Super El Nino-weak La Nina
- A prolonged ENSO evolution

La Nina in DJF (yr0), weak El Nino in DJF (yr1), IOD in SON (yr1), primary El Nino (yr2), followed by a fast decay to La Nina (yr3)

### 2017-2018

- Weak La Nina phase
- Late SCS monsoon onset, 2018





DJF: La Nina forcing→ weakened Aleutian low, weaker westerly jet, warm SSTA in North Pacific (PMM) MAM: PMM persisted and move southwestward (WES feedback)



# II. The Influences of the ISO and ER on the SCS monsoon onset in June 2018









# I.II Summary

- A Pacific meridional mode (PMM) is evident in the winter-spring 2018, along with an anomalous cyclone (AC) over SCS (DJF) and PS (MAM) during the decaying La Nina phase following the prolonged 2015/16 ENSO
- The AC and PMM are maintained by equatorial heating and extratropical WES feedback in boreal winter monsoon flow
- The AC brought cold-dry northeasterly to SE China coast and SCS in MAM, ISO heating in April and May over IO results in easterly over eastern IO
  → an anticyclonic circulation (AAC) extending from SCS to IO

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• The anomalous AAC, ISO and ER caused a late SCS monsoon onset in June 2018

Q1: PMM and sub-tropical AAC (or AC) over the NW & SCS accompanied with El Nino (La Nina) add to ENSO diversity. Do they also exist in non-ENSO years?Q2: Maintenance of persistent AAC and AC in NE vs SW monsoon flow remains to be further explored other than the feedbacks via WES, IO heating

### III. The Influences of Scale-Interactions Involving the MJO (KWs) on Rainfall Variability over SCS and MC in December 2016 (2017)

Chen et al. 2019 J. Climate (in review)





(b) Total (shading) and KW filtered OLR (contour) in 5°S-5°N

Fig. 3



Fig. 4



Fig. 5



# Scale interactions

☆I1 TC Vardah in ER1 enhanced southwesterly winds and coastal convection over W Sumatra → KW2 enhanced

 $\Rightarrow$ I2 MT1 over SCS → Borneo Vortex + KW2 equatorial westerly → TD

ArrowI3 KW2 + WP trade wind + ER2 → MT4/TC Nock-ten

☆I4 MT3 → enhanced westerlies and convection over W Sumatra → terrain effect of Sumatra → KW3 enhanced





# Australian summer monsoon indices: (OLR, U850) (Hung and Yanai 2004); U850 (Kajikawa et al. 2010)



(2°S-11°S, 115°E-150°E; NAU) the first day when U>2 ms<sup>-1</sup> (westerly) and lasts 10 days or more and the OLR<210 Wm<sup>-2</sup> for several days within the 10 days.

U850 (5°S-15°S, 110°E-130°E). the first day after 1 November and (1) U850 during the 5 days following the onset day >0 ms<sup>-1</sup>. (2) In the following four pentads, U850 must be positive in at least three pentads. (3) U850 >1 ms<sup>-1</sup> for the accumulative four-pentad mean.U Fig. 9

### The Influences of Scale-Interactions Involving Convectively-Coupled Kelvin Waves on Rainfall Variability over SCS and MC in December 2016

1. Three KWs were observed while the MJO was inactive.

Four interaction events between the KWs and the westward-propagating waves over the offequatorial regions were examined: two events led to KW enhancements and the other two contributed to the formation of tropical depression/tropical cyclone.

2. Over the KW convectiv region of MC, the diurnal cycle of precipitation was enhanced in major islands and neighboring oceans.

Over land, the DC hotspots were modulated depending on the background winds & terrain effects. Over ocean, the "coastal regime" of diurnal cycle appeared at specific coastal areas.

3. The Australian Summer Monsoon onset occurred with the passage of a KW, which provided favorable conditions of low-level westerlies and initial convection over southern MC and the Arafura Sea. This effect may be helped by the warm sea-surface temperature anomalies associated with the La Niña condition of this month.

4. The current results indicate the importance of scale-interactions of the KWs with respective to synoptic disturbances, diurnal cycle, and Australian monsoon onset, at least when MJO is absent.



